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CARBON RELEASE FROM WOODY PARTS OF TREES ALONG AN ELEVATION GRADIENT IN A TROPICAL MONTANE MOIST FOREST OF SOUTHERN ECUADOR

Dissertation

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Chapter

INTRODUCTION

1.1 General Introduction

With respect to carbon (C) fluxes, tropical montane forests (TMF) are among the least studied terrestrial ecosystems (Brujinzeel and Veeneklaas 1998, Clark 2007). Covering an area of 3.3 million km² globally or of 21.2% of the tropical forests worldwide (Bubb et al. 2004), TMFs extent along a large altitudinal zone (Jacobs 1988). The definition of the lowland border of montane forests is vague and can vary between 750 and 1650 m elevation (Jacobs 1988). Low canopy height, a high abundance of epiphytes and reduced amount of woody climbers usually distinguish TMFs from tropical lowlands. TMFs are characterized by high species diversity harbouring high numbers of endemic plants, which holds especially true for the tropical Andes (Brummit and Lughadha 2003). At higher altitudes, TMFs change into tropical montane cloud forests (TMCF) with their short-statured elfin forest vegetation. Further upslope, TMCFs are often displaced by sparse and open páramo vegetation as the upper limit.

Due to their large altitudinal extension, TMFs occur within a wide range of humidity and temperature regimes. However, all TMFs are characterized by prevailing high atmospheric humidity levels and a frequent cloud cover (Brujinzeel and Veeneklaas 1998). With regard to structure and functionality, TMFs differ considerably from tropical lowland forests. Most obvious is the marked decline in the aboveground biomass. The lower aboveground productivity of TMFs compared to the tropical lowlands has frequently been related to the lower radiation input, the lower air temperatures and the nutrient poor soils (Tanner 1985, Brujinzeel and Veeneklaas 1998, Leuschner et al. 2007, Soethe et al. 2007, Moser 2008). However, in contrast to tropical lowlands, but very similar to high latitude forests, a substantial carbon allocation to the root system was found in TMFs with increasing elevation (Brujinzeel and Veeneklaas 1998, Malhi et al. 1999, Moser et al. 2008). Our understanding of this astonishing carbon shift in TMFs is still incomplete (Brujinzeel and Veeneklaas 1998).

A comprehensive knowledge of the carbon allocation patterns in TMFs can not be achieved without information on the two key processes controlling productivity: the assimilatory CO_2 uptake and the respiratory CO_2 release of the forest ecosystem. Indirect information about assimilatory capacity of TMFs is implied from more readily available forest inventories with data on leaf area and productivity rates. Data on respiratory rates are very scarce and for TMFs almost non-existing. The study of CO_2 release from above- and belowground woody organs provides a convenient framework to assess the impact of altitude and climatic fluctuations on the above- and belowground productivity of TMFs.

Thesis aims & structure

The main objective of the present study was to quantify CO₂ losses from above- and belowground woody organs (stems, coarse roots) of representative tropical montane tree species in southern Ecuador. The study was conducted along a 2000-m elevation transect and aimed to determine the impact of altitude and of seasonal climate variations on patters of CO₂ release from functionally different woody organs of trees. The results will contribute to the understanding of structure and functionality of TMFs as well as provide valuable data for modeling the C dynamics of tropical ecosystems.

The following hypotheses have been tested in the context of this study:

(i) Apparent CO_2 release from stems increases with increasing elevation from 1050 m to 3050 m.

Previous studies found that the aboveground biomass production declined by a factor of 20 from 1050 m to 3050 m. At the same time, it is assumed that photosynthesis does not decline proportionally. The question arises, if respiratory CO_2 release is in fact related to the decreasing temperature with increasing elevation, or if increasing abiotic stress causes compensatory CO_2 release from woody tissue at higher altitudes.

(ii) Apparent CO_2 release from coarse roots increases with increasing elevation from 1050 m to 3050 m.

The 5-fold increase in coarse root biomass production from 1050 m to 3050 m reported in previous studies represents an enormous shift in carbon allocation. This shift should be reflected in the overall respiratory activity of the coarse root system.

(iii) Apparent CO_2 release rates from stems and coarse roots of TMFs at different elevation levels remain constant through the year and do not show seasonality.

Based on the general assumption of a strong temperature-dependence of respiration, an increase of respiratory activity with increasing ambient temperature should be expected. However, mean monthly temperatures fluctuate little in the study area. Additionally, tissue temperature of wood usually varies much less than air temperature.

(iv) With increasing elevation, woody tissue respiration shows an increasing relevance in the carbon balance of TMFs. At high elevations, maintenance respiration from stems and growth respiration from roots are predominant. Based on hypothesis (i) and (ii), it is expected that the shift in biomass between above and belowground components is reflected in the respiration rates at the stand level and in the carbon balance of the TMFs.

The present study was embedded in the DFG- research unit FOR 402 "Functionality in a tropical mountain rainforest: Diversity, dynamic processes and utilization potentials under ecosystem perspectives" (www.bergregenwald.de). As part of an interdisciplinary research project, the present study was closely linked to previous investigations conducted at the very same study sites. Data about above- and belowground biomass stocks and biomass productivity as well as detailed forest inventories (Moser 2008) could be used to estimate the stand level carbon losses from woody tree organs. Information about soil and fine root respiratory activity (Iost 2007) finally helped to derive first C-balances along the tropical montane forest gradient in southern Ecuador.

This work consists of six main chapters: Chapter 1 provides general information about plant respiration and forest gas exchange and gives an introduction of study area and sites.

Chapter 2 presents data on the variability in stem and coarse root respiration patterns with respect to altitude. Changes in the relationship between above- and belowground respiratory activity with increasing elevation are discussed.

Chapter 3 analyzes in further details the temperature response of stem respiration of montane tree species at 1890 m elevation under dry and wet season conditions. Unlike commonly expected, the changes in respiration could not be explained by differences in the thermal regime of dry and wet season. This chapter points out that respiration is driven by more factors than temperature alone and that diverging adaptation potentials of species might play an important role in explaining respiratory reaction to abiotic fluctuations in a highly diverse ecosystem.

Chapter 4 covers a whole annual cycle of respiration measurements along the elevation gradient. A clear trend towards higher CO_2 release under warmer and drier climate conditions is shown for stems, but not for coarse roots. Changes in growth and maintenance respiration of stems and roots with elevation are addressed and first estimates on the annual stand carbon release of stem wood and roots are presented. Chapter 2-4 are formatted for manuscript submission.

Chapter 5 provides a preliminary C balance of the three study sites along the elevation gradient. Estimates are based on own measurements, literature data and previous forest inventories and soil respiration data of the three study sites. Chapter 6 summarizes the presented results and provides an overall discussion of hypotheses.

1.2 Plant respiration & Components

1.2.1 Respiration in plants

Although of equal importance as photosynthetic CO_2 assimilation, the respiratory CO_2 release has received much less attention in the attempt to determine ecosystem productivity (Chamber et al. 2004, Trumbore 2006). Photosynthesis is a distinct, light-dependent process and restricted to chloroplast-containing plant organs. Respiration, in contrast, integrates various disparate components; it is an omnipresent process, which never stops.

The respiratory activity of plants mainly comprises mitochondrial (dark-) respiration, and photo-(light-) respiration; but also futile cycles such as the alternative pathway (cyanideresistant-oxidase or alternative-oxidase). In contrast to mitochondrial respiration, photorespiration is a very specific process, which is closely linked to photosynthesis. Therefore, photorespiration is restricted to chloroplast-containing tissue; it depends on Rubisco, does not provide energy delivering products and it ceases by night. This study is dedicated to mitochondrial respiration, which in the following is referred to as respiration.

At the biochemical level, respiration is defined as the CO_2 release or O_2 uptake associated with the activity of glycolysis, the oxidative pentose phosphate pathway and the tricarboxylic acid (TCA) cycle. The substrate of respiration is mainly glucose and energy equivalents (ADP, Pi, NAD(P)⁺); the glycolysis may also start with other carbohydrates, fat or sugar alcohols (Amthor 1994, Lambers et al. 2005). The product of respiration is energy (ATP, NAD(P)H) to sustain plant life and growth. Plant respiration is thought to be controlled by three processes: substrate availability, demand for energy and the potential enzyme capacity (Amthor 1995, Lambers et al. 1998, Atkin et al. 2005).

However, plants sometimes respire using the alternative oxidase, which converts most of the redox energy into heat, resulting in a much lower energy gain, but higher O_2 consumption and CO_2 loss than the mitochondrial respiration (Larcher 1998). Stress, injuries or senescence of tissue is supposed to stimulate alternative respiration pathways, though the significance and underlying mechanisms remain unclear (Larcher 1998, Lambers et al. 1998, Lambers et al. 2005).

Plant respiration can consume between 25 and 80% of the daily assimilated carbohydrates and hence represents an important factor in the plants' carbon balance (Amthor 2000, Lambers et al. 2005).

1.2.2 Growth & maintenance respiration

Over the past decades, respiration of plants has been partitioned into two major components: growth and maintenance respiration. Maintenance respiration (R_m) is dedicated to the plants' processes for keeping the existing tissue and growth respiration (R_g) to the processes for constructing new tissue (Amthor 1991). R_m is supposed to be strongly correlated to the protein content of the plant. Since protein turnover and enzymatic activity is highly temperature dependent R_m is thought to be more temperature-sensitive in contrast to the temperature-independent R_g (Penning de Vries 1975a). Traditionally, the costs for new tissue are calculated straightforward from tissue composition and the energy required for constructing the components (Penning de Vries 1975b). Estimating maintenance requirements is less clearly quantifiable and assumed to be mainly a function of temperature (Lavigne 1987), nitrogen content and the volume of the tissue (phloem and xylem) it supports (Ryan et al. 1994).

However, the simplifications in the growth-maintenance-paradigm such as the assumptions of constant growth rates, invariable costs of tissue or the N-dependence of R_m have been frequently criticized (Van der Werf et al. 1992, Ryan et al. 1995, Lambers et al. 1998, Amthor 2000, Cannell and Thornley 2000, Thornley and Cannell 2000). Furthermore, it is impossible to strictly separate the various energy demanding processes between the two components. In this regard, the different processes, which comprise respiration (e.g. growth, nutrient uptake, phloem loading, symbiotic N_2 fixation) and the way, these processes are partitioned among the two respiration components still require a more thorough understanding (Cannell and Thornley 2000).

1.2.3 Respiration & temperature

Temperature has long been known to directly influence respiration (e.g. Amthor 1989). This influence was assumed to be exponential with a constant Q_{10} of respiration ranging between 2.0 - 2.3 (e.g. Amthor 1989, Ryan 1991, Raich and Schlesinger 1992). The Q_{10} of respiration is the factor expressing the proportional change in respiratory CO₂ release per 10K rise in temperature and is calculated from the formula (e.g. Atkin et al. 2005):

(1) $Q_{10} = e^{10k}$

where k is the temperature coefficient derived from non-linear regression of respiration plotted against temperature (or, alternatively, by linear regression of log-transformed respiration plotted against temperature). Other studies have used the Arrhenius theory to describe the response of plant respiration to temperature fluctuations (e.g. Crawford and Palin 1981, Lloyd and Taylor 1994, Griffin et al. 2002). Here, log-transformed rates of respiration are plotted against the reciprocal of temperature (1/T). In case of an exponential temperature-respiration relationship, the regression should be strongly linear. The slope of the line is expressed as E_a/R_g , where R_g is the universal gas constant (8.314 J mol⁻¹ K⁻¹) and E_a is the activation energy (J mol⁻¹) for the reaction (Forward 1960). The Arrhenius model is commonly used in physical chemistry to calculate standard references of reactions. However, with regard to plant respiration this theory incorporates a series of assumptions, which were usually not matched by a complex physiological process such as respiration. Among these assumptions, the most critical are the substrate saturation of the reaction, which rarely happen in intact tissue and the implication that a single value of E_a fits to the various specific reactions respiration is composed of (Atkin et al. 2005).

Strong respiration-temperature relationships with constant Q_{10} values were found under certain controlled conditions (e.g. Maier et al. 1998, Tjoelker et al. 1999, Atkin et al. 2000, Tjoelker et al. 2001, Loveys et al. 2003, Armstrong et al. 2006, Atkin et al. 2006, Atkin et al. 2007), while in the field, plant respiration responded neither constant nor necessarily exponential to changes in temperature (e.g. Edwards and McLaughlin 1978, Negisi 1982, Lavigne 1987, Gunderson et al. 2000, Teskey and McGuire 2002). Until now, neither the Q_{10} model nor the Arrhenius theory has been able to describe the highly variable response of respiration to changes in temperature under field conditions (Tjoelker et al. 2001, Atkin and Tjoelker 2003, Atkin et al. 2005). The response of plant respiration to changes in environmental conditions may be especially complex and far-reaching with respect to highlydiverse tropical ecosystems. In this context, the impacts of global climate change on plant respiratory activity urgently warrant further investigation.

1.3 Forest – Atmosphere gas exchange

1.3.1 Forest ecosystems & CO₂ exchange

Quantifying the carbon exchange between forests and the atmosphere is a major topic of active research due to the key role terrestrial ecosystems play for the regulation of the anthropogenic increase in atmospheric CO₂ and the associated climatic changes. The CO₂ exchange of forest ecosystems is driven by two large fluxes, the photosynthetic carbon dioxide (CO₂) assimilation of green tissue and the respiratory CO₂ release of autotrophic and heterotrophic organisms. Carbon balances commonly derived from two reversely operating approaches – the so called top-down or bottom-up models. While the former is based on calculations of net atmospheric CO₂ fluxes, the latter comprises forest or land-use change inventories. Since the processes and carbon pools included considerably differ among both models, results upon the world's carbon balance differ as well (Malhi et al. 1999).

Estimates on the global carbon budget revealed a gross fixation of 90 - 130 Pg C yr⁻¹ (Bolin and Fung 1992), while soil respiration and the CO₂ efflux from terrestrial plants released 64 – 72 Pg C yr⁻¹ (Raich and Schlesinger 1992) and 40 – 60 Pg C yr⁻¹ (Bolin and Fung 1992), respectively, back to the atmosphere. Compared with these large fluxes, the contribution of anthropogenic disturbance to the atmospheric CO₂ concentration is a relatively small value (fossil fuel combustion 5.7 Gt; cement manufacturing 0.01 Gt; Land-use changes 3.5 Gt, http://www.whole-systems.org). Even small changes in one of the two major fluxes of carbon fixation or respiration could alter the accumulation of CO₂ in the atmosphere to a much largely extent than small changes in the anthropogenic disturbance would do (Malhi and Grace 2000). Consequently, it is especially challenging to reduce the large uncertainties, which still exist concerning the magnitude of ecosystem fluxes and the fluxes from single fractions the ecosystem is composed of.

The tropical forest biome, which encompasses moist equatorial evergreen lowlands, moist deciduous, dry deciduous, montane forests and woody savannas accounts for 50% of the global forest area (FAO 1993). In total, these biomes are estimated to annually sequester 42 Pg C in biomass, which equals c. 70% of global terrestrial photosynthesis. However, terrestrial carbon emission and carbon sink terms have often been estimated as a residue of the other components comprising the global carbon budget (Malhi and Grace 2000). Applied ecological research increasingly focused on the quantification of total ecosystem CO₂ release

 $1Pg (1Petagram) = 10^{15}g = 1Gt (Gigatonne)$

rates, which determine net ecosystem production (NEP). Consequently, estimating ecosystem respiration is an important item to identify the carbon source or sink strength of different terrestrial biomes (Malhi et al. 1999, Malhi and Grace 2000, Luyssaert et al. 2007).

1.3.2 Measuring forest CO₂ exchange

Principally, two technical approaches are considered for determining the amount of CO_2 exchange: the eddy covariance technique and chamber measurements. The recently developed eddy covariance (EC) method allows for a comprehensive measurement of CO_2 exchange between large terrestrial areas and the atmosphere. EC is a high-technology, equipment-intensive method, which facilitates the quantification of net CO_2 fluxes entering and leaving a system over temporally large scales (hourly to annually); highly elaborate, only few studies have been conducted in tropical regions. In addition, for capturing atmospheric CO_2 this technique requires sustained high atmospheric turbulences within the system. An accumulation of CO_2 below the EC measurement height due to low turbulences has often resulted in underestimates of ecosystem respiration especially during night-time, when airstreams cease. This factor particularly constraints night-time measurements in tropical forests (Grace et al. 1995, Malhi et al. 1999, Chambers et al. 2004).

While the EC technique provides valuable information about net exchange rates of the ecosystem as a whole, chamber measurements on a spatially limited area allows for the investigation of individual respiratory sources and hence, for a more physiological interpretation of carbon fluxes (Amthor 2000). Generally, the impacts of important environmental factors (e.g., temperature, humidity, radiation) can be deduced from net changes in the entire flux, but it remains unknown, which ecosystem component is the responsive one. Underlying mechanisms of flux dynamics and its driving forces can be better understood by breaking the net flux down into flux subcomponents. Within a forest ecosystem such components could comprise the two large ecosystem fluxes (heterotrophic and autotrophic), functional groups (e.g., understory, trees, palms, lianas) or single plant organs (e.g., leaves, branches, stems, roots). Information on the CO_2 exchange patterns of various components is crucial if we want to gain deeper insight into the physiologically meaningful processes within a plant community and its plastic response to changing environmental conditions.

1.3.3 CO₂ release from woody organs

The respiration of woody parts of trees (branches, stems, coarse roots) plays a significant role for the carbon balance of forest ecosystems, since wood volume comprises a large fraction of the total biomass and considerably increases when trees age (Ryan et al. 1994, Carey et al. 1997). In mature evergreen tropical and subtropical forests, 80 - 90% of the dry biomass is wood (Larcher 1998). Although constituting a substantial fraction of total biomass, the aboveground woody organs respire comparatively less than other plant tissues (e.g., fine roots and leaves), when based on the dry mass (Ryan et al. 1994). This is mainly a result of the large proportion of dead tissue dominating the stem biomass, whereas the respiratory active cambial tissue is restricted to the stem outer ring under the phloem.

Traditionally, three major problems of measuring woody tissue respiration are commonly discussed in the literature and are summarized according to Sprugel et al. (1995):

- CO₂ produced by respiratory activity in the sapwood may be carried away with the xylem sap flow, which, in turn, would result in an underestimation of the actual CO₂ release at the measured section.
- CO₂ respired by roots or taken up from the soil may be transported upwards with the xylem sap flow and additionally released somewhere further up in the stem. This would result in an overestimation of the actual CO₂ release at the measured section.
- Photosynthetically active tissue in the bark may re-assimilate the CO₂ produced by respiration resulting in an underestimation of the actual CO₂ release rate at the measured section.

Nevertheless, independent of knowing exact sources, measuring the apparent CO_2 release of woody tissue *in situ* is an invaluable tool to gain information about the contribution of wood to the overall carbon balance. Early studies in tropical forests estimated the fraction of wood CO_2 release to range between 23 - 50% of total plant respiration (Müller and Nielson 1965, Yoda 1967, Whitmore 1984). However, these measurements were made on excised plant organs and wound respiration most likely contributed a considerable amount to the respiratory CO_2 release (Ryan et al. 1994). More recent estimates of woody tissue respiration in the tropics ranged between 10-13% of gross production (Ryan et al. 1994, Meir and Grace 2002), though only few studies exist about the respiratory CO_2 release from tropical tree stems. Even less information is available about the CO_2 release of branches and coarse roots in tropical ecosystems. Small diameter wood of <10 cm (including branches) contributed 70%

to total wood respiration in an evergreen tropical lowland forest in Costa Rica (Cavaleri et al. 2006). Root respiration was found to strongly increase with decreasing diameter for *Pinus radiata* and *Pseudotsuga menziesii*. Further, seedling roots respired much faster than the roots of mature trees in *Pinus radiata* and *Pseudotsuga menziesii* (Hollinger et al. 1994).

1.4 Study sites

The three study sites are located within the Podocarpus National Park at the eastern slopes of the southern Andes of Ecuador (Figure **1.1**). The park is situated between the two provinces Loja and Zamora covering an area of 1460 km² (Calderón 2002).



Figure 1.1. Study area and location of the three study sites: 1 Bombuscaro (premontane) at 1050m asl, 2 ECSF (lower montane) at 1890 m asl, and 3 Cajanuma (upper montane) at 3050m asl. The whole area covered by the research station (ECSF) is also indicated.

At its northern limit, in the valley of the Rio San Francisco, the research station "Reserva San Francisco" (RSF, formerly ECSF) borders the National Park. The Reserva, covering an area of about 1000 ha, is owned by the foundation Nature and Culture International (NCI), Ecuador and is rented by the DFG for the ongoing investigations of the research unit FOR 402. Emerging between the humid Amazon basin and the dry inter-Andean valley, the area encloses various tropical montane forest ecosystems (Richter 2003). The

montane vegetation includes premontane tropical forest at its lower limit (c. 1000 m elevation) up to upper montane cloud forest (TMCF) and the páramo. In this area, the timberline is reached at 3100 m a.s.l. (\pm 200 m) (Richter 2003).

The vegetation is mainly evergreen broad-leaved. Few deciduous tree species (e.g. *Tabebuia chrysantha*, *Cedrela montana*) and one conifer species (*Podocarpus oleifolia*) occur in the study area (Homeier 2004).

The climate of the study area is strongly influenced by the extremely patchy topography of the Andean mountains. The rainfall gradient along the altitudinal extension of the Reserva is steep and amounts of annual precipitation increase tremendously from c. 2000 mm at 960 m a.s.l. (Zamora) to c. 7780 mm at 3185 m a.s.l. (Cerro del Consuelo, Las Antenas). Fog water input plays a minor role at 1800 m, but increases to 30% of total water input at Cerro del Consuelo (Fabian et al. 2005). The high cloud frequency at the eastern Andean slopes causes a humid climate throughout the year with peak rainfalls between May and July. The prevailing humid climate is interrupted by a short, less humid period between November and February, when westerly foehn winds induce dry and sunny weather conditions (Bendix and Lauer 1992). Mean annual temperature is 19.4 °C at 1000 m and decreases by 0.59K 100 m⁻¹ (Richer 2003). Mean monthly temperatures are more or less constant throughout the year; the warmest months November and the coolest months April differ by 1.9 - 2.4K (Röderstein et al. 2005). However, daily minimum and maximum temperatures during the less humid season can differ by more than 20K (Moser 2008). Temperature never reaches 0 °C throughout the study area.

The soils of the study area developed on metamorphic shale, quartzite or sandstone bedrock. According to FAO taxonomy, soils at 1000 m are classified as alumic Acrisols, at 1900 m as gleyic Cambisols and as Podzols at 3000 m (Iost 2007). At 3050 m, the organic layer is markedly thick (about 430 mm), whereas at 1050 m only the upper 50 mm are organic material. Characteristically for the soils of the study sites are the generally low cation exchange capacity and the low pH (2.9 - 3.9) resulting in an overall low fertility (Iost 2007).

The three sites chosen for this particular study are distributed along a 2000-m elevation gradient and are located with a distance of 30 km between stands (Figure 1.1). All sites are on gentle slopes $(26-31^{\circ})$ facing north-east or north-west (Moser 2008). The low-elevation site (1050 m, S 04°06`54``/ W 78°58`02``) is close to the north-eastern entrance of the Podocarpus National Park (Bombuscaro section) in the province of Zamora-Chinchipe (Figure 1.1). This premontane forest is dominated by tree species from the families

Myrtaceae, Sapotaceae (mainly *Pouteria*), Annonaceae (*Guatteria*), Moraceae (*Ficus*) and Mimosaceae (*Inga*) (Figure **1.2**). Trees are taller than further upslope, reaching a canopy heights of 31.8 m, with maximum stem length of 39.7 m (mean: 15.6 m \pm 0.7 SE, Moser 2008).

The mid-elevation site (1890 m, S $03^{\circ}58^{\circ}345^{\circ}$ / W $79^{\circ}04^{\circ}648^{\circ}$) is close to the Reserva San Francisco (RSF), 30 km from Loja on the road to Zamora, Province of Zamora-Chinchipe (Figure 1.1). In this lower montane vegetation the tree families Melastomataceae (mainly *Graffenrieda emarginata* and *Miconia*), Lauraceae (*Ocotea*), Euphorbiaceae (*Alchornea*), Rubiaceae (*Palicourea*) and Chletraceae (*Chlethra*) are most abundant (Figure 1.2). Canopy height is 18.9 m, with a maximum stem length of 24.8 m (mean: 10.1 m \pm 0.4 SE, Moser 2008).

The upper site (3050 m, S 04°06`711``/ W 79°10`581``) was situated in the the northwestern section of the Park, Cajanuma, Province of Loja (Figure **1.1**). This upper montane vegetation is charcterized by trees from the families Cunnoniaceae (*Weinmannia*), Rubiaceae (*Faramea*), Clusiaceae (*Clusia*), Ericaceae and Symplocacaceae (Figure **1.2**). Canopy height is 9.0 m, with a maximum stem length of 19.2 m (mean: 5.2 m \pm 0.3 SE, Moser 2008).



Figure 1.2. The three study sites at 1050 m (premontane), 1890 m (lower montane) and 3050 m (upper montane). Views from outside (above) and inside (below) the sites.

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Chapter 2

ELEVATIONAL CHANGES IN WOODY TISSUE CO₂ EFFLUX IN A TROPICAL MOUNTAIN RAIN FOREST IN SOUTHERN ECUADOR

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2.1 Abstract

Much uncertainty exists about the magnitude of woody tissue respiration and its environmental control in highly diverse tropical moist forests. In a tropical mountain rainforest in southern Ecuador, we measured the apparent diurnal gas exchange of stems and coarse roots (d: 1-4 cm) of trees from representative families along an elevational transect with plots at 1050, 1890 and 3050 m a. s. l.. Mean air temperatures were 20.8, 17.2 and 10.6 °C, respectively. Stem and root CO₂ efflux of 13 to 21 tree individuals per stand from dominant families were investigated with an open gas exchange system while stand microclimate was continuously monitored. Substantial variation in respiratory activity among different species and different tree individuals was found at all sites. Mean daily CO₂ release rates from stems (R_S) declined 6.6-fold from 1.38 μ mol m⁻² s⁻¹ at 1050 m to 0.21 μ mol m⁻² s⁻¹ at 3050 m. Mean daily CO_2 release from coarse roots (R_R) showed a decreasing tendency from 0.35 to 0.20 μ mol m⁻² s⁻¹ with altitude, but the differences were not significant. There was, thus, a remarkable shift from a relatively high respiratory activity of stems compared to coarse roots at lower montane elevation to an apparent equivalence of stem and coarse root CO₂ efflux rates at high elevation occurred. We conclude that stem respiration, but not root respiration, greatly decreases with elevation in this transect, coinciding with a substantial decrease in relative stem diameter increment and a large increase in fine and coarse root biomass production with elevation.

Keywords: altitudinal transect, coarse root respiration, infrared gas analysis, stem respiration, temperature dependence

2.2 Introduction

Plant tissue respiratory activity is thought to consume 30-80% of the daily assimilated carbon gain (Amthor 2000), constituting one of the main sources of CO_2 released to the atmosphere (Trumbore 2006). Although the carbon balance of forests is in the focus of current global change research, the respiration of stems, branches and coarse roots is one of the least studied processes (Sprugel and Benecke 1991). Knowledge of tropical forest respiration is particularly sparse, despite their acknowledged importance of tropical forests in the global carbon balance (Meir and Grace 2002, Chambers et al. 2004). For reliable modeling of the carbon sink strength of tropical forests in a changing climate, a detailed knowledge of plant respiration is needed, particularly its variability among forest types, and its dependence on the environment. The few *in situ* measurements of respiration of tropical forests trees indicate that woody tissue respiration accounted for 10-13% of gross photosynthesis (Ryan et al. 1994, Meir and Grace 2002). Earlier measurements in tropical forests based on observations of CO_2 release from excised plant organs yielded values between 23 and 50% (Müller and Nielson 1965, Yoda 1967, Whitmore 1984).

To our knowledge, only one gas exchange study has been conducted in tropical highelevation forests until now: Cavieres et al. (2000) measured leaf gas exchange of two tree species in the Venezuelan Andes. Studies quantifying woody tissue respiration along altitudinal transects in tropical mountain forests are lacking. Information from such studies would help to predict effects of temperature change on plant respiration in tropical ecosystems. Woody tissue CO₂ release rates can vary enormously, not only among different forest types (Lavigne et al. 1996, Ryan et al. 1997), but also among species within a stand and among individuals of the same species (Meir and Grace 2002). Information on the spatial variability of woody tissue respiration is indispensable when extrapolating gas flux data from tree to stand. This information is particularly important in tropical forests with their high species richness and large structural variability across environmental gradients (Meir and Grace 2002, Chambers et al. 2004).

The current study was undertaken to: (1) quantify species-specific differences in woody tissue respiration in tropical mountain forests; (2) compare the respiratory activity of stems and coarse roots; and (3) analyze changes in stem and root respiration along an altitudinal span of 2000 m in a tropical mountain rainforest in southern Ecuador.

2.3 Materials & Methods

2.3.1 Study sites

The study was carried out in Podocarpus National Park (PNP) in the surroundings of Loja on the eastern slopes of the southern Ecuadorian Andes. We chose three forest stands along an altitudinal gradient ranging from 1050 m to 3050 m a. s. l. The maximum distance between the stands was about 30 km. The low-elevation stand (1050 m, S 04°06`54``/ W 78°58`02``) is located in the northeastern part of PNP (Bombuscaro section) in the Province of Zamora-Chinchipe. The mid-elevation stand (1890 m, S 03°58`345``/ W 79°04`648``) is close to the Estacion Cientifica San Francisco (ECSF), 30 km from Loja on the road to Zamora, Province of Zamora-Chinchipe. The high-elevation stand (3050 m, S 04°06`711``/ W 79°10`581``) is in the Cajanuma area in the northwestern part of PNP, Province of Loja. All stands were selected on gentle slopes (26-31°) facing northeast to northwest, covering an area of 20 x 20 m.

The climate of the area is mainly influenced by easterly winds that bring frequent rainfall throughout the year with peaks from May to July. During our study, conducted from October to December 2005, westerly winds strongly influenced the local weather causing a relatively dry and sunny period in the study area.

The soils of the area developed either from grandiosities (low-elevation stand), or metamorphic shale, quartzite and sandstone bedrock (mid- and high-elevation stands). Throughout the study region, the soils are relatively infertile (Schrumpf et al. 2001).

Forest structure and selection of tree individuals

The stands were selected to: a) be representative of the vegetation type at each elevation; b) have a closed canopy within a surrounding area of 100 x 100 m; and c) be free of recent anthropogenic influence or landslide disturbance. The low-elevation stand (1050 m) represents the transitional zone between tropical lowland and lower montane rainforest. The mid-elevation stand (1890 m) is a typical lower montane rain forest, and the high-elevation stand (3050 m) is located close to the timberline and represents a typical "elfin forest" characterized by stunted trees with warped stem forms. Further details on forest structure are given in Table **2.1**. All plots have been previously studied and described by Röderstein et al. (2005), Leuschner et al. (2007) and Moser et al. (2008).

Table 2.1. Climate and stand structure characteristics of the study sites at 1050, 1890 and 3050 m (data from Moser et al. 2008, means with standard deviation in parenthesis). Mean air temperature and mean relative humidity were recorded inside the stands at a height of 2 m. Rainfall data are extrapolated from measurements in gaps at about 1050 m (G. Moser, in press), 1950 m and 3170 m (Emck 2007). Abbreviations: AGB, aboveground biomass; BGB, belowground biomass.

1050 m	1890 m	3050 m
20.8 (3.3)	16.8 (4.4)	10.6 (3.1)
87.3 (16.1)	87.4 (21.9)	91.0 (13.0)
2230	1950	4500
17.3 (1.3)	12.2 (0.8)	7.2 (0.4)
15.6 (0.7)	10.1 (0.4)	5.2 (0.3)
33.6	36.9	42.2
968	2333	8317
6.0 (0.4)	5.7 (0.5)	2.2 (0.2)
0.64 (0.15)	0.60 (0.16)	0.69 (0.15)
6.82 (0.44)	9.74 (0.83)	3.64 (0.29)
285.1	173.0	112.2
32.1	25.8	62.7
	1050 m 20.8 (3.3) 87.3 (16.1) 2230 17.3 (1.3) 15.6 (0.7) 33.6 968 6.0 (0.4) 0.64 (0.15) 6.82 (0.44) 285.1 32.1	1050 m $1890 m$ $20.8 (3.3)$ $16.8 (4.4)$ $87.3 (16.1)$ $87.4 (21.9)$ 2230 1950 $17.3 (1.3)$ $12.2 (0.8)$ $15.6 (0.7)$ $10.1 (0.4)$ 33.6 36.9 968 2333 $6.0 (0.4)$ $5.7 (0.5)$ $0.64 (0.15)$ $0.60 (0.16)$ $6.82 (0.44)$ $9.74 (0.83)$ 285.1 173.0 32.1 25.8

In each stand, a minimum of 80 canopy trees were identified at the species level. As long as they reached the canopy, trees smaller than 5 cm DBH (diameter at breast height) were included in the samples. To measure CO_2 release rates of woody organs, 13 to 21 trees per stand were selected (Appendix 1). We required that trees belong to abundant families in the particular stand and comprise a broad range of DBH classes in order to represent the floristic composition and size heterogeneity of the stand. Thus, each sample consisted of trees from 10-11 families. At the mid- and low-elevation stands, we included more trees (n = 20 and 21, respectively) to account for the wider diameter range than in the high-elevation stand (n = 13). Diameter of the selected trees ranged from 8.70 to 43.85 cm at 1050 m, from 3.02 to 26.47 cm at 1890 m and from 2.48 to 17.67 cm at 3050 m. We measured 4 - 8 coarse roots (diameter: 1-4 cm) at each site depending on accessibility.

2.3.2 Gas exchange measurements

Woody tissue CO_2 release rates were measured *in situ* between October to December 2005, which corresponds to the drier season of the year, although each month received at least 80 mm of rain. In each stand, the measurements were made over a 10-day period, during which continuous measurements were made of CO_2 efflux from woody organs (stems and coarse roots). Stem CO_2 release was monitored at breast height (1.3 m) using transparent plexiglas chambers (95.1 cm³ volume) tightly fitted onto the bark surface (Appendix 2). When necessary, mosses and lichens were cautiously removed from the measured stem segment using a soft brush, carefully avoiding damage to the bark. Segments of coarse roots were enclosed in transparent plexiglas chambers of 473.8 cm³ volume fixed around the organ with staunching rings and sealed using Terostat® VII (Teroson, Henkel AG, Düsseldorf, Germany). The cylindrical chamber design allows for the measurement of organ sections ranging in diameter from 1 to 4 cm. Both types of chambers have a relatively small volume and are designed with inlet and outlet nozzles at opposite sides in order to ensure adequate mixing of the incoming air. Air-tightness of the measurement chambers was controlled via electronic air flow meters.

The diameter of the stems or roots was measured in the middle of the organ section enclosed in the chamber. The surface temperature of the measured organ section was recorded with thermocouples attached on the outside of the stem surface next to the chamber. We selected coarse roots (d: 1-4 cm) growing a few centimetres beneath the soil surface and uncovered the root section to be measured with a soft brush.

Gas exchange system

Net exchange of CO_2 across stem or root surfaces was measured with a mobile 6chamber respiration system ANARESY 2 (Walz, Effeltrich, Germany; Appendix 2) with an integrated infrared gas analyzer for CO_2 and H_2O (LI-7000, Li-Cor, Inc., Lincoln, NE). The open gas exchange system was operated in differential mode (Ryan et al. 1995) and allows for continuous diurnal measurements of the apparent CO_2 release rate in six plexiglas chambers (Horna and Zimmermann 2000). Buffered incoming air permanently passed all six chambers with a maximum flow rate of $1 1 \text{ min}^{-1}$.

Every 6 min, the system switched automatically from one chamber to the next, thus recording about two CO_2 release values per chamber per hour. The six chambers were moved to different tree individuals after completing a 24-hour measurement cycle. Electrical power

was supplied by a generator that charged car batteries connected in series (24 V DC). The generator was placed at a distance of over 100 m from the measuring system to avoid any influence of the fumes on the measurement. For every 10-day measurement interval, the entire set up was moved to the next site. The weather conditions during the measuring period from October to December 2005 were sufficiently stable to allow for a comparison of the three stands.

Calculation of respiration rates

The LI-7000 infrared gas analyzer continuously determines both the absolute CO_2 concentration ([CO_2]) and the difference between ambient atmospheric CO_2 concentration [CO_2] and the concentration inside the chamber corrected for atmospheric pressure. Air flow rate is expressed as a molar flow rate. The woody tissue CO_2 release rates are then calculated as:

(1)
$$R = D [CO_2] * F/Ac$$

where *R* is the respiration rate in μ mol CO₂ m⁻² s⁻¹, *D* [CO₂] is the difference between ambient (reference gas) and chamber (sample gas) [CO₂] concentration, *F* is the air flow rate (mol s⁻¹) which passes through the chamber, and *Ac* is the surface area (m²) of the enclosed organ segment.

Depending on the diameter of the measured object, CO_2 release rates of woody biomass in tropical forests have been related either to surface area (Levy and Jarvis 1998, Chambers et al. 2004), tissue volume (Ryan et al. 1994), or to a combination of both (Meir and Grace 2002). Recently, Cavaleri et al. (2006) found the proper unit of expression of CO_2 release by tropical woody tissue is dependent on the position within the tree, with canopy rates related to surface area, but efflux rates in the bottom 2 m of the canopy related to both volume and surface area. Nevertheless, since the volume of living tissue in stems and woody roots may be differ considerably among tropical trees of different systematic classifications and ages (Meir and Grace 2002, Chambers et al. 2004), we used the surface area of the measured wood sections as a basis for calculation.

Despite evidence that dissolved CO_2 is transported in substantial quantities in the xylem sap of certain tree species (Levy and Jarvis 1998, Levy et al. 1999, Horna and Zimmermann 2000, Teskey and McGuire 2002, McGuire and Teskey 2002, Gansert and Burgdorf 2005), we ignored this flux in the current measuring program and interpreted our stem and root efflux data as woody tissue respiration rates.

Thermal regimes differ greatly along the altitudinal gradient, with little overlap in temperature between the low- and high-elevation sites. Converting respiration rates to a common temperature (e.g., 15-20 °C) would yield extrapolated release rates beyond naturally given amplitudes and would result in the comparison of efflux rates at lowest night temperatures at 1050 m with rates at the upper daytime temperature limit at 3050 m. Therefore, we decided to underpin our comparison of mean stand respiration rates with results of individual regression analyses rather than adjusting apparent efflux rates to a common temperature.

Stand microclimate

During the 10-day measurement intervals, air temperature and relative air humidity at 2 m height inside the stands were monitored synchronously using a Rotronic sensor (Rotronic AG, Bassersdorf, Switzerland) connected to the data logger of the ANARESY system (CR 10, Campbell Scientific, Logan, UT). Annual means of air temperature and air humidity for each site were computed from daily climate data from instruments located in each stand (1.5 m above ground).

Statistical analysis

Mean stem respiration rates of individual trees were log-transformed for homogeneity of variances before analysis of variance (ANOVA) testing for significant differences between sites (Scheffé test for unbalanced data sets, p < 0.05). Root respiration data matched parametric assumptions without transformation.

Carbon dioxide release rates were regressed against tissue surface temperature with the 45 half-hour respiration rates (0100-2300 h) of all measured stems and roots per site. Additionally, we determined the regression coefficients for each individual tree and root segment and for all stems and roots per site. We ran an ANOVA based on the results of the regression analysis of the various stem and root individuals to test for significant differences in respiration rates at 0 °C (intercept) and responsiveness to temperature (slope) between the three study sites. We used the coefficient of determination (r^2) to quantify the influence of the independent variable (temperature) on the dependent variable (CO₂ efflux).

2.4 Results

Microclimate

The study months October to December 2005 were characterized by westerly foehn winds causing relatively dry weather conditions. In this period, 20-30% less rain fell per month as compared to 2004. As a consequence, relative air humidity (RH) and air temperature (T_a) showed considerable variation during the 10-d measurement periods at all study sites (Table **2.2**).

Table 2.2. Stand mean air temperature (T_a) and relative humidity (RH) during the 10-daymeasurement campaigns at the study sites at 1050, 1890 and 3050 m elevation. Overall ranges are given in parenthesis.

Elevation (m)	$T_a (°C)$	RH (%)
1050 m	20.8 (16.1 - 30.3)	87.1 (39.6 - 99.9)
1890 m	17.2 (9.5 – 25.5)	77.6 (15.1 – 99.9)
3050 m	10.6 (4.5 – 19.9)	91.1 (43.4 - 99.7)

Woody tissue CO₂ efflux along the altitudinal transect

Mean daily CO₂ release rates from stems (R_S) differed significantly between all sites. Values decreased from 1.38 μ mol m⁻² s⁻¹ at 1050 m to 0.21 μ mol m⁻² s⁻¹ at 3050 m, with an average R_S of 0.76 μ mol m⁻² s⁻¹ at 1890 m (Table **2.3**). Thus, average stem respiration declined by a factor of 6.6 over an altitudinal span of 2000 m. Parallel to R_S, mean root CO₂ efflux (R_R) also tended to decrease along the elevational transect (Table **2.3**). Mean R_R decreased by 20% from 1050 to 1890 m and by 29% from 1890 to 3050 m, resulting in a 43% reduction along the whole transect, but differences were not significant.

There was considerable variation in R among the trees within the three stands (Figure **2.1 - 2.3**). The overall range of R values was higher at the lower and mid-elevation stand (two orders of magnitude) than at the upper site (one order of magnitude). Coarse root respiration varied over one order of magnitude within all study sites (data not shown).



Figure 2.1. Diurnal course of stem and coarse root (d: 1-4 cm) respiration together with air temperature (solid line) at the study sites at 1050 m (n=21 stems, n=8 roots), 1890 m (n=20 stems, n=7 roots) and 3050 m (n=13 stems, n=4 roots). Each daily course consists of 45 half-hour measuring points. The 45 half-hour respiration data points are the mean of the diurnal values of 13 to 21 trees or 4 to 8 roots. Error bars equal 1 standard deviation. Air temperature was measured at 2 m height inside the stands. Note the different scales of the y-axis for stem and root respiration.

Woody tissue CO₂ efflux and dependence on temperature

Mean diurnal R_S and R_R exhibited little if any change with temperature over the course of the day in all three stands (Figure 2.1). Despite contrasting diurnal temperature regimes at the lower and the upper montane sites, the slopes of the regression analysis of the integrated data sets of stems and coarse roots were similar and remarkably flat (Figure 2.2). ANOVA conducted for intercept and slope of the regression lines of every individual stem (n= 13-21) and root segment measured (n= 4-8) confirmed that significant differences in temperature responsiveness (coefficient b) for stems and coarse roots between the three study sites did not exist (Table 2.3). Parallel to the apparent respiration rates at ambient temperature, the mean yaxis intercept (coefficient a) of the individual regression lines significantly decreased for stem respiration between 1050 m and 3050 m. For coarse root CO₂ efflux, the intercept (i.e., the respiration rate at 0 °C) also tended to decrease with elevation, but the differences between the sites were not significant.



Figure 2.2. Dependence of the CO₂ efflux rates (R) of all investigated stems (a) and coarse roots (b) on air temperature at the three study sites (• = 1050 m, ∇ = 1890 m and \blacktriangle = 3050 m a.s.l.). Data points are 45-half-hour values of the 13 - 21 different stems or 4 - 8 roots measured during a certain measurement day at ambient temperature. The mean respiration rate at ambient temperature is marked (\Box) for each stands. Note the different scales of *R*-axis in panels a and b.

Although pooling data at stand level resulted in no major differences in temperature responsiveness between sites, regression analysis of individual trees and root segments showed exceptionally high variation in temperature response and respiration rates at 0 °C (intercept; data not shown) between individuals as well as within sites. Furthermore, we found remarkable discrepancies among individuals in the direction and strength of the temperature dependence of R_s and R_R . Highly significant positive as well as negative relationships were found at 1050 m and 1890 m, but also no temperature response was detected (Appendix 3a-b). Only in the high-elevation stand (3050 m) did the regression analysis for individual stems reveal significant positive correlations of R_s and temperature in most cases, but no negative relationships were detected (Appendix 3c). Similarly, R_R at 3050 m mainly showed a significant positive relation to temperature. In contrast at the two lower sites, most of the coarse roots did not show any relationship to temperature. The direction of response was not related to taxonomic groups (family) in any of the stands.



Figure 2.3. Stem CO_2 efflux rate (R) of individuals from various families in the stands at 1050, 1890, and 3050 m elevation. Data are daily means of R from all measured trees, each consisting of 45 half-hourly values during one day. Error bars equal 1 standard deviation. The families included are those with the greatest number of individuals at the particular site.

Abbreviations: Mora1-4 = Ficus sp; Mela1-2 and 7-9 = Miconia punctata; Mela3 = indet.; Sapo 1 = Chrysophyllum sp; 3-2 = Pouteria cf; Anno1-2 = indet.; Laur1-2 = indet.; Mimo1 = Inga sp; Mimo2 = indet.; Aral1 = Schefflera sp; Cecr = Pourouma cf; Euph1 = Alchornea sp; Myri = Virola cf; Myrt = indet.; Mela4-6 = Graffenrieda emarginata; Laur3 = Nectandra sp; Laur4 = Endlicheria oreocola; Laur5 = Ocotea aciphylla; Sapi1-3 = Matayba inelegans; Rubi1 = Palicaurea sp; Rubi2 = Ladenbergia cf oblongifolia; Aqui1 = Ilex cf amboroica; Aral2 = Schefflera sp; Clet = Clethra revoluta cf; Euph2 = Hyeronima morisiana; Myrs1= Myrsine cf; Sapo = Micropholis guyanensis; Clus1 = Clusia sp 1; Clus2 = Clusia sp 2; Eric1 = Cerotostema cf; Eric2 = indet.; Myrs2 = Myrsine sp; Mysr3 = indet.; Aqui2 = Ilex weberlingii; Chlo = Hedyosmum sp; Cuno = Weinmannia loxensis; Mela10 = Axinea sp; Rubi3 = Faramea sp; Styr = Styrax foveolaria; Symp = Symplocos sp.

Table 2. 1050, 18 1050, 18 with line with line parenthe changes roots $(n=breast he parenthe parenthe parenthe$	3. Magazina Magaz A Magazina Mag	ean daily CO ₂ eff nd 3050 m eleva efficients (i.e., a gression analysis Different letters i sal area (incremo er site) were sup of the stems or ro	flux rates (R, μmol m ⁻² tion. Analysis of varia = intercept, b = slope is for each individual st ndicate significant diff ent per existing basal a plied by G. Moser (Un oot diameter (d, cm) an	⁸ s ⁻¹) of stems and coar ance was performed o e of the increase in R tem and root segment. ferences between sites area per year in %) of niversity of Göttingen, id mean diameters of t	se roots at ambient temperature at in daily mean efflux rates and on with temperature) as determined Standard deviations are given in (p < 0.05, Scheffé test). Data on stems (n=80 per site) and coarse Germany). Ranges in diameter at he increment studies are shown in
Site	n	R	а	þ	Basal area change
Stems					
1050 m	21	1.38 (0.88) a	1.2899 (1.2011) a	0.0048 (0.0458) a	2.11 (d: 5.34-69.32, mean: 17.3)
1890 m	20	0.76 (0.52) b	0.6336 (0.6317) ab	0.0076 (0.0286) a	1.34 (d: 4.20-35.26, mean: 12.2)
3050 m	13	0.21 (0.12) c	0.1269 (0.1264) b	0.0084 (0.0072) a	0.47 (d: 2.91-16.47, mean: 7.2)
Roots					
1050 m	8	0.35 (0.23) a	0.3539 (0.2078) a	-0.0030 (0.0160) a	0.40 (d: 3.21-32.23, mean: 7.93)

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1.26 (d: 4.52-10.06, mean: 7.11)

1.92 (d: 3.57- 6.77, mean: 4.56)

0.0148 (0.0124) a

0.0004 (0.0256) a

0.2728 (0.4023) a

0.28 (0.20) a

~

1890 m

0.0434 (0.0781) a

0.20 (0.15) a

4

3050 m
2.5 Discussion

Elevational changes of stem and root respiration – evidence for shifts in the relative importance of root versus shoot growth with altitude

To our knowledge, our respiration data are the first reported values for tropical mountain forests. Along our 2000-m elevational transect from 1050 to 3050 m, mean R_S decreased more than sixfold, whereas R_R did not decrease significantly. Because we investigated a large number of stems from different species representing the most abundant families in the three forest stands, this marked decrease in CO_2 efflux rates must be a general trend across the altitudinal transect. Moreover, the altitudinal change in R_S across the transect coincides with a pronounced shift in the aboveground:belowground biomass ratio from 9:1 at 1050 m to 2:1 at 3050 m (Moser et al. 2008). Therefore, the apparent belowground shift of respiratory activity may reflect differences in the ratio of growth respiration versus maintenance respiration, with a priority of stem growth at lower elevations and a priority of root growth and root activity at higher elevations.

Dendrometric measurements of stem and coarse woody root growth at our study sites (G. Moser, unpublished data, Table **2.3**) revealed a 4.5-fold decrease of relative basal area increment per year for stems from 1050 to 3050 m, counteracted by a 4.8-fold increase in annual coarse root basal area increment along the altitudinal gradient.

The reduction in stem growth and thus the decrease in stem respiratory activity with increasing elevation is most likely a consequence of decreasing nitrogen availability and reduced photosynthetic gain as a result of decreases in leaf area index and foliar nitrogen concentration at high elevations (Leuschner et al. 2007). The observed coarse root biomass increment along the gradient was paralleled by a pronounced increase in fine root biomass (Moser et al. 2008). Increasing allocations of carbon and nutrients to belowground organs is likely an adaptation to decreasing nutrient availability (Bloom et al. 1985). Soethe et al. (2006) concluded that the large investment in the coarse woody root stock at the high-elevation site is an adaptation ensuring tree anchorage on waterlogged and steep slopes at 3050 m elevation.

The lower values of R_R relative to R_S that were observed at all stand elevations contrast with results commonly reported for coniferous forests. In pine forests, for example, coarse root respiration was found to exceed stem and branch CO₂ efflux two- to seven-fold (*Pinus radiata*: Ryan et al. 1996, Ryan et al. 1997) or even up to ten-fold (*Pinus strobus*: Vose and Ryan 2002). However, these data come from very different biomes. We are unaware of a study on stem or root respiration of trees along altitudinal transects in the Tropics with which our data can be compared.

Mean R_s was four times higher than R_R at 1050 m, 2.7 times greater at 1890 m, but about equal at 3050 m: a result of a decrease in R_s while R_R remained almost constant. The similarity in respiration rates of plants growing across a thermal gradient (Körner and Larcher 1988) has been attributed to acclimation (Amthor 1994) or thermal homeostasis (Larigauderie and Körner 1995). Such adaptation allows plants to meet their energy requirements even when temperatures are low. However, the relative constancy of the coarse root respiratory activity, which we observed with increasing altitude, despite a marked decline in stem CO_2 release, is most likely the result of a shift in recourse allocation reflecting the changing environmental conditions. Even though the number of roots measured was quite low, our results closely mirrored the dendrometric studies of Moser et al. (2008) at the same sites, showing an increase in coarse root growth with increasing altitude.

Temperature sensitivity of respiration in tropical mountain forests

By measuring CO_2 efflux in 13 to 21 tree stems and 4 to 8 roots per stand, we demonstrated a substantial within-plot variation in apparent respiratory activity (Figure 2.3). The coefficient of variation in R_S remained constant across the gradient, whereas only the overall range in respiratory release decreased with elevation by two orders of magnitude from 1050 to 1890 m and by one order of magnitude from 1890 to 3050 m.

Similarly, Chambers et al. (2004) reported a variation of two orders of magnitude (0.03 to 3.64 μ mol m⁻² s⁻¹) within a stand. However, they focused on ecosystem exchange rates without differentiating between tree families or species. Ryan et al. (1994) found that stem CO₂ release rates varied sevenfold between two tree species. In forest stands in central Cameroon and the Brazilian Amazon, stem CO₂ release rates from 14 and 13 tree families ranged from 0.2 to 5.2 μ mol m⁻² s⁻¹ and from 0.1 to 3.3 μ mol m⁻² s⁻¹, respectively (Meir and Grace 2002). Cavaleri et al. (2006) monitored different plant functional groups in a tropical lowland rain forest in Costa Rica and found substantial differences in the CO₂ efflux of stems and branches of dicotyledonous tree species, lianas and palms. The comparable variation among families in stem respiration along our altitudinal transect suggests that there is no decrease in tree functional diversity towards the harsher environment close to the tree line.

Homogeneity among sites in the slopes of the respiratory responses to temperature per site indicated that temperature sensitivity did not differ across elevations. Furthermore, the small slopes suggest that temperature responsiveness is low in the trees of this tropical mountain rain forest. By contrast, regressions for individual trees revealed highly variable response patterns within and between sites. The variability in response and the occurrence of inverse temperature responses among coexisting trees could not be explained by taxonomic status (family) or tree size (authors' unpublished observations). Negative respiration-temperature relationships as found at 1050 and 1890 m elevation may be attributable to climatic effects on xylem sap flow which, in turn, affected the amount of CO₂ released through the bark (Gansert and Burgdorf 2005). It is also possible that the exceptionally dry and hot weather during the measurement periods induced changes in respiratory activity. The capacity to acclimate to short- (hours) or medium-term (days) changes, or both, in temperature can vary greatly among individual plant species (Larigauderie and Körner 1995, Atkin et al. 2005). Additionally, plants or organs able to acclimate rapidly by short-term weather fluctuations may even show continuously changing temperature responsiveness as found by Atkin et al. (2000) for leaf dark respiration of *Eucalyptus pauciflora* Sieb. ex Spreng..

Based on absolute values, the mean slopes of the stem and coarse root response to temperature were steeper at the higher elevation site, indicating a more pronounced thermal sensitivity of the cold-grown plants. Higher Q_{10} values in plants grown in cold environments compared with warm-grown ones were found by Tjoelker et al. (2001), whereas Larigauderie and Körner (1995) found that the variability of leaf Q_{10} was unrelated to plant origin. Atkin et al. (2005) concluded that a systematic variation in Q_{10} values does not exist and that differences among contrasting biomes are relative and thus do not reflect an inherent variable temperature responsiveness of characteristic plant species. However, our results showed that the mean temperature sensitivity of stems and coarse roots was similar among the contrasting thermal environments along the gradient, whereas the analyses of individual plant responses showed that large differences in temperature responsiveness exist within the three study sites.

Given the homogeneity of mean slopes across the differing growth environments, any differences in CO_2 release rates across our gradient resulted from differences in the y-axis intercept or respiration rate at 0 °C. Removing temperature as a confounding factor still yielded a marked decrease in stem CO_2 efflux from 1050 to 3050 m, indicating that temperature might not be the primary influencing factor as underpinned by the low temperature sensitivity found along the altitudinal gradient.

Conclusions

Our respiration measurements along a tropical mountain transect indicate a shift from high respiratory activity of stems compared with coarse roots at lower elevation (1050 m) to an apparent equivalence of stem and coarse root CO_2 efflux rates at 3050 m. The observed decrease in the ratio of stem to root efflux rates with altitude may be explained by the substantial decrease in stem growth, while coarse root growth increased with increasing elevation. We found that responses of CO_2 release rates of woody tissue to changes in temperature differed greatly between study sites, as well as among species and plant organs, with the underlying mechanisms remaining unclear. The remarkable variation in respiratory activity and, most importantly, in temperature response of respiration suggests that predictions at the community level or even estimates for entire ecosystems on the basis of few point measurements of selected plant species must be interpreted cautiously. Because there was great variation in acclimation pattern among plants and between sites, questions about the response of plant community CO_2 efflux to climate change should principally be answered at the community level (Larigauderie and Körner 1995).

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Chapter 3

$\label{eq:constraint} \begin{array}{l} \text{Diverging Temperature Response of} \\ \text{Tree Stem CO}_2 \ \text{Efflux to Dry And Wet Season Conditions} \\ \text{In a Tropical Montane Moist Forest} \end{array}$

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SUBMITTED TO: TREES - STRUCTURE AND FUNCTION

3.1 Abstract

It is commonly presumed that plant respiratory CO₂ release increases with increasing temperature. However, we report on very contrasting stem CO₂ release (R_S)-temperature relationships of trees in a species-rich tropical montane forest of southern Ecuador under dry and wet season conditions. Rates of R_S were low and completely uncoupled from the dial temperature regime during the humid season. In contrast, during the dry season, R_s was generally higher and temperature sensitivity of R_s differed greatly in degree and even in the direction of response, indicating that temperature might not be the only determinant of R_S. In order to explain the heterogeneity of R_S, we related R_S to vapour pressure deficit, wind speed and solar radiation as important abiotic drivers influencing transpiration and photosynthesis. Stepwise multiple regression analyses with these meteorological predictors were either biased by high collinearity of the independent variables or could not enhance the ability to explain the variability of R_S. We assume maintenance respiration to dominate under humid conditions unfavourable for energy acquisition of the tree, thus explaining the pronounced uncoupling of R_s from atmospheric parameters. In contrast, the drier and hotter climate of the dry season seems to favour R_S via enhanced assimilatory substrate delivery and stem respiratory activity as well as elevated xylem sap CO₂ imports with increased transpiration. In addition, tree individual differences in the temperature responses of R_S may mirror diverging climatic adaptations of co-existing moist forest tree species which have their distribution centre either at higher or lower elevations.

Keywords: climatic adaptation, Ecuador, stem respiration, temperature sensitivity of respiration, tree species richness

3.2 Introduction

Temperature has long been identified as the most important abiotic factor influencing plant respiratory activity due to its well known affect on enzymatic reactions (e.g. Amthor 1989). A strong relationship between temperature and plant tissue respiration has been exhaustively proved and verified under controlled conditions (e.g. Maier et al. 1998, Tjoelker et al. 1999, Atkin et al. 2000, Tjoelker et al. 2001, Loveys et al. 2003, Armstrong et al. 2006, Atkin et al. 2006, Atkin et al. 2007). However, in the field, a consistent relationship is often not found (e.g. Edwards and McLaughlin 1978, Negisi 1982, Lavigne 1987, Gunderson et al. 2000, Teskey and McGuire 2002). Especially woody tissue respiration is reported not to be exclusively related to temperature since other sources of CO₂ than the respiratory activity of living wood cells (inner bark, cambium, xylem parenchyma) alone can influence the apparent amount of CO₂ escaping through the bark. Imports of dissolved CO₂ via the transpiration stream can add substantial amounts to the stem internal carbon dioxide. The quantitative contribution of the different sources to the internal flux of CO₂ was found to vary considerably on a seasonal as well as daily basis. Consequently, stem internal CO₂ concentrations can constantly fluctuate, which in turn can affect the amount of CO₂ diffusing out of the bark (e.g. McGuire and Teskey 2004, Teskey et al. 2008, Saveyn et al. 2008ab).

Stem respiration usually depends on substrate availability, the demand for energy equivalents and the enzymatic activity (Amthor 1995, Lambers et al. 1998, Pruyn et al. 2002, 2005, Atkin et al. 2005). However, it is still barely understood how stem CO₂ release (R_S) is finally controlled and how internally circulating CO₂ is influencing the apparent efflux. High rates of R_S coincided with high xylem sap fluxes (Levy and Jarvis 1998, Levy et al. 1999, Horna and Zimmermann 2000). Other studies reported a negative correlation between xylem flux and radial CO₂ efflux (Negisi 1979, Edwards and Hanson 1996, Teskey and McGuire 2002, Gansert and Burgdorf 2005, Teskey and McGuire 2007). All these studies suggested that R_S might be linked to canopy water use. However, removing the foliage had little effect on the dial pattern of R_s in an experiment of Maier and Clinton (2006). No relationship between sap flow and R_S was found by Ceschia (2001), Carey et al. (1996) and Edwards and Wullschleger (2000). Studies of xylem sap flow were mainly conducted on coniferous (Carey et al. 1996, Maier and Clinton 2006) or deciduous (Edwards and Hanson 1996, Ceschia 2001, Gansert and Burgdorf 2005, Teskey and McGuire 2007) tree species of temperate climates. Results of one single tree (Ceschia et al. 2006, Gansert and Burgdorf 2005, McGuire and Teskey 2002, Saveyn et al. 2008a) or of few individuals of the same species (McGuire and Teskey 2007, Maier and Clinton 2006, Saveyn et al. 2008b) have been used to generalize trends in the influence of internal CO_2 on apparent CO_2 efflux rates.

However, on the ecosystem level the picture gets more complicated due to the fact that co-occurring plants are responding individualistically to their specific environment and probably even more so to shifts in the environmental setting (Larigauderie and Körner 1995, Arnone and Körner 1997, Amthor 1989, Oren and Pataki 2001, Reich et al. 2003, Atkin et al. 2005, Kerkhoff et al. 2005, Enquist et al. 2007). A better understanding of the variability in R_s among different individuals in a population, different species in a stand, and different patches of the forest is therefore crucial to predict changes in the carbon dioxide exchange between tree wood and the atmosphere, in particular in highly diverse tropical forests (Kerkhoff et al. 2005, Enquist et al. 2007).

In this paper, we report on in situ-measurements of the diurnal rates of R_S of representative tree species in a tropical montane moist forest in southern Ecuador during two hydrologically contrasting seasons of the year. Under the prevailing moist conditions, mean annual temperature is relatively low and diurnal temperature amplitudes are small. In contrast, the short dry season climate is characterized by exceptionally high day-time and low nighttime temperatures. Trees of this ecosystem have to cope with constantly low temperatures, compared to lowland forests, while being exposed to unpredictable short-term microclimatic shifts. The objectives of this study were (i) to analyse the temperature-response of R_S of adapted moist forest tree species during two contrasting seasons, i.e sunny-dry versus cloudymoist weather conditions, and (ii) to compare patterns of R_S of eight co-existing tree species from seven families for quantifying variability between trees. Linear regression analyses and stepwise multiple regressions with selected meteorological factors (vapour pressure deficit, wind speed, solar radiation) were used to take account of abiotic drivers controlling transpiration and photosynthesis, which may affect R_S under contrasting climatic conditions. We hypothesize that plant inherent factors and external variables apart from temperature are playing an important role in determining the apparent R_s at the level of individual trees.

3.3 Materials & Methods

Study area and tree selection

The study was conducted in the Reserva San Francisco (RSF) adjacent to the Podocarpus National Park (PNP) on the eastern slopes of the southern Ecuadorian Andes. The study site (S 03°58'345''/ W 79°04'648'') is located at 1890 m elevation and covered by a species-rich lower montane forest of about 12 m in height. Forest structure and species composition are described in more detail by Röderstein et al. (2005) and Leuschner et al. (2007). The soils of the region developed from metamorphic shale, quartzite and sandstone bedrock; they are characterized by low fertility (Schrumpf et al. 2001).

Within the stand we selected 20 mature canopy trees representing abundant families and comprising a broad range of stem diameters. Further details on tree selection are given in Zach et al. (2008) (see Chapter 2 and Appendix 1).

For the exemplary study on the impacts of the dry season conditions on stem respiratory activity we restricted analysis to 10 out of the 20 mature canopy trees sampled (Table **3.1**). These 10 trees had been measured during a pronounced dry season period characterized by high ambient temperature, high vapour pressure deficits (D) and declining soil water availability (Figure **3.1**). The remainder 10 trees were measured one week earlier, when some rain fell, which did not allow for an appropriate comparison of wet and dry season respiratory patterns. The tree selection comprised canopy tree species of 7 families (Table **3.1**). Most of the tree species were represented by only one individual, except *Miconia punctata* and *Matayba inelegans* of which each two individuals of similar size were sampled (Table **3.1**). Trees were equipped with dendrometer bands for increment measurements.

Rates of stem CO₂ release (R_S) of the selected tree individuals were measured during four days in the dry period in November 2005(DOY 324-327; Figure **3.1**) and four days during the humid season in April 2006 (DOY 117-120). For each tree we conducted one 24hrs-course of R_S . Diurnal values were averaged to calculate the mean daily rate of R_S of each tree (Table **3.1**), while the entire dataset was used for regression analysis (Table **3.2**, Figure **3.2**). R_S was monitored at breast height (1.3 m) using the mobile 6-chamber respiration system ANARESY 2 (Walz, Effeltrich, Germany) and an integrated LI-7000 infrared gas analyzer for CO₂ and H₂O (Li-Cor, Inc., Lincoln, NE, USA) running in differential mode. Details of the technical equipment are provided in Zach et al. (2008; Chapter 2).

'able 3.1. Mean stem CO_2 release rate (R_s , μ mol CO_2 m ⁻² s ⁻¹) of individual trees at the study site
t 1890 m elevation as measured during the dry (R _s dry) and during the humid (R _s wet) season of
ne year. Relative stem diameter increment (in %) as determined between September 2005 and
pril 2006. $d=$ stem diameter in cm, $h=$ stem height in m, SD in parentheses.

Table at 189 the ye April 2	3.1. Mean stem CO ₂ release 0 m elevation as measured duar. Relative stem diameter i 2006. d= stem diameter in cn	rate (R _s , μ mol CO ₂ aring the dry (R _s dr ncrement (in %) as n, h= stem height in	e m ⁻ s ⁻¹) y) and du determi m, SD in	of individ uring the P ned betwe 1 parenthe	ual trees at the numid (R _s wet) een September ses.	study site season of 2005 and	
Abbr.	Species	Family	q	प	$\mathbf{R}_{\mathrm{S}}\mathrm{d}\mathbf{r}\mathbf{y}$	R _s wet	Stem diameter increment
Cr	Clethra revoluta	Clethraceae	23.45	12.69	0.99 (0.10)	0.41 (0.09)	0.68
Ео	Endlicheria oreocola	Lauraceae	20.80	10.00	0.87 (0.24)	0.14 (0.07)	0.05
Ge	Graffenrieda emarginata	Melastomataceae	26.11	16.54	0.61 (0.18)	0.07 (0.10)	0
Hm	Hyeronima moritziana	Euphorbiaceae	9.52	13.70	0.39 (0.13)	0.22 (0.07)	1.17
Ia	llex amboroica	Aquifoliaceae	9.63	7.50	0.18 (0.11)	0.33 (0.07)	ċ
Lo	Ladenbergia cf oblongifolia	Rubiaceae	9.48	7.92	0.28 (0.08)	0.19~(0.08)	ċ
Mil	Matayba inelegans	Sapindaceae	8.17	8.30	0.85 (0.22)	0.40 (0.11)	0.98
Mi2	Matayba inelegans	Sapindaceae	10.53	12.70	1.04 (0.34)	$0.64\ (0.18)$	ż
Mp1	Miconia punctata	Melastomataceae	10.62	14.05	1.34 (0.26)	0.52 (0.07)	ċ
Mp2	Miconia punctata	Melastomataceae	12.00	7.06	0.57 (0.11)	0.64 (0.11)	0

Air temperature (T_A) and relative air humidity at 2 m height inside the stand were monitored using a Rotronic sensor (Rotronic AG, Bassersdorf, Switzerland) connected to the data logger of the ANARESY system (CR 10, Campbell Scientific, Logan, UT, USA). Thermocouples (diameter: 3 mm, length: 20 mm, Siemens, Munich, Germany) for tissue temperature (T_T) measurement were installed at breast height (depth: 10 mm) at two arbitrarily chosen tree stems (DBH: 15 and 20 cm) in July 2006 for continuous measurements. We used data from July/August 2006 and November 2006 to establish relationships between the continuously recorded air temperature and tissue temperature:

(1) wet season:
$$T_T = 19.74 * (1 - \exp^{(-0.0888 T_A)})$$
 (adj $r^2 = 0.83$, p < 0.0001)

(2) dry season:
$$T_T = 23.13 * (1 - \exp^{(-0.0736 T_A)})$$
 (adj r² = 0.87, p < 0.0001)

where T_A is the air temperature measured inside the stand. These equations were used to calculate tissue temperature from air temperature in all periods where tissue temperature was not recorded.

Climatic conditions of the measurement periods

In general, the climate of the study area is humid throughout the year. However, an extremely wet period (April to July) is followed by several months with less frequent rainfall (September to December) (Bendix et al. 2006). Mean annual precipitation at 1960m a.s.l. is ca. 2200 mm (Emck 2007), annual mean air temperature is 15.7 °C and relative humidity 90.7% (Moser et al. 2008).

During November 2005, foehn winds caused an exceptionally dry and sunny period throughout the study area. Fires occurred frequently throughout the wider region during this period. Only 90 mm of rainfall were recorded in November, which fell during the first two weeks (M. Richter, unpublished data). When we started respiration measurements end of November, the rainless period had already lasted for seven days. The lack of rainfall resulted in a substantial decrease of air humidity and an increase of daily temperature amplitudes compared to the wet season conditions. Vapour pressure deficit (D) continuously increased during the three consecutive measurement days in November (Figure **3.1**, data recorded by M. Richter at a nearby climate station, unpublished). Substantially higher wind speeds were recorded during the dry period compared to the humid days (Figure **3.1**).

The average soil water matric potential (ψ_S) as measured in adjacent forest sites at 2000 m progressively decreased during the measurement days in November from -0.15 (± 0.13) to -0.24 (± 0.12) MPa at 15 cm depth and from - 0.10 (± 0.09) to - 0.19 (± 0.12) MPa at 30 cm (S. Engelhardt, unpublished data). Minimum values of – 0.60 MPa (15 cm) and – 0.49 MPa (30 cm) were recorded at the end of the measurements (Figure **3.1**).

The measurement campaign in the wet period was conducted at the beginning of the humid season in April 2006. During this month precipitation was frequent. Low diurnal temperature amplitudes and a moderate vapour pressure deficit were characteristic for this season (Figure **3.1**). Data of the soil matric potential were only available for the last two measurement days. This data indicates that soil moisture exceeded field capacity with potential values greater than -0.1 MPa (S. Engelhardt, unpublished data).

Data treatment

The stem CO_2 release rates (R_S) were calculated as:

 $(3) R_S = D [CO_2] * F/Ac$

where R_s is the CO₂ release rate (µmol CO₂ m⁻² s⁻¹), *D* [CO₂] is the difference between ambient (reference gas) and chamber (sample gas) CO₂ concentration, *F* is the molar air flow rate (mol s⁻¹) which passes through the chamber, and *Ac* is the surface area (m²) of the enclosed stem segment. In an earlier study (Zach et al. 2008, Chapter 2) we found better agreement when relating the apparent CO₂ efflux rates to stem surface area than to volume.

We chose the slope of the regression curve of R_s on temperature as a measure of the temperature sensitivity of respiration instead of the commonly used Q_{10} value. The Q_{10} of respiration (i.e., the factor expressing the proportional change in respiratory CO₂ release per 10 °C rise in temperature) basically compares the CO₂ efflux rate at a given temperature to that at a 10 °C lower temperature (Atkin et al. 2005). Plant respiration is a process influenced by various environmental factors. Consequently, neither the Arrhenius nor the Q_{10} model has been found to adequately describe the observed temperature response of respiration under field conditions. For that reason, reported measures of temperature responsiveness have to be treated with caution (Tjoelker et al. 2001, Atkin and Tjoelker 2003). In addition, there is evidence that the respiratory response to temperature is neither constant nor necessarily exponential (Atkin et al. 2000, Atkin and Tjoelker 2003, Atkin et al. 2005). Instead of focussing on two single data points on the temperature axis, we used the entire diurnal dataset of R_s for analysing the temperature dependency of R_s .

One important source of variation in diurnal and seasonal R_S is the internal transport of dissolved CO_2 in the transpiration stream influencing the apparent CO_2 release through the bark (Edwards and Hanson 1996, Levy and Jarvis 1998, Teskey and McGuire 2002, Gansert and Burgdorf 2005, Teskey and McGuire 2007). Because we did not measure xylem sap flow directly, we used D, wind speed and solar radiation as the main determinants of transpiration in our regression analysis in the attempt to explain the large heterogeneity of R_S .

By means of linear regression we analyzed the temperature responsiveness of R_S for each tree individual during the dry and humid season. For each tree, rates of R_S were plotted against the respective courses of T_T and T_A using either the diurnal dataset, or by considering only day-time or night-time values. To improve the explanatory power of the regression equation, we conducted linear regression analyses with the three additional meteorological predictor variables (D, wind speed, solar radiation). Using coefficient of determination (r²), we quantified the influence of these four variables on R_S . If more than one predictor showed significant effects we conducted stepwise multiple regression analyses ($\alpha = 0.05$ for tolerance, collinearity $|\zeta| < 0.6$) to determine main causes of diurnal, day-time and night-time stem CO₂ efflux variability. Subsequently, multiple regression analyses were repeatedly applied by removing factors with no significant influence (p < 0.05) until each individual tree model reached maximum likelihood (F-test, p-value).

3.4 Results

Stem CO₂ release and its temperature responsiveness under dry and humid conditions

Mean daily R_s was highly variable and ranged from 0.18 to 1.34 µmol m⁻²s⁻¹ during the dry, and between 0.07 and 0.64 µmol m⁻²s⁻¹ during the humid period (Table **3.1**). Mean daily T_T ranged between 15.7 and 16.8 °C during the dry (minimum 10.1 °C, maximum 19.9 °C), and between 14.5 and 14.9 °C during the humid season (minimum 10.8 °C, maximum 17.5 °C). In most cases, R_s was higher during the dry period than under cooler and more humid conditions. Between September 2005 and April 2006, stem diameter increment was very low (Table **3.1**). Due to dendrometer defects diameter increment could not be determined for all trees.

Among regression models, we found linear regression to give the most appropriate fit for the relationship between R_S and temperature for both periods (Figure **3.2**). The respective correlation between R_S and T_T or T_A did not differ (not shown).

Under dry season conditions, the response of R_S to T_T was highly divergent and showed strong discrepancies in the direction of response between different tree individuals. Moreover, the eight measured species showed contrasting patterns in their temperature response of R_S . Besides the commonly expected positive relationship, R_S was significantly negative correlated with temperature or completely uncoupled from changes in temperature in certain cases (Figure **3.2**). In the dry season, the temperature responsiveness of R_S differed between day-time and night-time periods. In most of the cases, we found better correlations with data separated into day-time and night-time values than with the diurnal datasets. While most of the tree stems were more responsive to temperature during day-time hours, we found three tree species showing a higher temperature coupling during the night (Ia, Mi1, Cr; Table **3.2**). The general temperature responsiveness of R_S was unexpectedly low. For most of the tree individuals, plotting diurnal R_S against T_T yielded r² values of less than 0.5. In two of the ten measured tree species, T_T could explain more than 70% of the variability in the diurnal R_S . One tree of the species, *Hyeronima moritziana* (Hm) showed no significant relation to T_T under dry season conditions neither during the day nor the night (Table **3.2**).

In the humid season, R_S was mostly uncoupled from temperature; only in some cases a very weak correlation was observed (Table **3.2**). Two trees showed significant relationships in the diurnal dataset (r^2 = 0.13 and 0.23). Two other trees were weakly, though significantly, correlated with T_T during night-time only (r^2 = 0.22 and 0.23). For the remainder species, no differences in the temperature sensitivity of R_S between day- and night-time were apparent in the humid season.



Figure 3.1. Climate data for the measurement periods in November 2005 (dry season) and April 2006 (wet season) as recorded at a nearby climate station at 2000 m a.s.l. (M. Richter, unpublished data). Vapour pressure deficit (D) was calculated from the climate station data (D_{out}) and from climate data recorded inside the forest site at 2 m above ground (D_{in}). Soil matric potential (ψ_S) was measured at adjacent forest sites at 2000 m (S. Engelhardt, unpublished data).

Dependence of stem CO_2 release on other meteorological predictors under dry and humid conditions

 R_S was linearly related to all four selected meteorological parameters (not shown). Nevertheless, neither D nor wind speed or solar radiation improved the explanatory power of the regressions with respect to R_S variability under dry or humid conditions. Trees showing a significant relationship to T_T also yielded comparable correlations to the other climatological parameters, most likely as a result of tight intercorrelation between the parameters. In most cases, stepwise multiple regression analysis was not applicable due to the high collinearity of the meteorological variables ($\zeta > 0.6$). An exception was wind speed, which was generally less correlated to T_T , D or solar radiation ($\zeta < 0.5$).



Figure 3.2. Stem CO₂ release rates of 10 tree individuals at 1890 m elevation plotted against tissue temperature. Tissue temperature was calculated from air temperature with equation (1) and (2). Separate regression lines are given for the dry season (November 2005, \circ) and the humid season (April 2006, \bullet) for each tree individual. Coefficients of determination are given in Table **3.2**. For key to species abbreviations see Table **3.1**.

Wind speed and radiation were the main determinants of the diurnal variation in R_S of *Graffenrieda emarginata* (Ge) during the dry season, explaining half of the variability (model adj. $r^2 = 0.497$, F = 23.25, p < 0.0001); the correlation with radiation was a negative one (Table **3.2**). Wind speed alone was the most important factor controlling the diurnal variation in R_S of *Matayba inelegans* (Mi1) during the dry season (model adj. $r^2 = 0.337$, F = 24.45, p < 0.0001); again, the relation was a negative one. Wind speed improved the regression model by explaining further 10.2% of the diurnal variation in R_S in *Miconia punctata* (Mp1). Together with T_T , both parameters could explain 57.7% of the variability in diurnal R_S in the dry period (F = 31.07, p < 0.0001). For the remaining tree species, multiple regression analysis either was not appropriate or could not improve the explanatory power of the model in comparison to a single-factor model with T_T alone (Table **3.2**).

Under humid conditions, the variability of R_S could not be explained by any of the meteorological parameters. A few tree species showed significant, yet very weak, correlations to some of the predictors (Table **3.2**). An exception was the night-time variability in R_S in *Matayba inelegans* (Mi1), which was strong and negatively related to D ($r^2 = 0.59$).

3.5 Discussion

Sources of variation in humid season stem respiration

During the humid measurement period, R_S was independent of T_T (Figure 3.2), and also seemed to be unaffected by any of the other investigated meteorological parameters (Table 3.2). Plant respiratory activity is thought to be determined by three processes, substrate supply, demand for respiratory products and potential enzyme capacity (Amthor 1995, Lambers et al. 1998, Atkin et al. 2005). Limited light availability due to cloud cover is known to strongly reduce net carbon gain of tropical forests on a seasonal as well as daily basis (Hollinger et al. 1994, Chen et al. 1999, Graham et al. 2003). In the humid season with frequent cloudiness and rainfall in the Ecuadorian Andes, assimilation rates must have been impeded, thus restricting carbohydrate supply. Porometric measurements conducted on mature canopy trees at nearby forest sites in the Podocarpus National Park showed that leaf transpiration and xylem sap flow of the tropical montane trees were substantially repressed during cloudy days and under high atmospheric humidity (Motzer et al. 2005). This may reduce the transport of dissolved CO₂ with xylem water from the roots to the stem.

al predictors	960 m a.s.l.,	m^{-2}). Given	the dry and	sured during	
n meteorologi	nate station at	le the forest, V	al measured fo	lata points me	
e of 10 trees o	ity from a clin	ecorded outsic	n tree individu	two outlying e	
em CO ₂ releas	and air humic	r radiation as 1	data set of eacl	ed by omitting	
analyses of st	m temperature	s^{-1} ; rad = sola	ight-time (nt) a	*- data calculat	
lear regression	calculated fro	e the forest, m	ime (dt) and n	l in bold. Mi1*	
ttion (r^2) of lir	sure deficit as	station outside	al (day), day-ti	05) are printed	e Table 3.1.
s of determine	= vapour pres	d at a climate	s of the diurna	lations ($p < 0$.	breviations se
n's coefficient	rature, °C; D	ed as recorde	ession analyse	gnificant corre	y to species ab
le 3.2. Pearsor	= tissue tempe	w = wind spe	esults of regre	vet season. Sig	t-time. For key
Tab	(T _T :	kPa;	are i	the v	nigh

		E			C			M		rad	
Dry	day	dt	nt	day	dt	nt	day	dt	nt	day	dt
Cr	0.38	0.26	- 0.30	0.56	0.42	0.07	0.31	0.09	0	0.55	0.38
Eo	0.49	0.53	0.34	0.45	0.51	0.31	0.18	0.19	0.21	0.19	0.16
Ge	- 0.31	0.12	0.05	- 0.24	0.03	0	0.25	0.36	0.18	- 0.21	0.03
Hm	0.02	0	0.01	0.03	0	0	0.01	0.04	0	0.03	0
Ia	- 0.22	0.14	- 0.56	- 0.13	0.09	- 0.49	- 0.13	0.10	- 0.23	0.01	0
Lo	0.10	0.12	0.02	0.11	0.16	0.04	0.01	0	0.03	0.01	0
Mil	0.23	0.23	0.36	0.05	0.04	0.1	- 0.35	- 0.37	- 0.32	0.02	0
Mi1*			0.80			0.09			- 0.62		
Mi2	0.87	0.89	0.67	0.68	0.66	0.05	0.27	0.05	0.03	0.49	0.30
Mp1	0.49	0.68	0.55	0.39	0.65	0.39	0.38	0.42	0.39	0.14	0.26
Mp2	0.77	0.76	0.31	0.64	0.53	0	0.22	0.02	0	0.37	0.11
Wat											
Cr	0	0.1	0	0.04	- 0.31	0.02	0.03	0.02	0	0	0.02
Eo	0.13	0.27	0.13	0.16	0.25	0.11	0.03	0.25	- 0.23	0.20	0.29
Ge	0	0.01	0.11	0	0.06	0	0.10	0.02	0.04	0.06	0.02
Hm	0	0.04	0.05	0.02	0.09	0.07	0.08	0.11	- 0.43	0.03	0.15
Ia	0.03	0	0.06	0.02	0	0.01	0	0	0.09	0.01	0
Lo	0	0.02	0.22	0	0	0.17	0.11	0.01	0.24	0	0
Mil	0	0.09	0.04	0	0.11	- 0.59	0	0.08	0.21	0.01	0.08
Mi2	0.03	0.03	0.23	0	0.04	0.03	0.03	- 0.36	0.07	0.09	- 0.41
Mp1	0.08	0.04	0.15	- 0.16	0.02	0.11	0.06	0.13	0	- 0.19	0.08
Mp2	0.21	0.28	0.06	0.24	0.42	0.03	0	0.20	0	0.18	0.26

However, amount and concentration of xylem sap CO_2 mainly depend on the respiratory activity of the root system involved, while smaller amounts of CO_2 may also be taken up with the soil water. Differences in fine root respiration among the measured tree individuals remain unknown as well as the quantitative contribution of root respiratory CO_2 release to the transpiration stream, which might finally lower or raise rates of CO_2 release through the bark. We assume that several factors have contributed to a slowed down R_S and a low temperature sensitivity of R_S , among which reduced substrate supply from photosynthesis, less demand for respiratory products under somewhat cooler temperatures, and a probably lower contribution of CO_2 from the transpiration stream are the most likely.

Sources of variation in dry season stem CO₂ release

The dry-season R_S exceeded the wet-season rates in most of the measured trees (Figure **3.2**). Overall higher rates of R_S in the dry season may in part be explained by an assumed higher photosynthetic carbon gain accompanied by a higher carbon investment under the sunnier and hotter conditions of the dry period. However, in the Ecuadorian montane forest, annual wood production was generally low (Moser et al. 2008; Table **3.1**) and distinct stem growth periodicities were not visible (Homeier 2004), implying that pronounced differences in growth related respiratory activity between dry and wet season were unlikely.

Temperature coupling was stronger during the dry season, though considerable scatter in the data was also evident (Figure **3.2**). Although the diurnal amplitude of T_T and T_A differed between the dry compared to the humid season (Figure **3.1**), the temperature response of R_S to variations in T_T and T_A did not differ within periods (data not shown). Hence, thermal differences between T_T and T_A could not account for the observed discrepancies in the temperature response of R_S between the two seasons. However, the actual T_T of our tree stems could differ from the calculated T_T values, since the continuous measurements of T_T on the two stems used for the extrapolation started later in the year. Higher actual values of T_T than the modelled ones could be one reason for the better temperature coupling of R_S in the dry season.

Attempts to explain deviations from the common relationship between R_s and temperature also focussed on time lags between temperature variation and CO_2 release (Ryan 1990, Lavigne 1996, Stockfors and Linder 1998), the significance of stem photosynthetic activity (Sprugel and Benecke 1991, Gansert 1995, Pfanz 1999, Strobel 2004), and / or the CO_2 transport with xylem sap flux (Edwards and Hanson 1996, Levy and Jarvis 1998, Teskey and McGuire 2002, Gansert and Burgdorf 2005).

In three of our trees (Cr, Ge, Hm; Table **3.1**), the correlation coefficients between diurnal T_T and R_S were higher when a time lag of 1.5 to 2 hrs was considered in the analysis (data not shown). At least in two of the stems (Cr, Ge), R_S might indeed show a delayed temperature response, which should be related to the bigger tree size prolonging the radial diffusion pathway of CO₂ (Lavigne 1996), while in the smaller stem, a lowered thermal conductivity could be the reason for the observed time delay (Hm) (Gries 2004). Bark photosynthetic activity as a significant factor could be excluded in our tree sample by comparative measurements of R_S in shaded and unshaded chambers (data not shown).

Taking account of a possible influence of xylem sap CO_2 to explain the large heterogeneity in R_S , we related R_S to D, wind speed and solar radiation. These parameters affect stomatal conductance and transpiration and thus the xylem sap flow with its assumed CO_2 transport capacity. Motzer et al. (2005) measured leaf transpiration and xylem sap flow in mature trees nearby found high stomatal sensitivities to D and strong stomatal control of sap flux in these trees. Even on short time scales (i.e. passing clouds) a sudden decline in leaf transpiration could be detected. Thus, the measurements of Motzer et al. (2005) evidenced a great diurnal and seasonal variation in sap flux rates which could well have resulted in large differences of CO_2 transport in the xylem sap. This could explain why D, radiation and wind speed indeed showed a significant effect on R_S in a number of tree species, in particular in the dry season (Table **3.2**). Xylem sap effects could also account for the high release rates despite low temperature sensitivities evident for most of the tree individuals.

Several authors have argued that reduced cell turgor in the stem of droughted trees could cause a transient reduction in the respiratory activity of the living tissue (e.g. Saveyn et al. 2007). This effect would offer an explanation for the apparent day-time depression of R_S in *Graffenrieda emerginata* and *Ilex amboroica* (Ge, Ia; Figure **3.2**) during the dry season.

However, eight of the ten tree individuals showed higher, and not lower, rates of R_S during the sunny/hot period. Although higher rates of R_S could not be directly related to growth in case of the measured trees, the enhanced R_S very likely indicated that the dry season climate principally favoured photosynthetic carbon gain and xylem sap flow, hence the delivery of substrate for plant respiratory activity as well as of dissolved CO₂ from the transpiration stream, and thereby enhancing stem CO₂ release.

Atmospheric uncoupling of stem respiration versus temporal efficiency in energy acquisition - evidence for climate-sensitive and climate-tolerant tree species?

Differences in rates of R_S and temperature sensitivity of R_S between the measured trees were much more pronounced in the dry than in the wet season (Table 3.1) indicating a diverging response of the trees to hotter weather conditions in this species-rich montane forest. The RSF forest with its high species diversity consists of a mixture of different tree functional types and includes typical climax and pioneer tree species co-occurring in the same stand (Homeier 2004). Moreover, our study site at 1890 m may represent a melting point of tree species which usually occur predominantly at higher or lower elevations, thus representing different climatic adaptations. This assumed variety in tree physiological types may offer an additional explanation for the large differences in patterns of R_S. For example, *Ilex amboroica* and *Graffenrieda emarginata*, the two individuals with a significant negative R_s-temperature response, are mostly recorded above 2000 m elevation in neotropical montane forests (www.mobot.org, Jorgensen and Yanis 1999) and thus are growing at their lower distribution limit in the RSF (J. Homeier, personal communication). Distribution preferences towards cooler and more humid climates at higher elevations would account for a stressinduced adverse reaction to the dry season conditions at RSF (Table 3.2, Figure 3.2). In contrast, Clethra revoluta is most abundant between 1500 and 3500 m a.s.l. in Ecuador (Jorgensen and Yanis 1999). This species seemed to be well adapted to the local climate variability, if the strong positive temperature response during the dry season is used as a criterion to assess performance under varying thermal regimes (Figure 3.2). Miconia punctata and Matayba inelegans are usually occurring in tropical lowland forests below 500 and 1000 m elevation, respectively (www.mobot.org, Jorgensen and Yanis 1999). Thus, they are growing at their upper distribution limit in a rather cold environment in the RSF. These species from hotter environments responded to the dry season by a relatively large increase in R_s with increasing temperature, probably revealing a better adaptation to warmer climates. In contrast, Ladenbergia oblongifolia, one of the more abundant tree species between 1800 and 1900 m a.s.l. in the RSF (Homeier 2004), is mainly recorded below 1000 m elevation in the Neotropics (www.mobot.org). This tree individual showed a very weak R_S response to increasing temperature (Table 3.2).

Conclusion

Our study reports on the substantial variability in stem CO_2 release between seasons and among tree individuals in a species-rich tropical montane forest. Xylem sap flow effects and species-specific differences in respiratory activity or the sensitivity of R_S to atmospheric factors could be influential. The plastic response in R_S of adapted moist forest tree species to a drier climate could help to gain insights into the acclimation potential of a mega-diverse plant community to cope with the expected climate warming. In this context, tropical montane forests are of increasing relevance, since the predicted warming by 2 to 4 K during this century will most likely force mesic lowland taxa to migrate upwards (Mahli and Phillips 2004, Mayle et al. 2004). Understanding and monitoring of the present forest community composition and the species-specific responses to environmental changes are a crucial basis for conservation options and will provide clues for reliable future projections.

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Chapter 4

WOOD CO_2 EFFLUX ACROSS AN ELEVATION TRANSECT IN AN ANDEAN MOIST FOREST: SEASONALITY, RESPIRATION COMPONENTS AND STAND LEVEL UPSCALE

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4.1 Abstract

The carbon (C) economy of tropical montane forests (TMF) is controlled by frequent cloudiness hampering assimilatory CO₂ uptake, while only short periods of clear-sky conditions throughout the year allow for high assimilatory C gain. However, TMFs are well known for their high C allocation shift from above- to belowground plant parts with increasing altitude. Altitudinal changes in C allocation patterns must be reflected in the respiratory CO₂ efflux from above- and belowground woody organs. Studying the impact of climate seasonality on wood CO2 release provides a convenient framework to understand the C use efficiency of TMFs. Seasonality of wood CO₂ release from tropical evergreen forest trees has received little attention, but has never been studied for tropical evergreen montane trees. We used a portable CO₂ measurement system, which allowed us to monitor the respiratory CO₂ release from stems (R_S) and coarse roots (R_R) across an elevation transect with study sites at 1050, 1890 and 3050 m a.s.l. in an Andean moist forest in Southern Ecuador. The study aimed to (1) assess the seasonal variability of R_S and R_R and the impact of elevation, (2) separate woody tissue respiration into the two functional components of growth and maintenance respiration, and (3) extrapolate wood C fluxes to the forest stands. R_S, but not R_R showed a clear seasonality within the measurement year. Highest rates of R_S were measured during the dry season, though the increase in R_S could not be simply related to temperature variation. We assumed a high degree of climate sensitivity of R_S, since increased R_S measured under dry season conditions could not satisfactorily be related to stimulated cell growth, but could also indicate C losses via futile cycles. The increasing C allocation to the root system with elevation was associated with a large increase of the coarse root carbon use efficiency from 0.17 at 1050 m to 0.55 at 3050 m. Annual carbon efflux from stems decreased from 167.1 g C m⁻² yr⁻¹ at 1050 m to 37.7 g C m⁻² yr⁻¹ at 3050 m, while coarse root carbon release changed little from 1050 m (40.9 g C m⁻² yr⁻¹) to 3050 m (36.8 g C m⁻² yr⁻¹) reflecting the increasing importance of belowground organs at high altitudes.

Keywords: altitudinal gradient, coarse root respiration, Ecuador, growth respiration, maintenance respiration, stem respiration, tropical montane forest

4.2 Introduction

At present, very limited data are available on the carbon (C) fluxes in tropical montane ecosystems (Brujinzeel and Veneklaas 1998). Studies quantifying the carbon dioxide (CO₂) efflux from woody parts of tropical montane forest trees are entirely missing. However, such information is essential if we are to gain insights into structure and productivity of the tropical montane biome.

Tropical montane forests (TMF) typically show lower aboveground productivity, lower nutrient concentrations in the soil and hence in plant organs and a slower nutrient turnover than tropical lowland forests. This has been related not only to lower temperatures, but also to the frequent cloud cover and high atmospheric humidity that limit photosynthesis and transpiration and hence assimilate supply and nutrient uptake (Bruijnzeel and Veneklaas 1998, Graham et al. 2003). On the other hand, TMFs are known to invest increasing amounts of C into the root system with increasing elevation (Brujinzeel and Veneklaas 1998, Röderstein et al. 2005, Leuschner et al. 2007, Moser et al. 2008). However, the underlying C allocation processes are still unclear (Brujinzeel and Veneklaas 1998). As is typical for the moist tropics, climate seasonality is not very pronounced in TMFs, which experience a perhumid climate throughout the year, interrupted by few drier weeks with less rainfall (Bendix et al. 2006). In this regard, the question remains of how do tropical montane communities manage to cope with prevailing low light conditions and soil nutrient restrictions. A thorough understanding of the environmental factors influencing carbon gain and carbon losses of evergreen montane trees is still lacking. To date, few studies exist on the influence of climate seasonality on physiological or ecological processes in evergreen tropical forests. For an Andean moist forest of southern Ecuador, phenology was found to be mainly triggered by cloudiness, hence by light intensity and temperature (Cueva et al. 2006). Recently, Bräuning et al. (2008) found highly inconsistent growth dynamics among several abundant tree species of the same study area, which could not simply be related to certain climatic events.

The present study aimed to analyze carbon losses from woody plant organs in relation to site characteristics and climate variations across an Andean moist forest transect in southern Ecuador. During a 1-year-measurement period we recorded stem and coarse root respiratory CO_2 release at three study sites at 1050, 1890 and 3050 m a.s.l.. Characterized by prevailing humid climatic conditions, we could take advantage of a pronounced dry season period during our measurements to study the seasonal variability in stem and coarse root respiration of several montane tree species. In addition, we were able to profit from previous inventory studies conducted at the same sites to make use of standing above-and belowground biomass and annual increment data of single components.

The study has the foci to: (1) examine the seasonal variability of stem and coarse root CO_2 release at three different elevations to learn about climate effects on the respiratory activity of woody organs, (2) separate woody tissue respiration into the two functional components of growth and maintenance respiration which may provide insight into patterns of C use of tropical montane tree species and (3) make first estimates on net annual C fluxes from woody parts of trees in an Andean moist forest.

4.3 Materials & Methods

Study area

The study was conducted within and close to Podocarpus National Park on the eastern slopes of the southern Ecuadorian Andes between August 2005 and September 2006. Across an elevational transect of 2000 m, we chose three forest sites of 400 m² each and with a maximum distance of 30 km between the sites. The sites were a premontane forest at 1050 m, a lower montane forest at 1890 m and an upper montane forest at 3050 m elevation (Table **4.1**). The three sites have been studied previously for above- and belowground biomass and productivity (Röderstein et al. 2005, Leuschner et al. 2007, Moser et al. 2008).

The study area experiences a humid climate with a rain peak from May to July and a less humid period between November and February (Bendix et al. 2006). While easterly winds are responsible for the prevailing humid climate for most of the year, westerly foehn winds frequently cause dry and sunny days between November and February. Foehn winds were especially pronounced during our measurement period between October and December 2005, causing a marked dry season throughout the study area. The nutrient-poor soils of the area developed from granodiorite (1050 m), or metamorphic shale, quartzite and sandstone bedrock (1890 m, 3050 m) (Schrumpf et al. 2001).

	1050 m	1890 m	3050 m
Forest type	premontane	lower montane	upper montane
Coordinates	S 04°06`54``	S 03°58`345``	S 04°06`711``
	W 78°58`02``	W 79°04`648``	W 79°10`581``
Inclination (°)	26	31	27
Rainfall (mm)	c. 1900	c. 2200	c. 4500
RH (%)	88.7	90.7	93.5
T_{Air} (°C)	19.4	15.7	9.4
Canopy height (m)	31.8	18.9	9.0
Tree height (m)	15.6 (0.7)	10.1 (0.4)	5.2 (0.3)
DBH (cm)	17.3 (1.3)	12.2 (0.8)	7.2 (0.4)
Basal area $(m^2 ha^{-1})$	33.6	36.9	42.2
Stem density (n ha ⁻¹)	968	2333	8317
Wood density $(g \text{ cm}^{-3})$	0.64 (0.03)	0.60 (0.04)	0.69 (0.03)
LAI $(m^2 m^{-2})$	6.0	5.7	2.2
SAI $(m^2 m^{-2})$	0.52	0.51	0.52

Table 4.1. Characteristics of the three study sites at 1050, 1890 and 3050 m elevation (means \pm SE; Moser 2008). Stem area index (SAI) is derived from a cone-shaped tree model.

Sampling

For measuring rates of wood CO₂ release we chose 13 - 21 canopy trees per site. The tree selection at each site comprised 13 - 16 different species belonging to 10 - 11 abundant families. The range of DBH (at 1.3 m height) was 8.7 - 43.8 cm at 1050 m, 3.0 - 26.5 cm at 1890 m, and 2.5 - 17.7 cm at 3050 m (Zach et al. 2008, Appendix 1). In addition, 4 - 8 woody coarse roots (diameter: 1 - 4 cm) growing a few centimetres below the soil surface were measured at each site. The trees were equipped with dendrometer bands at breast height (1.3 m) for stem increment measurements (accuracy: 0.01 cm). Tree diameter was taken initially and at the end of the last measurement campaign to determine annual increment rates. Coarse root diameter was measured in the middle of the section enclosed by the measurement chamber at the beginning of the study.

CO₂ efflux measurements

Between August 2005 and September 2006, we conducted four measurement campaigns at each site to monitor the diurnal rates of CO_2 release from stems (R_s) and coarse roots (R_R) (Figure 4.1). The individual campaigns lasted for 10 – 21 days with the measurement system rotating between the three sites. The annual cycle included a pronounced dry season between October and December 2005, which in the following will be referred to as D1, and three measurement campaigns conducted under more humid climate conditions
between August and September 2005 (H1) and from March to September 2006 (H2 and H3; Figure **4.1**). The measurement campaign conducted at 1890 m in April 2006 (H2) was done twice; we repeated measurements when the first unusually dry part of April (H2_{dry}) was followed by more typical rainy weather (H2_{wet}; Figure **4.1**). Due to equipment failure at the beginning of the campaigns in August 2005, we had to reject the dataset of the first campaign (H1) from the site at 3050 m, and the first measurements on coarse roots (H1) at 1890 m.

Rates of CO₂ release were measured using the mobile 6-chamber respiration system ANARESY 2 (Walz, Effeltrich, Germany; Appendix 2) with an integrated infrared gas analyzer for CO₂ and H₂O (LI-7000, Li-Cor, Inc., Lincoln, NE, USA) running in differential mode. Stem CO₂ release was measured at breast height (1.3 m) using tightly fitted plexiglas chambers (95.1 cm³); coarse root sections of 15 cm were enclosed in cylindrical plexiglas chambers (473.8 cm³). We recorded air temperature and relative air humidity at 2 m height inside the stand using a Rotronic sensor (Rotronic AG, Bassersdorf, Switzerland) connected to the data logger of the ANARESY system (CR 10, Campbell Scientific, Logan, UT, USA). For further details on the technical equipment see Zach et al. (2008) (Chapter 2).

We used tissue temperature data logged since July 2006 to extrapolate the stem (T_{TS}) and coarse root tissue temperatures (T_{TR}) of our measurement periods. Thermocouples for tissue temperature (\emptyset : 3 mm, length: 20 mm, Siemens, München, Germany) were installed at 10 mm depth in the stem wood (at breast height) of 2 - 3 randomly chosen trees at 1050 and 1890 m (DBH: 15 - 20 cm) and at 3050 m (DBH: 10 - 15 cm). We related tissue temperature to air temperature and used this relationship to calculate T_{TS} and T_{TR} of the measurement campaigns from site microclimate records (Zach et al. submitted; Chapter 3).

After measurements, we extracted wood cores from all tree individuals using an increment borer (5 mm diameter, Haglöf, Långsele, Sweden) by coring horizontally from the bark to the centre of the bole. Cores were taken in the section where CO_2 release was measured on the stem. Root samples were taken by cutting segments of 3 - 4 cm lenght. We recorded the fresh weight of all samples and used their length and diameter to calculate the sample volume. After drying the samples to constant weight at 70° C, we determined wood density as the dry mass per volume of fresh wood. The dry samples were analysed for their nutrient contents (Appendix 4).

Data treatment

Rates of CO₂ release were calculated as:

(1) $R = D [CO_2] * F / Ac$

where *R* is the CO₂ release rate (μ mol CO₂ m⁻² s⁻¹), *D* [CO₂] is the difference between ambient (reference gas) and chamber (sample gas) CO₂ concentration, *F* is the molar air flow rate (mol s⁻¹) which passes through the chamber, and *Ac* is the surface area (m²) of the enclosed stem or root segment. Due to a better relation of the CO₂ efflux to surface area than to volume, we based our release rates on the surface area (Zach et al. 2008; Chapter 2).

Separating growth and maintenance respiration

In tropical moist forests, where maintenance respiration (R_m) cannot be determined by dormant season measurements as in temperate forests (Ryan 1990, Sprugel 1990, Sprugel and Benecke 1991), R_m is estimated by subtracting the calculated growth respiration from total CO₂ efflux rates. Growth respiration (R_g) of the tree stems (R_{Sg}) was determined following Ryan et al. (1994):

(2) $R_g = 0.248 * p * V_G * (C_{sample} * (10^6 \,\mu mol/12g \,C) * (1/365 \,days *86400s \,day^{-1}))$ where 0.248 is the estimated mean carbon cost per gram carbon incorporated (Meir and Grace 2002), *p* is the wood specific gravity (g cm⁻³), V_G is the volume of annual wood growth under the measurement chamber (cm³m⁻²) and C_{sample} is the wood carbon content (g C g biomass⁻¹). The annual increment of the woody tissue under the chamber was estimated from diameter increment measurements and wood specific gravity as determined for each tree stem from the wood cores. We then calculated the production of woody biomass from the sample carbon content and by assuming an ash-free dry matter content of 99.3% (Ryan et al. 1994).

Estimates of coarse root growth respiration (R_{Rg}) were based on mean annual R_R efflux rates (unit: µmol CO₂ m⁻³ s⁻¹) and annual coarse root biomass production (Moser 2008). Coarse root biomass was assumed to contain 50% carbon in the ash-free dry matter and growth respiration was assumed to account for 25% of the carbon content of the dry-matter production (Penning de Vries 1975, Sprugel and Benecke 1991).

Up scaling of wood CO₂ efflux

Estimates of annual total carbon release at stand level were based on wood CO_2 efflux measurements and forest inventories of the study sites (Moser 2008). Because there were no significant relationships between basal area and R_S of the measured trees at any of the sites, we did not separate size classes (Figure **4.2a**). Instead, we assumed that rates of stem

respiration were constant across tree sizes for extrapolating stem C release to stand level. Carbon release for each tree stem was calculated assuming a cone-shaped stem and using tree height as a measure for stem length from the ground to the top of the tree. The calculated totals of carbon release for each tree were summed up to the stand. The cumulative vertical stand C fluxes were projected to the ground area by dividing the total site area (i.e., by multiplying vertical stand C flux with the stem area index, SAI). Site areas were corrected by slope angle (Table **4.1**). Up-scaling did not include estimates of branch or leaf carbon release.

Coarse roots total respiration estimates were based on mean annual R_R efflux rates (µmol CO₂ m⁻³ s⁻¹) and the standing coarse root biomass stock as determined by Moser (2008). Volume-based release rates were converted to biomass using the specific wood gravity as determined for the measured roots (data not shown). Coarse root biomass was assumed to consist by 50% of carbon, and ash-free matter to represent 99.3% of the biomass.

Statistical analysis

Differences between mean rates of stem respiration in the different measurement campaigns conducted at each site (1050 m, 1890 m, 3050 m) were tested for significance using analysis of variance (ANOVA). For the sites at 1050 m and 1890 m, the daily mean R_s values of the individual trees were log-transformed prior to ANOVA to achieve homogeneity of variances (Scheffé test, p < 0.05). Stem CO₂ release data of individual trees at 3050 m matched parametric assumptions without transformation. We tested for significant differences in coarse root CO₂ release (R_R) between the measurement campaigns at each site using the same procedure as for stem CO₂ efflux (Scheffé test, p < 0.05). Root respiration data matched parametric assumptions without transformation. We used coefficient of determination (r²) to quantify the influence of predictor variables (basal area, stem increment, nitrogen content) on stem respiration.

4.4 Results

Seasonality of wood CO₂ efflux across the elevation transect

Mean rates of R_s at 1050 m were significantly higher during D1 compared to more humid campaigns (Figure 4.1). Differences between campaigns were not significant at the sites at 1890 and 3050 m.





Figure 4.1. Meteorological conditions and stem (R_S) and coarse root (R_R) CO₂ efflux rates (mean ± SD) in the measurement campaigns conducted between August 2005 and September 2006at the three study sites at 1050 m, 1890 m, and 3050 m elevation. At each site, measurements were conducted during three humid season periods (H1 to H3) and during one dry season period (D1). At 1890 m, the measurement campaign H2 was done twice; we repeated measurements when the first unusually dry part of April (H2_{dry}) was follow by more typical rainy weather (H2_{wet}). For all sites, daily means of air temperature (T_{Air}) and relative air humidity (RH) and diurnal precipitation (P) are given for the study sites at 1890 m and photosynthetic active radiation (PAR) for the site at 3050 m. Given are mean daily CO₂ efflux rates for each individual tree and coarse root measured as well as the mean stand efflux per measurement campaign (\Box). Different letters indicate significant differences (Scheffé, p< 0.05) between measurement campaigns for mean stand rates of stem and coarse root CO₂ efflux at each study site.

Nevertheless, highest mean rates of R_S were measured during the driest and warmest period of the measurement year at all three sites (D1 and H2_{dry}; Figure **4.1**). During the measurement year, mean stand R_S ranged from 0.54 - 1.38 µmol CO₂ m⁻² bole surface area s⁻¹ at 1050 m, from 0.28 - 0.79 at 1890 m and from 0.16 - 0.21 at 3050 m; minimum rates of R_S were measured under more humid conditions (H1 to H3). Between tree individuals, the variability in rates of R_S was higher during D1 compared to H1, H2 and H3 at all sites (Appendix 5). This seasonal contrast was especially pronounced at 1890 m (Figure **4.1**). Differences in stand R_S between campaigns were less apparent at 3050 m (Figure 4.1). Annual mean rates of R_S decreased with increasing elevation from 0.86 ± 0.40 (SD) µmol $CO_2 \text{ m}^{-2}$ bole surface area s⁻¹ at 1050 m to 0.54 ± 0.23 at 1890 m and 0.19 ± 0.03 at 3050 m. Elevational differences were most pronounced during D1 and least marked during H1, H2 and H3 (Figure 4.1; Appendix 5).

Mean annual rates of coarse root respiration (R_R) decreased from 0.36 ± 0.09 µmol CO₂ m⁻² surface area s⁻¹ at 1050 m to 0.26 ± 0.04 at 1890 m and to 0.16 ± 0.01 at 3050 m. Rates of R_R did not show any seasonal trend at the three sites (Figure **4.1**; Appendix 5 and 6).

Growth and maintenance respiration

Based on absolute values stem growth respiration (R_{Sg}) decreased continuously with increasing elevation (Table **4.2**). However, the average proportion of R_{Sg} on total R_S decreased only slightly with altitude and accounted for 14.2% ($0.11 \pm 0.12 \mu mol CO_2 m^{-2} s^{-1}$, SD) at 1050 m, 13.2% (0.05 ± 0.06) at 1890 m and 10.3% (0.01 ± 0.02) at 3050 m. At tree level, the fraction of R_{Sg} on total R_S ranged from 0 - 75% at 1050 m, from 0 - 52% at 1890 m and from 0 - 70% at 3050 m. We did not find any correlation between stem maintenance respiration (R_{Sm}) and tissue nitrogen content at any of the sites (Figure **4.2b**). Mean annual R_S was weakly correlated with annual wood increment at 1890 m, whereas no relationship was found at 1050 and 3050 m (Figure **4.2c**).

The fraction of coarse root growth respiration (R_{Rg}) on total R_R continuously increased with increasing elevation and amounted to 5.2% at 1050 m, 14.6% at 1890 m and 29.9% at 3050 m (Table **4.2**).

Annual C losses from woody tissue across the elevation transect

Stand carbon release from stems accounted for 167.1 g C m⁻² ground area yr⁻¹ at 1050 m, 102.6 g C m⁻² ground area yr⁻¹ at 1890 m and 37.7 g C m⁻² ground area yr⁻¹ at 3050 m based on the stem surface area of a cone-shaped trunk model (Table **4.2**). Using cylinder-shaped tree stems would exactly double the annual carbon flux estimation derived from the cone model.

Stand carbon release from coarse roots was 40.9, 19.2 and 36.8 g C m⁻² root surface area y⁻¹ at 1050 m, 1890 m and 3050 m, respectively (Table **4.2**).



of y-axis at 3050 m.

(g C m⁻² yr¹) and mean amual stem respiration (R_s) at the three sites at 1050, 1890 and 3050 m elevation. Note different scale

4.5 Discussion

Seasonality of wood CO₂ efflux across the elevation transect

Few studies have been conducted on the annual variability in wood CO_2 efflux of tropical plants, and most are related to drought-deciduous tree species. In habitats with strongly seasonal rainfall, woody tissue respiration is highest during the wet season of the year. This is little surprising, since leaves are shed with the onset of the dry season and wood production and associated stem growth respiration ceases (Levy and Jarvis 1998, Nepstad et al. 2002, Silva et al. 2002, Chen et al. 2003, Chambers et al. 2004). In this respect, drought-deciduous forests are similar to cold-deciduous temperate broadleaved communities and do not allow for a direct comparison with the evergreen vegetation of the moist tropics. To our knowledge, the only study addressing the seasonality of wood CO_2 efflux from tropical evergreen trees was conducted in an aseasonal wet lowland forest and found no indication of seasonality in wood carbon release (Cavaleri et al. 2006).

Our study of stem CO_2 release across an Andean moist forest transect revealed that R_s considerably augmented under dry season conditions, while rates of R_s were comparatively low under the prevailing more humid climate (Figure **4.1**). The increase in R_s under dry season conditions was especially pronounced at 1050 m and 1890 m. The environmental sensitivity of R_s at 1890 m was corroborated by two consecutive measurement campaigns conducted in April 2006. While high rates of R_s were recorded during the first exceptionally dry part of April (H2_{dry}) with weather conditions comparable to the D1 campaign, mean R_s dropped by 65% of the previous intensity during the rainy second part of the month (H2_{wet}; Figure **4.1**). The results clearly indicate that the monitored trees were very responsive to climatic variations.

However, the higher respiration rates could not simply be explained by the higher ambient temperatures recorded during D1 (Figure 4.1). In contrast to general belief, the respiratory response to temperature was neither exponential nor clearly linear for the measured moist forest trees throughout our measurement campaigns (Appendix 3a-c). Even during D1, plotting R_s against temperature yielded a highly scattered picture at all three sites ($r^2 = 0.017 - 0.064$: Zach et al. 2008; Chapter 2). In a previous study on diverging temperature response patterns at 1890 m (Zach et al. submitted; Chapter 3) we found that the R_stemperature relationship varied considerably in degree and direction of response between tree individuals during the dry season campaign ($r^2 = 0.02 - 0.87$), while R_s was mostly uncoupled from temperature under wet season conditions ($r^2 = 0 - 0.21$). The diverging temperature reactions between dry and humid season were not related to lower temperatures or diminished diurnal temperature amplitudes under humid conditions, since R_S was related to stem tissue temperature and the diurnal amplitude of tissue temperature was equally weak for both seasons (Zach et al. submitted). Therefore, Zach et al. (submitted) proposed a complex interaction of abiotic factors (incoming PAR, mean daily temperature, wind speed, vapour pressure deficit, soil water availability), which affect assimilation, enzymatic activity or the demand for respiratory products and hence R_S in these tropical montane forests. Further, an influence of CO_2 supplied by xylem sap flux is possible (e.g., Teskey and McGuire 2002, Teskey and McGuire 2007). In conclusion, periods of highest respiratory activity were exceptionally dry and warm, but respiration was not related to temperature in a simple way.

At 3050 m, differences in rates of R_s between campaigns were less pronounced than at 1050 and 1890 m. However, the climatic differences in terms of incoming solar radiation or mean daily temperature between campaign D1 and H3 affected R_s at this high-elevation stand as well (Figure **4.1**). On the other hand, the R_s -temperature relationship was highly variable between campaigns and between individuals showing much scatter in the data of this site in most cases (data not shown). We measured one unexplainable high rate of R_s during H2 (Figure **4.1**), which substantially increased mean rates of R_s at 3050 m and hence lowered the difference between humid and dry season results. Omitting this value would yield a more clear trend of higher respiratory activity under favourable radiation and temperature conditions as are indicated by D1.

Coarse roots respiration did not vary between campaigns, probably as a result of low changes in soil temperature regime throughout seasons at the same depth roots were (S. Graefe, personal communication). To our knowledge, there are no comparable studies on R_R in tropical montane forests.

Growth and maintenance respiration

Tropical lowland forests that are dominated by slow-growing, tall climax tree species are assumed to use about 80% of the total respired carbon for maintenance processes (Meir and Grace 2002). Ryan et al. (1994) reported R_{Sg} values of 18% and 46% in slow- and fast-growing tropical lowland species, respectively. In an afrotropical lowland forest, 20% of the CO₂ released by tree stems was used for tissue construction (Meir and Grace 2002). Stem growth respiration in a seasonally flooded Varzea forest was below 15% of total CO₂ release for most of the trees measured (Horna 2002). A proportion of 10 - 14% for stem growth respiration as found in our study is in the lower range of literature data. We explain this low

figure mainly from the fact that the existing data are exclusively from tropical lowland sites with a much higher aboveground productivity than montane forests.

Table 4.2. Standing carbon stocks, productivity and C flux components of the sites at 1050, 1890 and 3050 m elevation. Total aboveground (AG) and belowground (BG) biomass, stem (S) and coarse root (CR) biomass (g C m⁻²), component and total net primary (NPP) productivity (g C m⁻² yr⁻¹) and annual stem (R_S) and coarse root (R_R) C fluxes (g C m⁻² ground area yr⁻¹). Given are stand-level C flux calculations (R_S and R_R) based on own CO₂ release measurements and forest inventory data (G. Moser, unpublished data), and estimates of the C costs of stem growth (R_{Sg}) and coarse root growth (R_{Rg}) respiration.

	1050 m	1890 m	3050 m
AG-biomass	14255	8650	5610
S-biomass	13915	8160	5430
Stand-level R _S (cone-based) *	167.12	102.62	37.66
S-growth	141	43.5	7
% of total S-biomass	1.0	0.5	0.1
R_{Sg} (%)	14.2	13.2	10.3
R _{Sg}	23.73	13.55	3.88
CUEs	0.46	0.30	0.16
R _S /S-biomass	0.012	0.013	0.007
BG-biomass	1605	1305	3135
CR-biomass	1470	995	2595
Stand-level R _R	40.9	19.2	36.8
CR-growth	8.44	11.40	44.20
% of total CR-biomass	0.6	1.2	1.7
Estimated R _{Rg}	2.11	2.80	11.00
Estimated costs (%)	5.2	14.6	29.9
CUE _{CR}	0.17	0.37	0.55
Total NPP	652	621	653

* Cylinder-based calculations would result in double the amount of the cone-based values

The proportion of R_{Sg} on total R_S of our montane tree species remained constant with elevation. This is surprising, since stem growth declined 10-fold from 1050 m to 3050 m (Table 4.2). At 3050 m, aboveground tree growth was rather low, which is thought to result mainly from unfavourable environmental conditions such as low light intensity, lower temperatures and waterlogged soils, reducing C gain and nutrient supply (Leuschner et al. 2007, Soethe et al. 2007). On the other hand, a higher wood density and growth-hampering conditions may result in higher construction costs of new tissue at 3050 m as found for high elevation trees (Brujinzeel and Veneklaas 1998). This could have counteracted low rates of R_{Sg} . As a consequence, rates of R_{Sg} were more uniform across the elevation transect. A different picture is found for rates of coarse root growth respiration (R_{Rg}). In parallel to the enormous five-fold increase in coarse root growth from 1050 m to 3050 m, the fraction of R_{Rg} on total R_R increased markedly from 5% (1050 m) to 30% (3050 m) in our transect, emphasizing the increasing importance of root growth with elevation (Table **4.2**). However, our root respiration estimates were based on very few records on roots with diameters between 1 and 4 cm, while the root biomass data and dry matter production values also included large woody roots (> 4 cm) as well as smaller size fractions (2-10 mm). Therefore, our calculated values can only provide a rough approximation of the actual amount of R_R .

Maintenance respiration is often strongly correlated with tissue nitrogen content (McCree 1983, Waring et al. 1985, Irving and Silsbury 1987, Ryan 1990, Ryan 1991), based on the assumption that metabolic activity for maintenance functions is dependent on the protein content of plant tissue (Penning de Vries 1975). However, the strong dependence of maintenance respiration on nitrogen is a theoretical construct based on biochemical principles, which still has to be validated under field conditions (Van der Werf et al. 1992). In fact, we could not find a close relationship between R_{Sm} and tissue N-content at our sites (Figure **4.2b**).

The calculation of growth respiration is typically based on the assumption of constant growth rates throughout the year. This assumption may be overly simplifying, since even tree species of the moist tropical lowlands show marked seasonal growth trends, triggered by changes in temperature, rainfall and nutrient availability (Clark and Clark 1994, Devall et al. 1995, Clark et al. 2003, Verheyden et al. 2004, Bräuning and Burchardt 2006). In the study area, cambial activity (and hence growth) typically ceased during periods of heavy rainfall and low light intensity (Bräuning et al. 2008, Bräuning et al. in press). On the other hand, Bräuning et al. (2008) found that cambial activity was mainly controlled by soil water availability. Even after only few rainless days, growth has halted and stem diameter started to shrink as a result of a decrease in the stem water status. Interestingly, the stem respiratory activity of our trees decreased under wet season conditions, but seemed to be stimulated by drier and warmer periods when the soil matrix potential may drop to -0.5 MPa (at 30 cm depth; S. Engelhardt, unpublished data). In this context, it is unclear if the high stem respiratory activity measured during the driest and warmest campaign (D1, Figure 4.1) was indeed related to enhanced cambial activity as previously assumed, or if higher rates of woody tissue respiration were the result of internal stress due to fluctuations in cell water status.

In a tropical evergreen lowland forest of Costa Rica, growth declined with increasing temperature and the decrease in productivity was strongest during the hottest year (El Niňo).

At the same time, the ecosystem CO_2 efflux increased (Clark and Clark 1994). Other recent studies consistently confirmed the trend of decreasing NPP in tropical forests during warmer years (Worbes 1999, Clark et al. 2003, Feeley et al. 2007). In addition, carbon respired via the alternative pathway of respiration was found to increase exponentially with increasing temperature in the humid tropics (Keller and Lerdau 1999, Lerdau and Throop 1999). A high climate sensitivity of tropical forest productivity would have serious implications for future rates of CO_2 enrichment in the atmosphere under the ongoing global warming.

Error propagation in extrapolating chamber measurements

Methods for estimating stand C release are not standardized and uncertainties exist concerning the choice of scalar (ground area, surface area, sapwood volume) and of the allometric equations used or the representativeness (tree diameter size distribution, tree density) of the study site (Levy and Jarvis 1998, Meir and Grace 2002, Cavaleri et al. 2006). Consequently, errors can propagate considerably when chamber measurements are extrapolated to the stand level (e.g., Damesin et al. 2001). Chave et al. (2004) illustrated that the largest source of error in estimating tropical forest biomass was the allometric tree model used. The vertical stem CO₂ efflux is commonly projected to a ground surface unit by multiplying with the stem area index (i.e., unit bark surface area per unit ground area; SAI). Differences in the tree model used to calculate the SAI exist (cylindrical: e.g., Levy and Jarvis 1998; cone-shaped: e.g., Meir and Grace 2002; truncated cone: e.g., Yoneda 1993) and can considerably alter efflux results. The cone-based tree model yielded 50% lower values of stem C efflux than the cylindrical trunk model (Table 4.2, Table 4.3). In our study, the cone-based calculation may slightly underestimate the actual ground-projected stem C efflux. However, the latter must conceivably overestimate rates of bole C release per unit ground. Any other alternative method or allometry is not available for accurate determination of stem surfaces.

We found no significant relationship between rates of R_S and stem diameter size or basal area at any of the sites (Figure **4.2a**). This may indicate that tree size is less important for rates of R_S than species identity or demographic positions in this species-rich tropical montane forest. Several studies calculating stand-level CO₂ release in temperate (Strobel 2004, Gansert 2002, Gries 2004) and tropical forests (Ryan et al. 1994, Horna 2002, Meir and Grace 2002, Nepstad et al. 2002, Cavaleri et al. 2006) distinguished different diameter size classes prior to up-scaling. However, Ryan et al. (1994) found stem diameter to explain only 20% of the variability in the stem CO₂ release of one single tropical lowland tree species.

temperatu	and one boreal forest si re (T _{Air} , °C), leaf area ind	te. Standing w lex (LAI), stan	d basal ar	ass is as ass is as ea (m ² ha	sumed ¹ and	to contei mean sta	nt 50 % c ind tree he	arbon. Given	n are annu n). Altituc	al precipitatio e in m a.s.l	n (PPT, mm), mean annual
					I	3asal			Stem	R _S / Stem	
Altitude	Forest type	Site	PPT	T _{Air}]	LAI a	ırea	$\mathbf{h}_{\mathrm{mean}}$	R _{Samual} I	3iomass	biomass	Source
1050	premontane	Ecuador	1900	19.4	6.0	33.6	15.6	167.1*	13915	0.012	this study
1890	montane	Ecuador	2200	15.7	5.7	36.9	10.1	102.6*	8160	0.013	this study
3050	upper-montane	Ecuador	4500	9.4	2.2	42.2	5.2	37.7*	5430	0.007	this study
	lowland, secondary	Cameroon	1520	24.5	4.4		36	214	8700^{**}	0.030^{**}	Meir & Grace 2002
	lowland	Brazil	2200	25.5	5.5	29	30	390	11700	0.033	Meir et al. 1996
50-150	lowland, old-growth	Brazil	c. 2900	26.7	5.7			420			Chambers et al. 2004
50	wet, old-growth	Costa Rica	3880	26.2	6.5			220-360	12450	0.018-0.029	Ryan et al. 1994
37-150	wet, old-growth	Costa Rica	4000	26	6.5	24	20	508	8100	0.063	Cavaleri et al. 2004
	wood <10 cm							359	1200	0.299	
	wood >10 cm							149	0069	0.022	
	Sahelian-shrub-fallow	Niger	562	33-41	0.63		2.2	47.2			Levy & Jarvis1998
	temperate, mixed	Alberta	1400	5-23	4.9	20.1	26	196	5710	0.034	Mahli et al. 1999
300	temperate, beech	France	820	9.2	5.6		13	225			Damesin et al. 2001
	temperate, mixed	Tennessee	1400	13.3	5.0	21.5		149-204	6500	0.023-0.031	Edwards & Hanson 1996
	boreal, coniferous	Canada	400	-20-17	3.3	31.5	6	87	3630	0.024	Lavigne & Ryan 1997
* Cylinde ** Given ai	r-based calculations result in d re aboveground biomass and R	ouble the amount aboveground / Bioma	of the prese SSaboveground ⁻	ented cone- ratio	-based va	lues					
		2	,								

Seasonality of wood CO₂ efflux

Annual wood carbon release across the elevation transect

In terms of standing carbon stocks, wood productivity and rates of stem respiration, our premontane site (1050 m) was similar to a wet tropical lowland forest in Costa Rica (Ryan et al. 1994, Cavaleri et al. 2006). We are not aware of any study on wood C release in a tropical montane forest. We found annual R_s at 1890 m slightly lower than data reported for broadleaved temperate forests (Table **4.3**). Annual R_s at 3050 m was even lower than found for other sites with low aboveground productivity such as semi-arid shrub forests in the Sahelzone (Levy and Jarvis 1998) or a Canadian boreal forest (Lavigne and Ryan 1997). The unexpectedly large gap between our results from tropical montane forests and the results from temperate, boreal and subtropical forests may partly be a consequence of the use of different up-scaling procedures.

In Ecuador, annual R_s declined 4.4-fold over 2000-m elevation distance. In contrast, the annual R_R did not change between 1050 m and 3050 m as the lower specific rates of R_R measured at 3050 m were compensated by the high coarse root biomass stock (Table 4.2). This shift in carbon partitioning was mirrored by a sharp increase of the coarse root carbon use efficiency (CUE: i.e., ratio of production [e.g., coarse root] to gross carbon fixation [e.g., coarse root production $+ R_R$). On the other hand, stem CUE declined with elevation (Table 4.2). At ecosystem level, CUE gives the fraction of carbon invested into growth relative to the total amount of carbon assimilated (i.e., ratio of net primary production to gross primary production). Based on our data of annual R_S and R_R and annual stem and coarse root increment, we calculated a CUE at the stem level of 0.46 and a coarse root CUE of 0.17 at 1050 m. An opposite relationship was found at 3050 m, where stem CUE was 0.16, whereas coarse root CUE amounted to 0.55 (Table 4.2). Low CUE values could indicate restricted nutrient availability which would hamper the use of assimilates for growth. Nutrient limitation typically leads to higher C costs of constructing new tissue (Chambers et al. 2004). A preferential C investment into stem growth and respiration in the premontane forest is an expression of more favourable growth conditions and more fierce competition for light at 1050 m. The upper montane forests are thought to be primarily limited by N shortage (Soethe et al. 2007, Leuschner et al. 2007). A high CUE of the coarse roots at 3050 m reflects that tree anchorage and nutrient acquisition are priority tasks in this environment.

Conclusions

- (1) The evergreen tree species of this tropical montane moist forest showed a clear seasonal pattern in stem respiratory activity.
- (2) Periods of highest rates of stem respiration were particularly warm and dry, but changes in respiratory activity showed no simple relation to temperature.
- (3) Our findings raise the question if the elevated rates of stem respiration measured under dry season conditions are a consequence of stimulated cell growth as was previously assumed, or if excessive carbon was respired via the alternative pathway of respiration.
- (4) The degree of climate sensitivity found in this study indicates that tropical montane forests could be strongly affected by global climate change.
- (5) The increasing carbon allocation to the root system with increasing elevation is associated with a large increase of the coarse root carbon use efficiency (CUE), indicating a higher belowground sink activity at high elevations.

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Chapter 5

THE C BALANCE OF THREE TROPICAL MONTANE FORESTS, SOUTHERN ECUADOR

5.1 Introduction

The carbon (C) balance of forest ecosystems is primarily controlled by assimilatory CO_2 uptake (A_C) and the release of CO_2 via autotrophic respiration (R_a). Part of the C is transferred to the soil either in form of aboveground tree litter or via belowground processes such as root turnover, the exudation of organic compounds to the rhizosphere (Luyssaert et al. 2007) and mycorrhizal hyphal turnover (Godbold et al. 2006). This fraction maintains the soil organic carbon (SOC) pool and the microbial biomass, which, in turn, releases CO_2 back to the atmosphere via heterotrophic respiration (R_h). Although comprehensive estimates exist of the C balance of differing forest biomes most of these studies combine data from a number of sources using different methods in order to reveal a generalized C balance of the tropical, temperate and boreal biome (Malhi et al. 1999, Malhi and Grace 2000, Luyssaert et al. 2007).

Few studies have been conducted determining all above-and belowground components of a forests' C balance within one study site, especially with regard to tropical forests (Malhi and Grace 2000). Even today, establishing a complete C balance is mainly constrained by the lack of data on the C dynamics belowground (Malhi and Grace 2000). While aboveground biomass is usually directly measured empirical data on the belowground counterpart are often missing. Instead, their estimates are based on simple assumptions. For example, Fearnside (1997) assumed belowground biomass to be 33% of aboveground biomass in a tropical lowland forest. Root production was assumed to be proportional to aboveground productivity and the above- to belowground biomass ratio (Malhi et al. 1999). However, it is commonly accepted that such assumptions have often led to underestimate the fraction of belowground carbon fluxes in forest C balances. Despite the lack of exact data on belowground C stocks and dynamics, it can be expected that a large fraction of assimilated C has to be transferred to the root system in order to allow forest ecosystems to maintain steady state conditions (Malhi and Grace 2000); this may especially apply to tropical montane forest (TMF) ecosystems, which are known to substantially invest into their belowground biomass with increasing altitude similar to high latitude forests (Brujinzeel and Veeneklaas 1998, Moser 2008). Boreal forests sequester 84% of their C below the ground and only 16% in the aerial parts, whereas in tropical lowland forests above-and belowground biomass is more or less equally partitioned (Malhi et al. 1999). The question as of how high-elevation tropical montane forests manage this substantial belowground productivity is a central topic in the ongoing debate about TMF forest structure and functionality.

This chapter presents first C balances for the three tropical montane forest sites described in the previous chapters (Chapters 1 to 4) in an attempt to shed some light on the frequently discussed above- and belowground C partitioning patterns of TMFs and their changes with altitude. Therefore, the CO_2 efflux measurements presented above are combined with detailed forest inventory data (Moser 2008) and data on soil respiration (Iost 2007).

5.2 Estimating C balance components of three tropical montane forests

Above- and belowground biomass and biomass production

Data on standing above- and belowground biomass (Table **5.1**) and annual biomass production (Table **5.2**) were provided by G. Moser (Moser 2008, unpublished data). These data refer to the very same study sites and have been assessed during an intensive measurement period from May 2003 to July 2004. Biomass data are given as Mg C ha⁻¹ and increment data as g C m⁻² yr⁻¹ assuming a carbon content of dry matter of 50% (Penning de Vries 1975).

Leaf area index

The leaf area index (LAI) of the study sites was determined using two optical (LAI-2000, hemispherical photographs) and one direct method (leaf mass-related, Table **5.3**; Moser et al. submitted, Moser 2008). Moser et al. (submitted) pointed to several constraints associated with the optical methods and concluded that the direct approach (based on leaf litter production, leaf lifespan and specific leaf area data) provided the most reliable estimate of LAI for the three study sites. Consequently, we based our calculations on the directly determined LAI (Table **5.2** and Figure **5.1**). Since the results of the three methods differ considerably, we summarized changes in LAI-dependent values (canopy CO_2 assimilation, ecosystem respiration) in Table **5.3**.

Canopy CO₂ assimilation

Photosynthetic capacity (A_{max}) was measured at the three sites using a portable open system infrared gas analyser with an integrated blue-red light source inside the leaf chamber (Li-6400, Li-Cor, Inc.). Concentration of CO₂ was kept constant at 360 ppm and air flow was 0.5 1 min⁻¹. We measured A_{max} at a light saturation of 1500 µmol photons m⁻² s⁻¹ (photosynthetic photon flux density, PPFD). Data were recorded when values stabilized (after 2-3 minutes). At 3050 m, we were able to measure attached, sun-exposed canopy leaves from trees within the study site for A_{max} measurements. At 1050 m and 1890 m, the tall stature of trees did not allow for *in situ* measurement of A_{max} . We therefore used representative, small-stature trees in the vicinity of the study sites for A_{max} measurements on attached, sun-exposed leaves (Appendix 7).

The A_{max} of sun-exposed leaves averaged 6.67 \pm 1.83 (SD) $\mu mol~CO_2~m^{-2}$ leaf $s^{-1},\,6.39$ \pm 0.98 and 4.66 \pm 1.74 at 1050, 1890 and 3050 m, respectively, and did not differ significantly between sites (Scheffé, p < 0.05). Canopy CO₂ assimilation (A_C) of the sites was extrapolated based on the measured mean A_{max} values of sunlit leaves and LAI (Table 5.2 and 5.3). In dense tropical forests, most of the canopy leaves are not directly sun-exposed. Based on a "sun/shade" model for photosynthesis simulation, Mercado et al. (2006) calculated the fraction of shaded leaves to account for 70-85% of the canopy LAI in a mature tropical lowland forest in Brazil (LAI 5.7). We therefore assumed a fraction of 80% of the LAI to be shade-leaves and 20% to be sunlit-leaves at 1050 m (LAI 6.0) and 1890 m (LAI 5.7). In contrast, the lower leaf density at 3050 m (LAI 2.2) allows most of the leaves to receive full sunlight (personal observation). Here, we assumed sunlit-leaves to account for a fraction of 80%, while the shaded leaf fraction constitutes 20% of the canopy LAI. We could not find any reference for the relationship between Amax of sun leaves and Amax of shade leaves in multilayered tropical canopy trees. To account for the lower photosynthetic activity of shade leaves, we instead used the proportional difference found between the Amax of leaves from tropical tree seedlings grown under high (1000 μ mol photons m⁻² s⁻¹) and low (100 μ mol photons $m^{-2} s^{-1}$) light intensities (Veneklaas and Poorter 1998). Here, A_{max} of leaves from low-light environments was 23% lower than the Amax of leaves grown under high-light conditions. We therefore reduced the A_{max} of the sun-exposed leaves as measured at the sites for 23% and used this value as a measure of the A_{max} of the shaded canopy fraction. Since we used data on photosynthetic capacity (A_{max}), we assumed a 6-h-day for assimilation activity instead of using the entire 12 hours of daylight in order to partially compensate for the high assimilation rates.

Leaf dark respiration

At 3050 m, leaf dark respiration rates were used as derived from light-response curves of photosynthesis (Appendix 8). We recorded light-response curves of 35 attached leaf samples of 10 canopy tree species abundant at the study site. Leaf dark respiration (R_L) was measured in the dark chamber and recorded after values have stabilized. The R_L values were averaged for mean stand R_L . At 3050 m, mean R_L was 0.53 ± 0.49 (SD) µmol CO₂ m⁻² leaf s⁻¹.

For the sites at 1050 m and 1890 m we calculated R_L by using the relationship between leaf dark respiration and photosynthetic capacity as described by Cavaleri et al. (2008):

(1)
$$A_{max} = (10.9*R_L) / (0.52*R_L)$$

We calculated rates of leaf dark respiration of 0.47 and 0.44 μ mol CO₂ m⁻² leaf s⁻¹ at 1050 and 1890 m, respectively. To account for differences in R_L of sun and shade leaves, we used the proportional difference in R_L measured by Veneklaas and Poorter (1998) for tropical tree seedlings as mentioned in the previous section. The authors found R_L to be 53% lower in seedlings grown under low-light than under high-light conditions. Leaf dark respiration of the sun- and shade-leaf fraction was summed and extrapolated to the annual foliage dark respiration by using the LAI and assuming constant 24-h respiration rates.

Woody tissue respiration

Stem wood tissue efflux values as derived from our measurements at the three study sites were used (Chapter 4). Since branch respiration rates had not been assessed, estimates of branch CO₂ release for the three study sites are based on a rough approximation. Data on wood biomass of the three sites did not distinguish between stem and branch wood, therefore a biomass-based extrapolation was not possible. Branch CO₂ efflux differs from stem CO₂ efflux (Sprugel and Benecke 1991); branch respiration was found to increase with increasing height in the canopy (Maier et al. 1998, Ryan et al. 1996, Cavaleri et al. 2006) and was higher than efflux rates of stems of the same diameter (Cavaleri et al. 2006). Cavaleri et al. (2006) found wood <10 cm to account for 70% of the total wood CO₂ efflux from various functional groups (trees, lianas, palms) in a tropical lowland forest. Lianas contributed substantially to this portion. However, half of total wood CO₂ efflux derived from small diameter wood (0-2 cm) from the upper-, mid- and lower-canopy. Based on these finding we assumed branch respiration to account for 50% of total wood respiration (i.e., stem and branches), which presumably underestimate the actual branch CO₂ release rates at the three sites. However, equal respiration rates of branch and stem respiration rates has also been assumed by other studies (e.g., Ryan et al. 1995).

Root and soil respiration

We used estimates of annual coarse root carbon release as measured and extrapolated above (Chapter 4). Data of total soil organic carbon (SOC) stocks, soil respiration, heterotrophic respiration and fine root respiration have been gained at the very same study sites (Iost 2007). Annual root litter production was assumed to equal annual root production at the three sites (G. Moser, personal communication).

Missing components

The forest C balance consists of several more components which are difficult to quantify and were not available for the study sites. Such "missing values" include e.g., data on coarse woody debris, understory growth and respiration, respiration of reproductive organs, C losses through herbivory, exudation from roots, transfer of C to mycorrhiza and emission of volatile organic compounds or methan. These fractions may account for 11-20% of total NPP in tropical forests (Luyssaert et al. 2007), but were omitted from our C balance approach.

Structure of results

The standing biomass stocks of the three sites are summarized in Table **5.1**. Annual C fluxes and flux components are given in Table **5.2**. Differences in total CO_2 influx (A_C) and total ecosystem CO_2 efflux (R_{total}) when based on the three different LAI are presented in Table **5.3**. Using different LAI would result in changes of items (1), (6), (11), and (29) to (41) of Table **5.2**. Figure **5.1** showed simplified diagrams of the annual C balance of the three study sites (based on Table **5.2**). Rather than presenting precise values, the flux diagrams aim to illustrate altitudinal changes and to demonstrate principle differences between the three diverging montane forest systems.

	1050 m	1890 m	3050 m
Above ground:			
Tree foliage	3.4	4.9	1.8
Tree wood	139.2	81.6	54.3
Total AG	142.6	86.5	56.1
Below ground:			
Coarse roots	14.7	10.0	26.0
Fine roots	1.3	2.9	5.6
Total BG	16.1	13.1	31.4
SOC	48.9	68.0	126.2
Total tree biomass	158.6	99.4	87.5
Root/total biomass	10%	13%	36%

Table 5.1. Estimated stocks of C (Mg ha⁻¹) at the three sites at 1050, 1890 and 3050 m elevation. Biomass data from Moser (2008). Stocks of soil organic carbon (SOC) from Iost (2007).

(41) R_{total}/ A_C

		1050 m	1890 m	3050 m
Above ground:	AG			
Leaf area index	LAI	6.0	5.7	2.2
(1) Canopy CO_2 assimilation	A_{C}	3064	2789	919
(2) Production of foliage	P_L	253	248	90
(3) Production of reproductive organs	P _{rep}	46	18	4
(4) Production of twigs	P_{B}	57	44	20
(5) Production of wood	P_{W}	141	44	7
(6) Dark respiration of foliage	R_L	605	540	389
(7) Respiration of reproductive organs	R _{rep}	nk	nk	nk
(8) Respiration of twigs and branches	R _B	167	103	38
(9) Respiration of stems	R _s	167	103	38
(10) Tree AG-NPP $(2)+(3)+(4)+(5)$	ΔAG	496	354	120
(11) Total AG-C-efflux (6)+(7)+(8)+(9)		939	745	464
(12) Total AG-C-consumption (10)+(11)		1435	1099	584
(13) Transport to roots (1)-(12)	Т	1629	1690	335
Below ground:	BG			
(14) Production of coarse roots	PCP	9	12	45
(15) Production of fine roots	PER	114	149	470
(16) Respiration of coarse roots		41	19	37
(17) Respiration of fine roots	REP	536	144	30
(18) Tree BG-NPP $(14)+(15)$		123	161	520
(19) root respiration (16)+(17)	R _R	576	164	67
Sell.				
Soll: (20) total soil respiration	D	1202	037	282
(20) total soli respiration (21) beterothrophic respiration (20) (22)	R _{soil}	767	932 788	363
(21) neterotinophic respiration (20)-(22) (22) root ($<2mm$) contribution ($\%$ of (20))	R _h	/0/	/00	555
(22) find root litter	т	41	140	8 470
(24) approx root litter	L _{BG}	0	149	470
(24) coarse root inter	L _{BG}	255	210	45
(25) total the fine litter	L _{AG}	333	107	115
(20) non-tree line litter (27) SOM shares $(22) + (24) + (25) + (26) + (21)$		33 255	211	14
(27) SOM change $(25)^+(24)^+(25)^+(20)^-(21)$ (28) Total PG C afflux (10)+(21)	ΔSOM	-233	-211	200 420
(28) Iotai BG-C-eintux (19) $+(21)$		1544	931	420
(29) Autothrophic respiration (11)+(19)	R _a	1516	909	531
(30) Tree NPP $(10)+(18)$		618	515	640
(31) Tree NPP (1)-(29)		1548	1880	388
(32) Tree GPP (30)+(29)		2134	1423	1171
(33) Ecosystem respiration (21)+(29)	R _{total}	2283	1696	884
(34) Difference $(1)-(30)+(33)+(26)$		128	471	-619
or (1)-(31)+(21)+(26)		714	14	164
$(35) \% R_a \text{ of } A_C$		49	33	58
$(36) \% R_h \text{ of } A_C$		25	28	38
$(37) \% R_W \text{ of } A_C$		11	7	8
(38) % NNP of A _C (32)/(1)		20	18	70
or (1)-(31)/(1)		49	33	58
(39) % root transfer of $A_{\rm C}$ (15)/(1)		53	61	36
(40) % RL of A _c		20	19	42

0.75

0.61

Table 5.2. Annual C fluxes (g m⁻² yr⁻¹) at the three study sites at 1050, 1890 and 3050 m. Data sources are given in the text. SOM = soil organic matter, NPP = net primary production.

0.96

5.3 The C balance of three tropical montane forests

Canopy CO₂ assimilation

Canopy CO₂ assimilation (A_C) was considerably lower at 3050 m than at the other two sites (item (1) in Table **5.2**). Including foliage respiration to estimate gross photosynthesis (A_C + R_L) resulted in values very similar to values of gross photosynthesis found for humid boreal forests (653-999 g C m⁻² yr⁻¹) (Luyssaert et al. 2007). However, the measured A_{max} of sunlit leaves did not decline significantly along the gradient, indicating that the capacity of single leaves to assimilate C is not the major constraint of C gain at high altitudes. Instead, annual A_C at 3050 m appeared to be restricted by a combination of several limitations, such as a low LAI, light availability due to more frequent cloudiness (Bendix et al. 2007, Leuschner et al. 2007, Moser 2008).

Depending on the LAI used for calculation, A_C continuously declined with increasing elevation (items (B) in Table **5.3**) or showed little changes from 1050 to 1890 m ((A) and (C)). Estimates of gross photosynthesis ($A_C + R_L$, Table **5.2**) based on the mass-related LAI (A)were in the range of values on gross photosynthesis reported for tropical lowland forests (Malhi et al. 1999: 3040 g C m⁻² yr⁻¹; Luyssaert et al. 2007: 3145 and 3735 g C m⁻² yr⁻¹). LAI of 5-6 are typical for tropical lowland forests (Ryan et al. 1994, Malhi et al. 1999, Meir and Grace 2002). Chambers et al. (2004) even reported a LAI of only 4.7 for a tropical lowland forest has been found to result in large uncertainties when extrapolating photosynthetic CO₂ assimilation or leaf dark respiration to the stand (Cavaleri et al. 2008). However, the comparatively high A_C estimated for the two TMF sites at 1050 and 1890 m most likely derived from the theoretical assumptions used to calculate the cumulative assimilatory C gain of sunlit and shaded leaves in the complex, multilayered canopy strata of tropical forest trees.

Net primary production (NPP)

More surprising than the decline in A_C with increasing elevation was the contrasting result with respect to NPP along the elevational transect. Previous studies based on direct measurements of above- and belowground biomass increments have suggested that NPP remains largely constant (30) with increasing elevation from 1050 to 3050 m (Moser 2008), while our estimates based on A_C and R_a suggest that NPP declines considerably from 1050 to 3050 m (31). Moser (2008) has found an enormous shift in C allocation from predominately

Table 5.3. Canopy carbon gain (A_C) and ecosystem respiration (R_{total}) based on differently determined leaf area indices (LAI; A, B, C) or mean LAI (D) for the three study sites at 1050, 1890 and 3050 m. The differences between both C fluxes represent the ultimate gap in the C balance (Δ Flux) of the three study sites. See text for data sources. Units are g m⁻² yr⁻¹and m⁻² (LAI).

	1050 m	1890 m	3050 m
(A) Leaf-mass related:			
LAI	6	5.7	2.2
A _C	3064	2789	919
R _{total}	2283	1696	884
Δ Flux	781	1093	35
(B) LAI-2000 (Li-Cor):			
LAI	5.1	3.9	2.9
A _C	2604	1908	1211
R _{total}	2192	1526	1007
Δ Flux	412	382	204
(C) Hemispherical photographs:			
LAI	2.8	3.0	2.2
A _C	1430	1468	919
R _{total}	1960	1440	884
Δ Flux	-530	28	35
(D) Mean (A)-(C):			
LAI	4.6	4.2	2.4
A _C	2349	2055	1002
R _{total}	2141	1554	919
Δ Flux	208	501	83

aboveground production at 1050 m towards a huge annual increment of coarse and fine root biomass at 3050 m. However, our flux measurements indicate that this enormous belowground biomass production at the upper montane site (3050 m, Moser 2008) may be an overestimation. Basically, the amount of C available for maintenance and growth of the root system can be derived from A_C and the aboveground C fluxes, with the difference between both terms assumed to be available for the belowground allocation (13). Carbon transfer to the roots was 45% in tropical lowland forests (Malhi et al. 1999). At our study site at 3050 m, this translocation of carbon accounted for 36% (39), which corresponds to 335 g C m⁻² yr⁻¹ (13). In contrast, directly measured root production was 520 g C m⁻² yr⁻¹ (18) (Moser 2008), and total root respiration accounts for additional 67 g C m⁻² yr⁻¹ (19), indicating a gap of 252 g C m⁻² yr⁻¹ between C demand and supply. Additional C demanding processes such as rhizodecomposition, root exudation and the C transfer to mycorrhizae not considered here would even augment the C gap. Moreover, rates of fine root respiration and root contribution to soil respiration declined with increasing altitude, while the opposite would be expected if the transfer of C to below the ground was to increase with elevation (Iost 2007). On the other hand, when A_C and R_L were based on the higher LAI derived from LAI-2000 (B, Table **5.3**), carbon delivery from aboveground amounted to 504 g C m⁻² yr⁻¹ and would hence match the C demand of belowground processes slightly better (data not shown).

The fine root productivity at the three study sites was quantified by means of minirhizotrons (Moser 2008). At 3050 m, some of the measurements may have been taken before a new equilibrium was reached, since the time needed to establish steady state is especially long at the upper montane site due to unfavourable growth conditions (i.e., low temperature, periodically waterlogged soils, low nutrient availability). This could have resulted in an overestimation of the fine root productivity at 3050 m (G. Moser, personal communication). Furthermore, extrapolating root growth from the uppermost soil layer to the total soil profile may have overestimated fine root productivity (Moser et al. submitted).

Gaps in the C balance and implications

Considering the two major fluxes A_C and R_{total} , the upper montane site (3050 m) was close to equilibrium (Table **5.2**; (A) and (C) in Table **5.3**). A close to steady state condition may indicate that at 3050 m the flux-calculated NPP (31) may be more robust than the directly measured one (item (30); Figure **5.1**). Compared to the site at 3050 m, differences between A_C and R_{total} at 1050 and 1890 m were much higher, ranging between -37% (i.e., efflux is greater than influx) and 25% of A_C at 1050m and between 2 and 39% at 1890 m dependent on LAI. Using the averaged LAI seemed to be most robust in terms of a balanced C exchange for all three study sites (Table **5.3**).

Negative changes in the SOM stock at 1050 m and 1890 m (27) indicate that R_h exceeded annual litter production. Soil respiration at 1050 m is indeed within the upper range of values reported for the humid tropics (Iost 2007). At 1890 m, litter accumulation was visibly higher than at 1050 m (personal observation), so that the calculated depletion could be the result of missing litter data (e.g., coarse woody debris). At 3050 m the calculated changes in SOM fairly agree with the observed accumulation of SOM at the upper montane forest mainly resulting from lower temperatures and the lack of primary decomposers (Maraun et al. 2008).



Figure 5.1. Estimated annual carbon fluxes for the three study sites at 1050, 1890 and 3050 m elevation. All units are $g C m^{-2} yr^{-1}$. Abbreviations are given in Table 5.2.

Conclusions

All three study sites are considered intact, old-growth forests, which are assumed to be in C flux equilibrium. Consequently, C efflux should more or less reflect C influx. While at 3050 m this assumption was almost matched, values calculated for the two lower sites implied a large C sink (Figure 5.1). Apart from the mentioned sources of errors (e.g., extrapolation, LAI), the missing values of several components of the C balance (see above) and uncertainties in the quantity of other components urgently require additional and more exact observations, such as on fine root production, branch wood production and respiration, daytime and night-time respiration of sunlit and shaded leaves, and, most importantly, on the canopy CO₂ fixation. Establishing complete C balances is challenging and the (artificial) gaps to close in C balances can be substantial; ranging in magnitude between 10-60% of A_C for different studies and forest biomes (Luyssaert et al. 2007). A balanced C budget of a forest is the equilibrium between NPP and R_h, since changes in the former results in changes in the latter after a certain time lag (Malhi and Grace 2000). However, on time scales shorter than the carbon residence times of biomass or SOM, an increase in one of the two fractions (higher nutrient input and favourable PAR increases NPP, or increasing soil temperature increases R_h) will result in either a source or sink character of a given stand, especially in stands of small sizes. The magnitude of uncertainties in the balance may be related negatively to plot size, which is small in the case of our study when compared to the extent of ecosystems a C balance is usually referred to. Interannual fluctuations in the regional climatic or the use of data derived from different years may have also caused uncertainties in the annual net C flux of the investigated TMF ecosystems. In this context, short-time variability in the components of TMFs' C balance as a result of local climate variation still needs to be resolved if we want to assess the possible impact of global climate changes on C fluxes in TMFs or the degree of the expected C fertilization effect of increasing atmospheric CO₂ on NPP.

5.4 References

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Synthesis

6.1 Discussion of hypotheses

In Chapter 1, a number of hypotheses were introduced concerning altitudinal patterns in the CO_2 release from above- and belowground plant organs to be tested in this study. The present section aims to discuss these hypotheses with respect to the previously described results and recent findings. The four hypotheses were as follows:

- Apparent CO₂ release from stems increases with increasing elevation from 1050 m to 3050 m.
- (ii) Apparent CO₂ release from coarse roots increases with increasing elevation from 1050 m to 3050 m.
- (iii) Apparent CO_2 release rates from stems and coarse roots of tropical montane forests (TMF) at different elevation levels remain constant through the year and do not show seasonality.
- (iv) With increasing elevation, woody tissue respiration shows an increasing relevance in the carbon balance of TMFs. At high elevations, maintenance respiration from stems and growth respiration from roots are predominant.

Previous intensive studies on above- and belowground biomass production found that with increasing elevation an increasing proportion of the available carbon (C) was translocated below the ground. While the aboveground annual biomass increment decreased 4.2-fold from 1050 to 3050 m, the annual C allocation to the root system increased 4.2-fold from the premontane (1050 m) to the upper montane forest (3050 m). This enormous shift in the relative importance of above- to belowground plant organs should have been reflected in the respiratory activity of the respective plant organs. Mean annual rates of stem respiration (based on the bark surface area) declined 4.5-fold from 1050 to 3050 m, while coarse root respiration decreased only by a factor of 2.2. The dissimilar decline in stem and coarse root respiration with increasing elevation resulted in a very similar level of respiratory activity of stems and coarse roots at the upper montane site, indicating the higher relative importance of belowground organs at higher altitudes (Chapter 2). Since absolute root biomass increased with increasing altitude, the extrapolation of coarse root respiration to the stand finally yielded a similar amount of C losses from the coarse root biomass at 1050 and 3050 m. Annual stem wood C losses at the stand-level, in turn, showed the same 4.5-fold decline as found at tree-level as a result of the parallel decline in the aboveground biomass with increasing altitude. We therefore reject hypothesis (i) and accept hypothesis (ii).
Hypothesis (i) was initially based on the assumption of a stress-induced compensatory CO_2 release counterbalancing an assimilatory carbon gain which was assumed to be similar to the CO_2 fixation at lower altitudes. In fact, annual C gain declined from 1050 to 3050 m. On the other hand, this decline was mainly a function of the lower leaf area index (LAI) at high elevations accompanied by several less favourable abiotic conditions (e.g., low light intensity, temperature, nutrient availability), whereas the photosynthetic capacity of single leaves did not change significantly along the elevation transect (Chapter 5). At 3050 m, a proportionally high C allocation to below the ground despite of a lower C gain compared to the lower-elevated sites can only be guaranteed at the expense of aboveground growth and hence aboveground respiration. We found no indication for a generally higher aboveground respiration at 3050 m as a result of abiotic stress throughout the year. However, the role of abiotic factors in driving the respiratory CO_2 release from tree stems could not be resolved satisfactorily (Chapter 3) - leading to hypothesis (iii).

Respiration of a certain tissue type is commonly expected to increase strongly with increasing temperature. This response is more strongly related to variations in tissue temperature than to changes in ambient air temperature (Chapter 3). In a prevailing humid climate of TMFs, where seasonality is usually less pronounced and mean monthly air temperature, and hence tissue temperature fluctuates little throughout the year, we hypothesized that the respiratory activity of stems and coarse roots remain more or less constant through the year. Taking advantage of an exceptional dry season period within the 1year-measurement cycle we could show clear seasonal patterns in the respiratory CO₂ release of stems, though not of coarse roots, along the elevation gradient (Chapter 4). Highest rates of stem respiration were found under the drier and warmer conditions of the dry season period at all three study sites, whereas measurements during the more humid periods yielded much lower stem CO₂ release rates. Root respiration was less responsive as a result of minor fluctuations in the abiotic conditions (temperature, relative humidity, vapour pressure deficit) at the buffered environment of the soil surface where coarse roots were measured. Based on our findings we can prove a climate-sensitivity of stem respiration along the elevation transect that leads us to reject hypothesis (iii).

On the other hand, higher rates of stem respiration could not be simply related to temperature alone, rather we proposed that a complex interaction of several climate factors (light incidence, vapour pressure deficit, wind speed, soil water availability) might play a role in controlling plant respiratory activity (Chapter 3). Even more challenging is the remaining uncertainty with respect to the underlying mechanisms driving stem respiration to augment

under dry season conditions. Stem cell growth in various tree species of the study area has been found to cease when soil water availability drops below a certain threshold during the dry season. In this context, we were not able to conclusively relate higher rates of stem respiration to higher cambial activity. On the contrary, the possibility exists that the higher respiratory activity might indeed be stress-induced resulting in an increasing release of CO_2 via the less productive alternative pathway of respiration, with plant growth being unaffected. A decline in net primary production despite higher CO_2 efflux under higher temperatures has been frequently observed for tropical lowland forests during the last decades. Higher rates of CO_2 release without implications for the biomass increment under warmer climate conditions would have marked consequences for the atmospheric CO_2 accumulation over tropical biomes under the expected global climate change.

Based on hypothesis (i) and (ii), we initially assumed that the shift in the relative importance of above- and belowground plant components with elevation should be reflected in the carbon balance of the TMFs and hypothesized that at higher elevations, maintenance respiration from stems and growth respiration from roots are predominant. We indeed found the fraction of coarse root growth respiration to increase markedly from 1050 to 3050 m (Chapter 4). On the other hand, the proportional partitioning between stem growth and stem maintenance respiration changed only slightly with increasing altitude, resulting in comparatively low fractions of stem growth respiration at the three sites than values found in the literature. Considering the entire C balance of the three TMFs, a shift in the relevance of above- versus belowground plant organs could not be found (Chapter 5). The proportion of C available for the belowground components (i.e., the proportion of C allocated to the roots as percentage of total CO₂ fixation) did not increase with increasing elevation as it would be expected in order to maintain a higher below- than aboveground biomass and biomass production at 3050 m. Our first estimate of above- and belowground C partitioning at the three TMFs indicated that the belowground production might have been overestimated to a certain extent, at least at the upper montane site. However, our C balances were not complete and some estimates of C balance components were based on very simple assumptions. In this regard, additional and more exact investigations are still needed until hypothesis (iv) can be conclusively answered, especially with respect to the total assimilatory C gain, the respiratory CO₂ release from twigs and branches or from sunlit and shaded leaves, and the belowground biomass production.

6.2 Concluding remarks

Measuring the respiratory CO_2 release from stems and coarse roots was a convenient tool to reproduce patterns of above- and belowground C allocation of the three Andean moist forests. Furthermore, the organ-level CO_2 efflux measurements could be used to quantify whole tree as well as stand-level C losses; though existing up scaling methods still lack standardization, which hampers comparison of literature data.

The impact of altitude and climatic fluctuations on the respiratory activity of woody plant organs studied at the level of single trees offered additional insight into the diverse response patterns of different tree species in a species-rich tropical forest community. However, the importance of diverging adaptation strategies of species and the affect of species origin warrant further investigation if we are to better understand species co-existence in the current and future structure of highly diverse tropical ecosystems. Such information would be particularly crucial with respect to global warming expecting to force mesic lowland taxa to migrate upwards in the tropics. In this context, tropical montane forests are of particular relevance and the portable chamber technique was shown to provide a suitable framework to assess species-specific patterns of woody tissue CO_2 release to changes in climatic conditions in logistically difficult montane regions.

The highly differing climate-sensitivity found for the measured tree species already indicated that the present forest structure and functionality of tropical montane forests could be strongly affected by global climate change. A better understanding of plant respiratory responses towards warmer conditions is therefore challenging.

Chapter 7

SUMMARY

7.1 Summary

Tropical montane forests (TMF), which account for 21.2% of tropical forests worldwide, are among the least studied ecosystems with respect to their C balance. TMFs are mainly characterized by high atmospheric humidity and frequent cloudiness driving the structure and functionality of these highly diverse ecosystems across a large altitudinal extension. One of the most marked changes with increasing elevation is the shift in C allocation from above- to belowground plant organs leading to a considerable decline in average tree height towards higher altitudes. Information on the key processes controlling the plant internal C use, the assimilatory C gain and the respiratory CO_2 loss from single plant organs, is missing.

The present study aimed to quantify CO_2 losses from above- and belowground woody organs of representative tree species of a tropical montane moist forest in southern Ecuador. Moreover, we wanted to extrapolate rates of woody tissue respiration to the stand level and bring them in the frame of a carbon balance. We used a portable CO_2 measurement system to monitor the respiratory CO_2 release from stems (R_S) and coarse roots (R_R) along a 2000 melevation transect with three study sites at 1050, 1890 and 3050 m a.s.l. In an intensive 1year-measurement period we determined the impact of altitude and of seasonal climate variations on patterns of CO_2 release from woody organs. We found substantial variation in the stem respiratory activity among different species and different tree individuals at all three study sites. Mean R_S declined significantly from the premontane forest (1050 m) to the upper montane site (3050 m). Mean R_R did not change significantly with altitude, though showing a decreasing tendency. The results corroborated the remarkable shift in the relative importance of above- to belowground plant organs with increasing altitude as it has already been found for the biomass allocation patterns along the elevation gradient.

Temperature has long been known to be the most important abiotic driver of plant respiration. In the field, however, a consistent relationship between temperature and respiratory CO_2 efflux is often not found. Comparing dry and wet season patterns in stem CO_2 release we found R_S to be largely uncoupled from changes in the dial temperature regime under humid season conditions. During the dry season, the respiration-temperature relationship was generally stronger, though temperature sensitivity of R_S differed greatly in degree and even in the direction of response among individual trees. Integrating additional influencing abiotic factors (vapour pressure deficit, wind speed and solar radiation) could not enhance the ability to explain the variability of R_S . We assumed maintenance respiration to

dominate under humid conditions unfavourable for photosynthetic carbon gain of the tree, whereas the dry season conditions principally favoured stem respiratory activity, and most likely energy acquisition. Differences in species distribution centres and hence in the level of climatic adaptation of co-existing moist forest tree species could provide an alternative tool to explain diverging temperature responses of R_s during the dry season.

Information about seasonal patterns of woody tissue CO_2 release from evergreen tropical montane trees has received little attention. We found R_S , but not R_R to show a clear seasonality within the measurement year. Highest rates were measured during the dry season, though the increase in R_S could not be simply related to changes in the temperature regime. On the other hand, the high degree of climate sensitivity of R_S of the studied montane forest trees could also indicate C losses via the alternative pathway (cyanide-resistant oxidase), since the higher respiratory activity could not satisfactorily be related to stimulated cell growth. Along the elevation transect, annual carbon efflux from stems decreased from 167.1 g C m⁻² yr⁻¹ at 1050 m to 37.7 g C m⁻² yr⁻¹ at 3050 m, while coarse root carbon release changed little from 1050 m (40.9 g C m⁻² yr⁻¹) to 3050 m (36.8 g C m⁻² yr⁻¹). Stem growth respiration accounted for a comparatively small fraction of total stem respiration at all three sites, whereas coarse root growth respiration was of increasing importance with increasing altitude.

A first C balance of the three study sites confirmed the expected decline in the total canopy carbon gain from premontane to upper montane forests. However, with respect to the altitudinal changes in above- and belowground C allocation patterns, discrepancies between flux measurements and direct biomass assessments emerged. The enormous C translocation to the root system at 3050 m as previously found could not be matched by the amount of C available as derived from C influx and efflux estimates. On the other hand, the C balances were incomplete and uncertainties still exist with regard to the magnitude of important flux components such as the canopy C gain and the canopy respiratory C losses (including branches and twigs).

7.2 Zusammenfassung

Die Bergregenwälder der Tropen bedecken etwa 21.2 % der Fläche tropischer Wälder weltweit; im Hinblick auf ihre Kohlenstoff-(C)-Bilanz gehören sie dennoch zu den am wenigsten untersuchten terrestrischen Ökosystemen. Die Struktur und Funktionalität tropischer Bergregenwälder wird über weite Flächen- und Höhenausdehnungen vorwiegend geprägt durch vergleichsweise hohe atmosphärische Feuchtigkeit und häufige

Wolkenbedeckung. Eine der markantesten Veränderungen mit zunehmender Meereshöhe stellt die vermehrte Verlagerung von Kohlenstoff von oberirdischen zu unterirdischen Pflanzenorganen dar, welche in eine beachtliche Abnahme der durchschnittlichen Baumhöhe mit zunehmender Meereshöhe resultiert. Informationen über die Schlüsselprozesse der internen C-Nutzung der Pflanze, der assimilatorische C-Gewinn und der respiratorische C-Verlust einzelner Pflanzenorgane fehlen weitestgehend.

Die vorliegende Arbeit hat zum Ziel, die respiratorischen Kohlenstoffdioxid-(CO₂)-Verluste verholzter ober- und unterirdischer Pflanzenorgane repräsentativer Baumarten in einem tropischen Bergregenwald im Süden Ecuadors zu quantifizieren. Weiterhin sollten C-Abgaberaten verholzter Organe auf Bestandesebene hochgerechnet werden und in eine C-Bilanz eingebunden werden. Mit Hilfe einer tragbaren CO₂-Messapparatur wurde entlang eines 2000-m-Höhentransektes an drei Untersuchungsflächen auf 1050, 1890 und 3050 m üNN die respiratorische CO₂-Abgabe von Stämmen und Grobwurzeln gemessen. In einer intensiven 1-Jahres-Messkampagne wurde der Einfluss der Meereshöhe und der saisonalen Klimaunterschiede auf die respiratorische Aktivität der verholzten Organe untersucht. An allen drei Untersuchungsflächen wurde eine beträchtliche Variabilität in der Stammatmung zwischen den verschiedenen Baumarten und zwischen einzelnen Individuen gefunden. Die mittlere Stammatmung nahm mit zunehmender Höhe vom prämontanen (1050 m) zum hochmontanen (3050 m) Standort signifikant ab. Die mittlere Grobwurzelatmung verringerte sich mit zunehmender Meereshöhe zwar tendenziell, diese Veränderung war jedoch nicht signifikant. Somit bestätigen die vorliegenden Ergebnisse eine Verschiebung der relativen Wichtigkeit von oberirdischen zu unterirdischen Pflanzenorganen mit der Meereshöhe, wie sie bereits in vorangegangenen Studien zur Verlagerung der Biomasse mit zunehmender Meershöhe gefunden wurde.

Es ist seit langem bekannt, dass die Respiration der Pflanze vor allem durch die Temperatur als grössten abiotischen Einflussfaktor bestimmt wird. Jedoch konnte gerade im Feldversuch häufig keine konsistente Beziehung zwischen Temperatur und Atmung bestätigt werden. In einer vergleichenden Studie der Stammatmung unter Regen- und Trockenzeitbedingungen auf den Untersuchungsflächen wurde während der Regenzeit eine weitestgehende Entkopplung der Atmung von Veränderungen im täglichen Temperaturverlauf festgestellt. Unter Trockenzeitbedingungen war der Temperaturbezug der Atmung generell stärker ausgeprägt, jedoch fanden sich grosse Unterschiede in der Temperatursensitivität der einzelnen Baumindividuen im Hinblick auf Ausprägung und Richtung der Temperaturantwort. Auch unter Einbezug weiterer einflussnehmender abiotischer Faktoren (Sättigungsdefizit der Luft, Wind, Strahlung) konnte die Variabilität der Stammatmung nicht erklärt werden. Es wurde daher schlussfolgernd angenommen, dass unter den strahlungsextensiven und damit assimilationshemmenden Witterungsverhältnissen der Regenzeit die Erhaltungsatmung dominiert. Dahingegen scheint unter strahlungsgünstigen, assimilationsfördernden Witterungsbedingungen und daher der Trockenzeit die respiratorische Aktivität angeregt zu werden. Weiterhin könnten Unterschiede im Verbreitungsschwerpunkt einzelner Arten und daraus folgende Unterschiede in Adaptionsmechanismen einen alternativen Erklärungsansatz für die grosse Variabilität in der Temperaturantwort der Atmung co-existierender Baumarten unter extremen Trockenzeitbedingungen bieten.

Saisonale Unterschiede in der CO₂-Abgabe von verholztem Gewebe tropischer immergrüner Baumarten wurde bislang wenig untersucht. Im Verlauf der einjährigen Messkampagne wurde eine stark ausgeprägte Saisonalität der Stammatmung, jedoch nicht der Wurzelatmung gefunden. Die höchsten Stamm CO₂-Abgaberaten wurden während der Trockenzeit gemessen, jedoch konnten die erhöhten Atmungsraten nicht ausschliesslich auf die höhere Umgebungstemperatur zum Messzeitpunkt zurückgeführt werden. Vielmehr könnte die erhöhte Atmungsaktivität auch einen Hinweis auf C-Verluste über die alternative Oxidase (cyanidresistente Oxidase) liefern, da die erhöhte Respiration nicht überzeugend mit stimuliertem Zellwachstum in Verbindung gebracht werden konnte. Entlang des Höhentransektes nahm die jährliche C-Abgabe der Stämme von 167.1 g C m⁻² yr⁻¹ auf 1050 m zu 37.7 g C m⁻² yr⁻¹ auf 3050 m ab, wohingegen sich die jährliche C-Abgabe der Grobwurzeln im Höhenverlauf von 1050 m (40.9 g C m⁻² yr⁻¹) auf 3050 m (36.8 g C m⁻² yr⁻¹) wenig änderte. Die Wachstumsatmung der Stämme nahm nur einen vergleichsweise geringen Anteil der Gesamtatmung ein, wohingegen die Wachstumsatmung der Grobwurzeln mit zunehmender Meereshöhe an Wichtigkeit gewann.

Eine erste Gesamt-C-Bilanz der drei Bergregenwald-Standorte bestätigte die erwartete Abnahme des assimilatorischen C-Gewinns von der prämontanen zur hochmomtanen Untersuchungsfläche. Im Hinblick auf die Veränderungen in der ober- und unterirdischen C-Verlagerung mit zunehmender Meereshöhe zeigten sich jedoch Diskrepanzen zwischen der direkten Bestimmung der Biomasse und der indirekten Berechnung über gemessene C-Flüsse. Die enorme C-Verlagerung zum Wurzelsystem, wie sie in vorausgegangen Studien ermittelt wurde, konnte quantitativ nicht durch die ein- und ausgehenden C-Flüsse bestätigt werden. Jedoch ist zu beachten, dass die aufgestellten C-Bilanzen nicht vollständig sind und dass weiterhin grosse Unsicherheiten bestehen im Hinblick auf die Quantität der C-Flüsse wichtiger Komponenten wie etwa der C-Gewinn und –Verlust des Kronendaches (inklusive der Äste und Zweige).

7.3 Resumen

Los bosques tropicales montanos (TMF), cubren el 21.2% de los bosques tropicales a nivel mundial, sin embargo son uno de los ecosistemas menos investigados especialmente con respecto a su balance de carbono. TMFs se caracterizan primeramente por la alta humedad atmosférica y la frequente nubosidad costituyendo los factores que determinan la estructura y funcionamiento de este ecosistema de diversidad complexa y amplia distribución altitudinal. Uno de los cambios más remarcables con el aumento de altitud es la redistribución de C entre los órganos aéreos y subterráneos, conduciendo a una reduccción considerable en la estatura de los árboles que crecen a mayores elevaciones. En general, existe muy poca información sobre los procesos clave en las plantas que controlan la asimilación de C y su emisión a través de la actividad respiratoria de diferentes órganos. El presente estudio tuvó como objetivo la cuantificación de las pérdidas de CO₂ a través de los troncos (R_S) y de las raíces gruesas (R_R) de especies de árboles representativas del bosque tropical montano húmedo del sur de Ecuador. Para las medidas de CO₂ usamos un sistema portátil de intercambio de gases y monitoreamos las emisiones de CO₂ de troncos y raíces gruesas a lo largo de un transecto de 2000 m-de elevación y en tres localidades ubicadas a 1050, 1890 y 3050 m s.n.m. Durante el periodo de un año, llevamos a cabo medidas intensivas para determinar el impacto del cambio altitudinal y variación climática estacional en los patrones de emisión de CO₂ de diferentes órganos leñosos.

En cada una de las tres localidades estudiadas encontramos considerable variación en la actividad respiratoria de troncos entre las diferentes especies e individuos seleccionados. La tasa promedio de emisión de CO_2 de los troncos (R_S) disminuyó significativamente del bosque premontano (1050 m) al bosque montano alto (3050 m). La tasa promedio de emisión de CO_2 de las raíces gruesas (R_R) no mostro variación altitudinal significativa, sin embargo se pudo observar una tendencia decreciente. Los resultados de este estudio corroboran el cambio remarcable en la importancia relativa de los órganos aéreos con respecto a los órganos subterráneos de árboles del bosque tropical montano con el aumento altitudinal. Tal cambio ha sido reportado previamente mediante observaciones sobre los patrones de distribución de biomasa a lo largo del mismo gradiente altitudinal.

La temperatura ha sido siempre identificada como el factor abiotico más importante responsable por la variación en las tasas de respiración. Sin embargo, en el campo, no siempre es possible encontrar una relación consistente entre la temperatura y las tasas de emisión de CO₂. Mediante la comparación de los patrones de emisión de CO₂ de los troncos durante las estaciones húmeda y seca, observamos que Rs no respondió a los cambios del regimén diario de temperatura bajo las condiciones de la estación húmeda. Durante la estación seca, la relación repiración-temperatura fue más fuerte, aunque la sensibilidad de R_S fué altamente variable en magnitud e incluso en la dirección de la respuesta. Aún con la adición de información sobre la variación de otros factores abióticos, (deficit de presión de vapor de agua, dirección del viento y radiación solar), no se logro obtener una mejor estimación para los valores de R_S. Bajo condiciones húmedas, poco favorables para la actividad fotosintética de asimilacion de carbono, se puede asumir que la respiración de mantenimiento domina la tasa total de emisión de CO2. Mientras que condiciones secas con alta radiación y temperatura favorecen la actividad respiratoria y adquisición de energía. Diferencias en los centros de distribución y consecuentemente en la adaptación climática de especies arbóreas podrían proporcionar una herramienta alternativa para explicar la divergencia en la respuesta de R_S a las variaciones en temperatura durante la época seca.

En las investigaciones de especies arbóreas del bosque tropical montano, se ha dado poca atención al estudio de los patrones estacionales de emisión de CO_2 de la biomasa leñosa. En este estudio y durante el periodo de un año de observacion se pudo identificar una clara variabilidad estacional en R_S, pero no en R_R. Las tasas más altas de emisión de CO_2 fueron registradas durante la época seca, aunque el aumento en R_S no pudo ser relacionado únicamente a cambios en temperatura. Por otro lado, el alto grado de sensibilidad climática de R_S de los árboles del bosque montano estudiado podrían indicar pérdida de C por medio de ciclos vanos, ya que la tasa más alta de actividad respiratoria no pudo ser relacionada satisfactoriamente con una estimulación del crecimiento a nivel celular. A lo largo del transecto altitudinal, la emisión anual de carbono de los troncos disminuyó de167.1 g C m⁻² a⁻¹ a 1050 m a 37.7 g C m⁻² a⁻¹ a 3050 m, mientras que la emisión anual de las raíces gruesas se mantuvo constante entre 1050 m (40.9 g C m⁻² a⁻¹) y 3050 m (36.8 g C m⁻² a⁻¹). La respiración de los troncos en las tres elevaciones. A nivel de raíces gruesas, la fracción correspondiente a la respiración de crecimiento mostró un incremento con la altitud.

Un primer balance de C de los tres sitios confirmó una disminución en la tasa neta de asimilación de carbono del bosque premontano al bosque montano alto. Sin embargo, con

respecto a los cambios altitudinales en los patrones de distribución de C entre la biomasa aérea y subterránea, diferentes discrepancias emergieron entre las estimaciones de flujo de C basadas en medidas de asimilación de CO_2 y las estimaciones basadas en medidas de biomasa. La asignación de grandes cantidades de carbono al sistema radicular a 3050 m observada en previas investigaciones, no pudo ser explicada por la cantidad de C disponible de acuerdo a nuestras estimaciones del balance entre el flujo de C. Por otro lado, se debe tener en cuenta que nuestro balance de C no esta completo ya que todavía se tienen muchas incertidumbres sobre la magnitud e importancia de componentes importantes del flujo de C, como es el caso de la asimilación y pérdida de C en el dosel (incluyendo ramas gruesas y finas).



APPENDIX

Appendix 1. Species list of the measured tree stems at 1050, 1890 and 3050 m a.s.l. Some trees could not be determined to species level (indet). Given are DBH (cm) at the beginning (August 2005) and end (September 2006) of the measurement campaigns, tree height (tree, m; G. Moser) and clear bole length (bole, m; own measurements).

Abbr.	Species	Family	DBH	DBH	tree	bole
	•	·	2005	2006		
	1050m					
An1	indet	Annonaceae	10.05	10.47	9.90	5.8
An2	indet	Annonaceae	21.37	21.53	14.50	9.2
Sc1	<i>Schefflera</i> sp.	Araliaceae	27.67	27.74	18.80	9.6
Pou	Pourouma cf.	Cecropiaceae	14.32	14.65	13.40	9.9
Al	Alchornea sp.	Euphorbiaceae	8.75	9.3	13.30	5.7
La1	indet	Lauraceae	13.66	13.67	12.30	9.9
La2	indet	Lauraceae	17.07	17.07	11.90	10.7
Mp1	Miconia punctata	Melastomataceae	9.8	-	14.04	-
Mp2	Miconia punctata	Melastomataceae	11.93	12.03	10.51	10.5
Mp3	indet	Melastomataceae	15.35	15.53	11.80	5.2
In	<i>Inga</i> sp.	Mimosaceae	16.34	16.41	17.80	10.5
Mi	indet	Mimosaceae	10.08	10.1	18.10	13.3
Fi1	Ficus sp.	Moraceae	14.16	-	16.20	5
Fi2	Ficus sp.	Moraceae	14.64	14.64	14.20	9.9
Fi3	Ficus sp.	Moraceae	16.03	16.43	14.20	13.1
Fi4	Ficus sp.	Moraceae	43.85	43.99	29.90	11.85
Vi	Virola cf.	Myristicaceae	14.17	14.17	15.00	12
Myt	indet	Myrtaceae	10.96	11.1	11.20	8.95
Ch	Chrysophyllum sp.	Sapotaceae	8.84	9.08	11.20	7.6
Po1	<i>Pouteria</i> cf.	Sapotaceae	20.79	20.78	16.80	11.1
Po2	Pouteria cf.	Sapotaceae	43.72	43.72	26.60	15
	1000					
	1890 m	A 'C 1'	0.62	0.64	7.50	F 7 F
la	Ilex amboroica	Aquitoliaceae	9.63	9.64	7.50	5.65
Sc2	Schefflera sp.	Araliaceae	9.16	-	14.55	6.70
Cr	Clethra revoluta	Clethraceae	23.45	23.65	12.69	4.98
Hm	Hyeronima moritziana	Euphorbiaceae	9.52	9.53	13.70	8.24
Eo	Endlicheria oreocola	Lauraceae	20.8	20.82	10.00	6.17
Ne	Nectandra sp.	Lauraceae	11.64	11.71	12.45	1.60
Oa	Ocotea aciphylla	Lauraceae	10.21	10.26	10.50	6.43
Gel	Graffenrieda emarginata	Melastomataceae	24.91	25	17.40	10.54
Ge2	Graffenrieda emarginata	Melastomataceae	26.11	26.1	16.54	8.72
Ge3	Graffenrieda emarginata	Melastomataceae	26.47	26.5	19.82	8.18
Mp4	Miconia punctata	Melastomataceae	10.62	10.64	14.05	6.28
Mp5	Miconia punctata	Melastomataceae	12	12	7.06	12.14
Mp6	Miconia punctata	Melastomataceae	14.45	14.56	12.53	8.84
Mc	Myrsine coriacea	Myrsinaceae	3.76	3.82	7.10	4.38
Lo	Ladenbergia oblongifolia	Rubiaceae	-	12.14	9.70	4.10
Pa	<i>Palicourea</i> sp.	Rubiaceae	3.02	-	3.70	2.14
Mi1	Matayba inelegans	Sapindaceae	4.46	4.5	5.60	3.92
Mi2	Matayba inelegans	Sapindaceae	8.17	8.25	8.30	4.20
Mi3	Matayba inelegans	Sapindaceae	10.53	10.67	12.70	6.50
Mg	Micropholis guyanensis	Sapotaceae	8.35	8.46	7.80	4.60

Abbr.	Species	Family	DBH	DBH	tree	bole
	-		2005	2006		
	3050 m					
Iw	Ilex weberlingii	Aquifoliaceae	8.2	8.2	5.72	4.49
He	Hedyosmum sp.	Chloranthaceae	9.36	9.36	4.19	1.81
Cl1	<i>Clusia</i> sp.1	Clusiaceae	17.67	17.67	5.80	1.50
Cl2	<i>Clusia</i> sp.2	Clusiaceae	8.94	9.01	6.79	1.63
Wl	Weinmannia loxensis	Cunoniaceae	10.5	10.5	6.13	3.34
Ce	Cerotostema cf.	Ericaceae	3.6	3.61	3.74	1.93
Er	indet	Ericaceae	15.6	15.6	6.22	1.64
Ax	<i>Axinea</i> sp.	Melastomataceae	8.99	8.99	4.80	2.60
Ma	Myrsine andina	Myrsinaceae	9.55	9.55	7.74	1.27
Mys	indet	Myrsinaceae	3.92	3.92	4.80	2.87
Mo	<i>Monnina</i> sp.	Polygalaceae	2.48	2.58	3.51	2.32
Sf	Styrax foveolaria	Styracaceae	12.06	12.08	3.43	1.71
Sy	Symplocos sp.	Symplocaceae	9.8	9.81	6.68	3.30

Appendix 2. The mobile 6-chamber respiration system ANARESY 2 (Walz, Effeltrich, Germany) with measurement equipment at one of the study sites (1050 m) (top left); tree stem respiration chamber with bark surface temperature sensor on the left hand side of the chamber at the study site at 1890 m (top right); coarse root respiration chamber with bark surface temperature sensor (bottom left). The author and a portable infrared gas analyzer (Li-6400, Li-Cor, Inc.) at 1050 m (bottom right).







Appendix 3b. Diurnal courses of stem CO₂ release (R_s) and the relationship between temperature and carbon release from 3 selected tree individuals during the dry season campaign (D1, left panel) and during a more humid campaign (H3, right panel) at 1890 m. Error bars represent 1 standard deviation. Species abbreviation is given in Appendix 1. Note different scales of the y-axis for the three tree individuals.





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Stems 1050 m 1051 m 5 An1 88.08 0.42 0.81 3.79 0.06 0.30 0.18 An2 91.28 0.50 0.64 3.13 -0.04 0.31 -0.27 Sc1 97.53 0.60 1.18 2.63 0.00 0.29 -0.09 Pou 89.21 0.40 1.33 1.96 0.00 0.70 -0.09 All 77.49 0.37 1.44 2.22 -0.02 0.70 -0.18 Lal 88.91 0.54 0.79 0.37 -0.04 0.65 -0.44 Mp3 79.48 0.49 2.33 2.45 -0.02 0.18 -0.27 In 89.39 0.67 1.24 2.89 -0.05 0.22 -0.44 Mi2 9.33 2.45 -0.02 0.17 -0.09 Fi1 85.96 0.35 2.12 9.16 0.11 0.63 -0.35 <th></th> <th>С</th> <th>Ν</th> <th>Ca</th> <th>K</th> <th>Na</th> <th>Mg</th> <th>Р</th>		С	Ν	Ca	K	Na	Mg	Р
1050 m An1 88.08 0.42 0.81 3.79 0.06 0.30 0.18 An2 91.28 0.50 0.64 3.13 -0.04 0.31 -0.27 Sc1 97.53 0.60 1.18 2.63 0.00 0.29 -0.09 Pou 89.21 0.40 1.33 1.96 0.00 0.70 -0.018 La1 88.91 0.54 0.79 0.71 -0.05 0.07 -0.26 La2 87.82 0.59 2.54 1.65 -0.04 0.14 0.00 Mp1 0.71 -0.05 0.07 -0.26 La2 87.82 0.59 2.54 1.65 -0.02 0.18 -0.27 Im 89.39 0.67 1.24 2.89 -0.05 0.02 -0.44 Mp3 79.48 0.49 2.33 2.45 -0.02 0.19 -0.26 Fi1 85.96 0	Stems							
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	1050 n	n						
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	An1	88.08	0.42	0.81	3.79	0.06	0.30	0.18
Sc1 97.53 0.60 1.18 2.63 0.00 0.29 -0.09 Pou 89.21 0.40 1.33 1.96 0.00 0.70 -0.09 Al 77.49 0.37 1.44 2.22 -0.02 0.70 -0.18 La1 88.91 0.54 0.79 0.71 -0.05 0.07 -0.26 La2 87.82 0.59 2.54 1.65 -0.04 0.14 0.00 Mp1 0.02 0.18 -0.27 In 89.39 0.67 1.24 2.89 -0.05 0.02 -0.44 Mi 94.33 0.65 2.26 4.04 -0.02 0.19 -0.26 Fi1 85.96 0.35 2.12 9.16 0.01 0.63 0.35 Fi2 83.50 0.43 1.97 2.74 0.05 0.28 -0.09 Fi3 88.69 0.38	An2	91.28	0.50	0.64	3.13	-0.04	0.31	-0.27
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	Sc1	97.53	0.60	1.18	2.63	0.00	0.29	-0.09
Al 77.49 0.37 1.44 2.22 -0.02 0.70 -0.18 La1 88.91 0.54 0.79 0.71 -0.05 0.07 -0.26 La2 87.82 0.59 2.54 1.65 -0.04 0.14 0.00 Mp1 0.07 0.65 -0.44 Mp3 79.48 0.49 2.33 2.45 -0.02 0.18 -0.27 In 89.39 0.65 2.26 4.04 -0.02 0.19 -0.26 Fi1 85.96 0.35 2.12 9.16 0.01 0.63 0.38 Fi2 83.50 0.43 1.97 2.74 0.05 0.28 -0.09 0.28 0.09 0.64 -0.03 3.39 -0.18 Fi4 94.83 1.14 1.60 1.30 -0.02 0.77 -0.35 Myt 83.47 0.33 2.09 0.23 0.57 0.35	Pou	89.21	0.40	1.33	1.96	0.00	0.70	-0.09
La188.91 0.54 0.79 0.71 -0.05 0.07 -0.26 La2 87.82 0.59 2.54 1.65 -0.04 0.14 0.00 Mp2 88.61 0.54 0.97 3.37 -0.04 0.65 -0.44 Mp3 79.48 0.49 2.33 2.45 -0.02 0.18 -0.27 In 89.39 0.67 1.24 2.89 -0.05 0.02 0.44 Mi 94.33 0.65 2.26 4.04 -0.02 0.19 -0.26 Fi1 85.96 0.35 2.12 9.16 0.01 0.63 0.35 Fi2 83.50 0.43 1.97 2.74 0.05 0.28 -0.09 Fi3 88.69 0.38 5.09 6.64 -0.02 0.17 -0.09 Vi 89.15 0.64 1.77 1.74 0.00 0.23 -0.35 Myt 83.47 0.33 2.09 2.23 0.05 0.76 -0.27 Ch 90.50 0.58 0.72 3.16 -0.05 0.27 -0.35 Po1 89.49 0.69 0.44 1.43 -0.03 0.14 -0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 </td <td>Al</td> <td>77.49</td> <td>0.37</td> <td>1.44</td> <td>2.22</td> <td>-0.02</td> <td>0.70</td> <td>-0.18</td>	Al	77.49	0.37	1.44	2.22	-0.02	0.70	-0.18
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	Lal	88.91	0.54	0.79	0.71	-0.05	0.07	-0.26
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	La2	87.82	0.59	2.54	1.65	-0.04	0.14	0.00
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	Mpl	00 (1	0.54	0.07	2.27	0.04	0.65	0.44
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Mp2	88.61	0.54	0.97	3.37	-0.04	0.65	-0.44
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Mp3	/9.48	0.49	2.33	2.45	-0.02	0.18	-0.27
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	In M:	89.39	0.67	1.24	2.89	-0.05	0.02	-0.44
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		94.33	0.65	2.20	4.04	-0.02	0.19	-0.20
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	F11 E:2	85.90	0.35	2.12	9.10	0.01	0.03	0.35
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	F12 E:2	83.30	0.43	1.97	2.74	0.05	0.28	-0.09
P1494.831.141.001.30 -0.02 0.17 -0.09 Vi89.150.641.771.740.000.23 -0.35 Myt83.470.332.092.230.050.76 -0.27 Ch90.500.580.723.16 -0.05 0.27 -0.35 Po189.490.690.441.43 -0.03 0.14 -0.09 Po2104.140.820.441.040.010.200.09Sc291.090.300.993.680.030.570.35Cr82.690.210.691.690.030.330.26Hm82.960.260.250.98 -0.03 0.110.27Eo91.760.720.381.73 -0.02 0.150.35Ne91.050.470.251.34 -0.03 0.150.26Ga89.680.410.000.86 -0.04 0.070.35Ge193.890.270.311.490.000.210.26Ge281.860.290.041.300.170.170.18Ge380.820.230.341.200.000.140.00Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09 -0.02 <	F13 E:4	88.09	0.38	5.09	0.04	-0.03	5.39 0.17	-0.18
N1 83.13 0.04 1.77 1.74 0.00 0.23 -0.33 Myt 83.47 0.33 2.09 2.23 0.05 0.76 -0.27 Ch 90.50 0.58 0.72 3.16 -0.05 0.27 -0.35 Pol 89.49 0.69 0.44 1.43 -0.03 0.14 -0.09 Po2 104.14 0.82 0.44 1.04 0.01 0.20 0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.21 0.69 1.69 0.03 0.33 0.26 Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.	Г14 V;	94.83	1.14	1.00	1.30	-0.02	0.17	-0.09
Myt 33.47 0.33 2.09 2.23 0.03 0.70 -0.27 Ch 90.50 0.58 0.72 3.16 -0.05 0.27 -0.35 Po1 89.49 0.69 0.44 1.43 -0.03 0.14 -0.09 Po2 104.14 0.82 0.44 1.04 0.01 0.20 0.09 1890 m Ia 89.96 0.26 0.52 1.92 0.02 0.60 0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.21 0.69 1.69 0.03 0.33 0.26 Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Ga 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.39 0.00 Mp5 91.67 0.45 0.30 1.52 0.00 0.27 0.44 Mp6 91.68 0.45 0.26 2.27 <t< td=""><td>VI Mart</td><td>89.13 82.47</td><td>0.04</td><td>1.//</td><td>1.74</td><td>0.00</td><td>0.25</td><td>-0.55</td></t<>	VI Mart	89.13 82.47	0.04	1.//	1.74	0.00	0.25	-0.55
Ch 90.30 0.33 0.72 3.10 -0.03 0.27 -0.33 Po1 89.49 0.69 0.44 1.43 -0.03 0.14 -0.09 Po2 104.14 0.82 0.44 1.04 0.01 0.20 0.09 1890 m Ia 89.96 0.26 0.52 1.92 0.02 0.60 0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.21 0.69 1.69 0.03 0.33 0.26 Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Oa 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.39 0.00 Mp5 91.67 0.45 0.30 1.52 0.00 0.27 0.44 Mp6 91.68 0.45 0.26 2.27 0.04 0.26 0.18 Mc 82.07 0.28 0.46 1.93 0.00 0.17 $0.$	Ch	00.50	0.55	2.09	2.25	0.05	0.70	-0.27
Po1 89.49 0.09 0.44 1.43 -0.03 0.14 -0.09 Po2 104.14 0.82 0.44 1.04 0.01 0.20 0.09 1890 m Ia 89.96 0.26 0.52 1.92 0.02 0.60 0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.21 0.69 1.69 0.03 0.33 0.26 Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Ga 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.17 0.18 Ge3 80.82 0.23 0.34 1.20 0.00 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.27 0.44 Mp6 91.68 0.45 0.26 2.27 0.04 0.26 0.18 Mc 82.07 0.28 0.46 1.93 0.00 0.17 0.44 Lo 94.88 0.35 0.22 2.09 -0.02 0.15 0.3	CII Do1	90.30	0.38	0.72	5.10	-0.03	0.27	-0.55
HO2 HO4 H	P01 Po2	69.49 104.14	0.69	0.44	1.45	-0.03	0.14	-0.09
1890 m Ia 89.96 0.26 0.52 1.92 0.02 0.60 0.09 Sc2 91.09 0.30 0.99 3.68 0.03 0.57 0.35 Cr 82.69 0.21 0.69 1.69 0.03 0.33 0.26 Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Oa 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.39 0.00 Mp5 91.67 0.45 0.26 2.27 0.04 0.26 0.18 Mc 82.07 0.28 0.46 1.93 0.00 0.17 0.44 Lo 94.88 0.35 0.22 2.09 -0.02 0.15 0.35 Pa 91.61 0.60 1.49 5.81 0.01 1.32 0.35 Mi1 77.70 0.43 0.14 2.96 0.00 0.32 0.35 Mi2 89.77 0.55 1.92 1.48 -0.01 0.57 <t< td=""><td>F02</td><td>104.14</td><td>0.82</td><td>0.44</td><td>1.04</td><td>0.01</td><td>0.20</td><td>0.09</td></t<>	F02	104.14	0.82	0.44	1.04	0.01	0.20	0.09
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	1890 n	n						
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Ia	89.96	0.26	0.52	1.92	0.02	0.60	0.09
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Sc2	91.09	0.30	0.99	3.68	0.03	0.57	0.35
Hm 82.96 0.26 0.25 0.98 -0.03 0.11 0.27 Eo 91.76 0.72 0.38 1.73 -0.02 0.15 0.35 Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Oa 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.17 0.18 Ge3 80.82 0.23 0.34 1.20 0.00 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.39 0.00 Mp5 91.67 0.45 0.30 1.52 0.00 0.27 0.44 Mp6 91.68 0.45 0.26 2.27 0.04 0.26 0.18 Mc 82.07 0.28 0.46 1.93 0.00 0.17 0.44 Lo 94.88 0.35 0.22 2.09 -0.02 0.15 0.35 Pa 91.61 0.60 1.49 5.81 0.01 1.32 0.35 Mi1 77.70 0.43 0.14 2.96 0.00 0.32 0.35 Mi2 89.77 0.55 1.92 1.48 -0.01 0.57 0.35 Mi3 86.68 0.36 0.38 1.38 0.01 0.17 0.53 <td>Cr</td> <td>82.69</td> <td>0.21</td> <td>0.69</td> <td>1.69</td> <td>0.03</td> <td>0.33</td> <td>0.26</td>	Cr	82.69	0.21	0.69	1.69	0.03	0.33	0.26
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	Hm	82.96	0.26	0.25	0.98	-0.03	0.11	0.27
Ne 91.05 0.47 0.25 1.34 -0.03 0.15 0.26 Oa 89.68 0.41 0.00 0.86 -0.04 0.07 0.35 Ge1 93.89 0.27 0.31 1.49 0.00 0.21 0.26 Ge2 81.86 0.29 0.04 1.30 0.17 0.17 0.18 Ge3 80.82 0.23 0.34 1.20 0.00 0.14 0.00 Mp4 91.07 0.44 0.46 3.30 0.00 0.39 0.00 Mp5 91.67 0.45 0.30 1.52 0.00 0.27 0.44 Mp6 91.68 0.45 0.26 2.27 0.04 0.26 0.18 Mc 82.07 0.28 0.46 1.93 0.00 0.17 0.44 Lo 94.88 0.35 0.22 2.09 -0.02 0.15 0.35 Pa 91.61 0.60 1.49 5.81 0.01 1.32 0.35 Mi1 77.70 0.43 0.14 2.96 0.00 0.32 0.35 Mi2 89.77 0.55 1.92 1.48 -0.01 0.57 0.35 Mi3 86.68 0.36 0.38 1.38 0.01 0.17 0.53 Mg 83.25 0.46 0.50 0.85 0.01 0.17 0.53	Eo	91.76	0.72	0.38	1.73	-0.02	0.15	0.35
Oa89.680.410.000.86-0.040.070.35Ge193.890.270.311.490.000.210.26Ge281.860.290.041.300.170.170.18Ge380.820.230.341.200.000.140.00Mp491.070.440.463.300.000.390.00Mp591.670.450.301.520.000.270.44Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.170.53Mg83.250.460.500.850.010.170.53	Ne	91.05	0.47	0.25	1.34	-0.03	0.15	0.26
Ge193.890.270.311.490.000.210.26Ge281.860.290.041.300.170.170.18Ge380.820.230.341.200.000.140.00Mp491.070.440.463.300.000.390.00Mp591.670.450.301.520.000.270.44Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Oa	89.68	0.41	0.00	0.86	-0.04	0.07	0.35
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Ge1	93.89	0.27	0.31	1.49	0.00	0.21	0.26
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	Ge2	81.86	0.29	0.04	1.30	0.17	0.17	0.18
Mp491.070.440.463.300.000.390.00Mp591.670.450.301.520.000.270.44Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Ge3	80.82	0.23	0.34	1.20	0.00	0.14	0.00
Mp591.670.450.301.520.000.270.44Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.170.53Mg83.250.460.500.850.010.170.53	Mp4	91.07	0.44	0.46	3.30	0.00	0.39	0.00
Mp691.680.450.262.270.040.260.18Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Mp5	91.67	0.45	0.30	1.52	0.00	0.27	0.44
Mc82.070.280.461.930.000.170.44Lo94.880.350.222.09-0.020.150.35Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Мрб	91.68	0.45	0.26	2.27	0.04	0.26	0.18
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	Mc	82.07	0.28	0.46	1.93	0.00	0.17	0.44
Pa91.610.601.495.810.011.320.35Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Lo	94.88	0.35	0.22	2.09	-0.02	0.15	0.35
Mi177.700.430.142.960.000.320.35Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Pa	91.61	0.60	1.49	5.81	0.01	1.32	0.35
Mi289.770.551.921.48-0.010.570.35Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Mi1	77.70	0.43	0.14	2.96	0.00	0.32	0.35
Mi386.680.360.381.380.010.280.27Mg83.250.460.500.850.010.170.53	Mi2	89.77	0.55	1.92	1.48	-0.01	0.57	0.35
Mg 83.25 0.46 0.50 0.85 0.01 0.17 0.53	Mi3	86.68	0.36	0.38	1.38	0.01	0.28	0.27
	Mg	83.25	0.46	0.50	0.85	0.01	0.17	0.53

Appendix 4. Contents of carbon and nitrogen and mineral concentrations (mg g^{-1}) of wood cores from the measured tree stems of the three sites at 1050, 1890 and 3050 m elevation. See Appendix 1 for species abbreviation. Coarse roots are described by their diameter (d, cm).

	С	N	Ca	K	Na	Mo	P
3050 m	C	T.A.	Ca	IX	ina	IVIE	1
Iw	93.02	0.27	0.72	1.93	0.04	0.94	0.26
Не	84.64	0.23	0.35	1.62	0.00	0.28	0.26
Cl1	80.22	0.17	0.79	1.72	0.06	0.70	0.26
Cl2	79.57	0.15	0.65	1.30	0.01	0.16	0.26
W1	86.97	0.18	1.04	2.67	0.52	0.72	0.27
Ce	92.22	0.22	0.69	1.61	-0.06	0.33	0.09
Er	84.30	0.16	1.39	2.70	0.01	0.42	0.18
Ax							
Ma	92.05	0.20	0.26	1.54	0.00	0.14	0.18
Mys	90.91	0.20	0.59	1.61	0.05	0.27	0.26
Mo	76.55	0.25	0.31	2.44	0.02	0.18	0.18
Sf	88.94	0.19	0.55	1.25	0.01	0.20	0.26
Sy	86.02	0.31	0.65	1.93	0.02	0.27	0.09
Roots							
d							
1050 m							
1.1	110.3	2.68	7.42	2.16	0.04	0.22	0.53
1.4	122.2	3.71	4.20	3.73	0.08	0.33	0.53
1.7	151.1	2.32	1.06	1.23	0.00	0.50	0.44
2.0	152.4	2.17	1.36	3.86	-0.01	0.37	0.44
2.5	128.0	1.57	11.90	-0.56	0.12	1.50	0.26
2.6	132.3	1.62	9.48	6.48	-0.03	0.49	0.53
3.0	103.4	1.38	8.50	4.09	0.13	0.77	0.53
3.7	136.9	3.26	4.48	1.33	0.00	0.60	0.71
1890 m							
1.3	143.1	3.01	0.72	1.86	0.00	0.33	0.53
1.5	113.9	1.22	1.66	2.59	-0.04	0.58	0.62
1.6	113.6	0.85	0.53	1.85	-0.05	0.42	0.70
2.0	122.7	1.23	0.78	6.34	-0.03	0.66	0.44
2.2	139.7	1.61	0.58	2.02	0.00	0.27	0.53
2.9	134.7	1.45	3.82	3.55	-0.02	0.77	0.44
3.0	122.4	1.54	0.66	2.59	-0.03	0.31	0.53
3050 m							
1.5	114.1	0.97	1.77	4.04	0.02	2.79	0.44
1.8	125.4	1.07	6.01	5.49	0.00	1.02	0.53
2.3	106.9	0.62	1.68	3.99	-0.02	1.04	0.44
2.9	130.4	0.80	0.30	0.74	0.00	0.15	0.18

Appendix 5. Rates of CO₂ release (μ mol CO₂ m⁻² s⁻¹) of the measured tree individuals and coarse roots (SD in parenthesis). Given are mean diurnal rates of CO₂ efflux from each measurement campaigns (H1, D1, H2 and H3). At 1890 m one measurement campaign was done twice, during the first unusual dry part of April 2006 (H2dry) and when more typical rainy weather started two weeks later (H2_{wet}). See Appendix 1 for species abbreviation. Coarse roots are described by their diameter (d, cm).

	H1		D1		H2		H2 _{wat}		H3	
					$(H2_{drv})$		wei			
Stems										
1050 m										
An1	1.02	(0.18)	2.40	(0.47)	0.63	(0.12)			0.25	(0.07)
An2			1.82	(0.13)	0.78	(0.05)			0.69	(0.07)
Sc1	0.79	(0.14)	2.18	(0.37)	0.93	(0.12)			0.80	(0.10)
Pou	0.63	(0.28)	0.22	(0.10)	0.29	(0.08)			0.37	(0.04)
Al			0.41	(0.10)	0.75	(0.13)			0.41	(0.10)
La1	0.27	(0.26)	1.71	(0.33)	0.38	(0.08)			0.28	(0.08)
La2			1.38	(0.25)	0.29	(0.07)			0.19	(0.05)
Mp1	0.90	(0.19)	0.41	(0.14)						
Mp2			1.65	(0.17)	0.80	(0.06)			0.68	(0.15)
Mp3	1.39	(0.13)	0.84	(0.29)	0.49	(0.08)			0.73	(0.13)
In	1.06	(0.19)	0.34	(0.15)	0.68	(0.13)			0.88	(0.09)
Mi	1.19	(0.21)	1.32	(0.18)	0.44	(0.09)			0.58	(0.07)
Fi1	0.44	(0.17)	2.32	(0.51)						
Fi2	0.79	(0.17)	0.49	(0.15)	0.20	(0.06)			0.08	(0.10)
Fi3	0.66	(0.09)	0.54	(0.12)	1.14	(0.08)			1.33	(0.14)
Fi4	1.71	(0.25)	2.71	(0.30)	0.51	(0.09)			0.37	(0.11)
Vi			0.78	(0.18)	0.21	(0.08)			0.20	(0.09)
Myt	1.34	(0.14)	1.00	(0.15)	0.90	(0.06)			1.00	(0.12)
Ch	1.40	(0.15)	1.70	(0.20)	0.31	(0.06)			0.53	(0.07)
Po1	1.17	(0.15)	3.20	(0.38)	0.40	(0.09)			0.67	(0.10)
Po2	0.77	(0.22)	1.60	(0.29)	0.25	(0.08)			0.30	(0.08)
1890 m										
Ia	0.12	(0.20)	0.18	(0.11)	0.40	(0.09)			0.29	(0.06)
Sc2	-0.005	(0.11)	0.11	(0.06)			0.17	(0.17)		
Cr	0.93	(0.15)	0.99	(0.10)	0.65	(0.08)	0.42	(0.10)	0.48	(0.11)
Hm	0.35	(0.10)	0.39	(0.13)	0.38	(0.21)	0.24	(0.08)	0.11	(0.06)
Eo	0.11	(0.08)	0.87	(0.24)	0.44	(0.16)	0.15	(0.07)	0.15	(0.08)
Ne			2.12	(0.33)	1.70	(0.19)	0.32	(0.09)	1.05	(0.11)
Oa	0.53	(0.11)	1.13	(0.27)	1.68	(0.18)	0.37	(0.07)	0.19	(0.09)
Ge1	0.73	(0.13)	0.55	(0.14)	0.51	(0.12)	0.03	(0.07)	0.03	(0.05)
Ge2	0.74	(0.24)	0.61	(0.18)	1.21	(0.14)	0.05	(0.13)	0.19	(0.22)
Ge3	0.22	(0.09)	0.96	(0.08)	0.31	(0.09)	0.14	(0.08)	0.04	(0.03)
Mp4	0.79	(0.33)	1.34	(0.26)	1.05	(0.10)	0.52	(0.07)	0.14	(0.04)
Mp5			0.58	(0.11)	1.02	(0.11)	0.66	(0.11)	0.38	(0.06)
Mp6	0.17	(0.10)	0.91	(0.13)	0.97	(0.13)	0.31	(0.07)	0.23	(0.03)
Mc	0.02	(0.05)	0.09	(0.07)	0.00	(0.04)	0.06	(0.05)	0.06	(0.03)
Lo			0.28	(0.09)	0.60	(0.12)	0.19	(0.08)	0.64	(0.09)
Pa	0.03	(0.07)	0.18	(0.06)	0.17	(0.05)	0.17	(0.07)		
Mi1	0.18	(0.06)	1.02	(0.21)	0.59	(0.11)			0.51	(0.12)
Mi2	2.13	(0.22)	0.85	(0.22)	1.24	(0.17)	0.40	(0.11)	0.91	(0.13)
Mi3	1.19	(0.20)	1.04	(0.34)	1.78	(0.42)	0.64	(0.18)	0.53	(0.07)
Mg	0.42	(0.11)	1.06	(0.11)	0.37	(0.11)	0.29	(0.19)	0.18	(0.06)

	H1		D1		H2		H2 _{wet}		H3	
					$(H2_{dry})$					
3050 m										
Iw			0.22	(0.06)	0.04	(0.03)			0.12	0.02
He			0.22	(0.05)	0.90	(0.08)			0.09	0.02
Cl1			0.23	(0.05)	0.02	(0.02)			0.51	0.02
C12			0.03	(0.04)	-0.04	(0.03)			0.24	0.04
Wl			0.30	(0.05)					0.14	0.02
Ce			0.19	(0.06)					0.20	0.03
Er			0.11	(0.05)	0.05	(0.05)			0.02	0.01
Ax			0.39	(0.06)	0.26	(0.03)			0.28	0.04
Ma			0.11	(0.03)	0.07	(0.03)			0.02	0.01
Mys			0.09	(0.01)	0.09	(0.04)				
Mo			0.19	(0.10)	0.27	(0.04)			0.05	0.02
Sf			0.42	(0.05)	0.27	(0.02)			0.07	0.03
Sy			0.28	(0.04)	0.29	(0.10)			0.18	0.03
Roots										
d										
1050 m										
1.1	0.45	(0.24)	0.37	(0.15)	0.40	(0.18)			0.37	(0.17)
1.4	0.07	(0.32)	0.76	(0.14)	0.23	(0.15)			0.30	(0.15)
1.7	0.17	(0.15)	0.22	(0.12)	0.10	(0.09)			0.35	(0.09)
2.0	0.31	(0.15)	0.38	(0.19)	0.35	(0.09)			0.43	(0.13)
2.5	0.73	(0.19)	0.09	(0.10)						
2.6	0.48	(0.13)	0.47	(0.09)	0.43	(0.14)			0.39	(0.07)
3.0			0.31	(0.10)	0.21	(0.06)				
3.7	0.34	(0.10)	0.14	(0.05)	0.82	(0.22)				
1890 m										
1.3			0.15	(0.13)	0.12	(0.16)	0.08	(0.14)	0.05	(0.13)
1.5			0.50	(0.16)	0.24	(0.12)			0.46	(0.12)
1.6			0.07	(0.15)						
2.0			0.14	(0.04)	0.08	(0.07)	0.15	(0.10)	0.12	(0.05)
2.1			0.52	(0.07)	0.15	(0.07)	0.33	(0.15)		
2.9			0.23	(0.08)	0.15	(0.04)	0.18	(0.08)	0.18	(0.06)
3.0			0.34	(0.06)	0.78	(0.07)	0.59	(0.12)	0.42	(0.06)
3050 m										
3030 m 1 5			0.11	(0.05)	0.33	(0.07)			0.17	(0,04)
1.5			0.11	(0.03)	0.55	(0.07)			0.17	(0.04)
1.0			0.50	(0.11)	0.20	(0.04)			0.02	(0.02)
2.3			0.20	(0.00)	0.20	(0.04)			0.20	(0.03)
∠.9			0.06	(0.05)	0.01	(0.04)			-0.02	(0.01)

Appendix 6. Diurnal courses of coarse root CO_2 release (R_R) from four selected root individuals measured during the dry season (D1) and during one of the more humid campaigns (H3) at the three study site at 1050, 1890 and 3050 m. Error bars represent 1SD. d = diameter.



Species	Family	n	A_{max}	N	C/N
1050 m					
indet	indet (Leguminoseae)	3	5.09 (0.73)	8.82 (1.24)	15.47
Lectandra cf	Lauraceae	5	6.31 (0.73)	9.18 (0.79)	15.78
<i>Miconia</i> sp	Melastomataceae	4	8.31 (2.49)	4.37 (0.36)	30.16
<i>Psammisa</i> cf	Ericaceae	4	3.73 (1.37)	2.57 (0.50)	53.52
<i>Hedyosmum</i> sp	Chlorantaceae	4	8.86 (2.35)	4.66 (0.34)	27.89
indet	Myrtaceae	4	6.58 (1.78)	4.45 (0.19)	29.78
Vismia sp	Clusiaceae	4	7.83 (1.08)	3.56 (0.58)	43.11
2000 m					
Cavendishia zamorensis	Ericaceae	3	8.47 (1.10)	3.43 (0.03)	46.24
Myrica pubescens	Myricaceae	3	6.43 (2.61)	5.01 (0.61)	26.37
indet	Asteraceae	3	4.28 (0.76)	4.66 (0.24)	29.42
3050 m					
Miconia	Melastomataceae	4	5.98 (0.85)	4.15 (0.13)	28.99
indet	Chlorantaceae	3	3.84 (1.43)	4.45 (0.52)	27.36
indet	Cunnoniaceae	1	1.76	3.26	38.01
Faramea	Rubiaceae	2	6.10 (0.24)	5.45 (0.43)	21.67
indet	indet	4	3.40 (1.81)	6.27 (2.65)	26.36
Clusia sp1	Clusiaceae	3	4.36 (1.84)	3.19 (0.34)	44.71
Axinea	Melastomataceae	4	5.45 (2.02)	3.77 (0.46)	37.87
indet	Melastomataceae	1	5.97	2.84	53.62
Cerotostema cf	Ericaceae	2	3.92 (1.82)	2.58 (0.06)	57.15
Styrax foveolaria	Stryracaceae	3	8.69 (2.70)	2.82 (0.38)	55.74
Cybianthus maginatus	Myrsinaceae	1	4.21	4.02	31.46

Appendix 7. Mean photosynthetic capacity (A_{max} , $\mu mol CO_2 m^{-2}$ leaf s⁻¹), leaf nitrogen content (mg g⁻¹) and C/N ratio of measured leaves nearby the study sites at 1050, 1890 m and within the site at 3050 m elevation (SD in parenthesis). n= number of leaves measured

Appendix 8. Light response curve of photosynthesis (A, μ mol CO₂ m⁻² leaf s⁻¹) at 3050 m. Given are mean values of light curves of 35 measured leaves (10 species). Error bars represent 1SD. PAR = photosynthetic active radiation.



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10. CURRICULUM VITAE

PERSONAL DETAILS

Date of birth Place of birth	21. July 1976 Coburg, Germany
Inationality	German
EDUCATION	
2005 - 2008	PhD study at the Department of Plant Ecology, Albrecht von Haller Institute for Plant Sciences, University of Göttingen, Germany
2002 - 2004	M.Sc. Program "Tropical and International Agriculture" at the University of Göttingen, Germany Degree obtained: M. Sc. agr.
2000 - 2002	Study of "Ecological Agriculture in the Tropics and Subtropics" at the University of Kassel, Germany Degree obtained: Dipl. Agr. Ing.
1997 – 2000	Study of Agricultural Sciences at the Christian Albrechts University of Kiel, Germany
1986 – 1996	Gymnasium Albertinum, Coburg, Germany High school graduation: Allgemeine Hochschulreife

WORK EXPERIENCE

2005 - 2007	PhD field work in Podocarpus National Park (Loja, Ecuador)
2003 - 2004	MSc. thesis in Santa Rosa (Argentina)
2001	Field assistance in Adana (Turkey)
1998 – 2000	Study-related (agricultural sciences) internships in South-Africa, Nicaragua and Germany
2005 – 2008	Scientific assistant at the Department of Plant Ecology, University of Göttingen, Germany
2003 - 2005	Graduate assistant at the Institute of Agricultural Chemistry, University of Göttingen, Germany
1998 – 1999	Graduate assistant at the Institute of Plant Nutrition and Soil Science, University of Kiel, Germany